

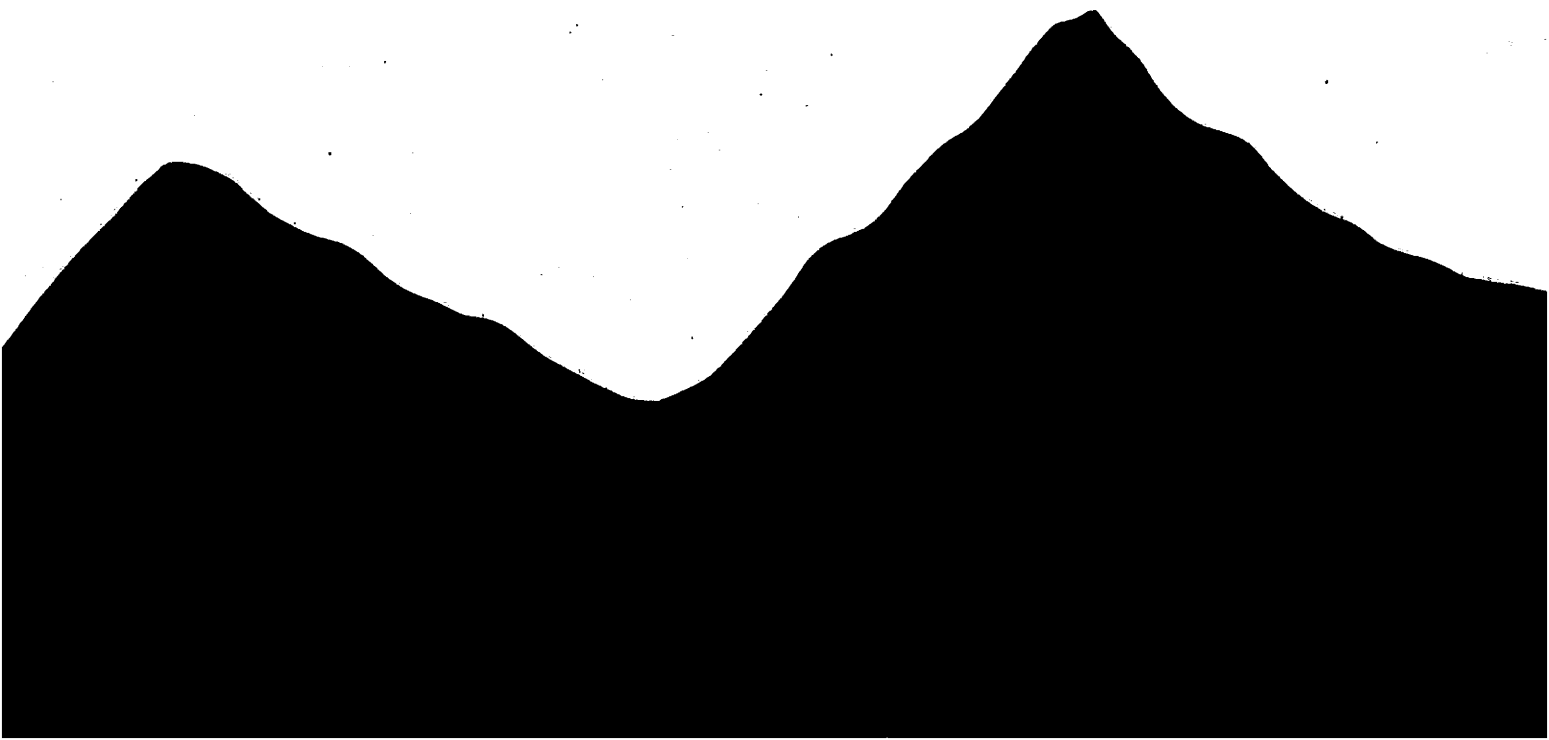
**DEGLACIAL CHRONOLOGY AND UPLIFT HISTORY:
NORTHEASTERN SECTOR, LAURENTIDE ICE SHEET**

Arthur S. Dyke

1974

Occasional Paper No. 12

INSTITUTE OF ARCTIC AND ALPINE RESEARCH • UNIVERSITY OF COLORADO



DEGLACIAL CHRONOLOGY AND UPLIFT HISTORY:

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by

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ABSTRACT

Using data contained in published literature and the Radiocarbon Data Bank of the Institute of Arctic and Alpine Research, an isochrone map is constructed which describes the pattern of deglaciation of the northeastern sector of the Laurentide Ice Sheet from the time of the late Wisconsin maximum (8000 yr to 8500 yr B.P.) to the present. The change in area and volume of the Northern Baffin Island Ice Cap from 7000 yr B.P. to the present is calculated using the isochrone map and two models relating ice area to volume. The volume measurements are then used to determine the contribution of the ice cap to world sea level rise since 7000 yr B.P.

Based on 325 shoreline locations, radiocarbon dated between 250 yr and 8750 yr B.P., eight isobase maps of the study area are produced depicting the amounts of uplift accomplished since 8000 yr B.P. and 1000 yr intervals thereafter. The pattern of isobases on the 8000 yr B.P. surface shows good agreement with the outline of the late Wisconsin terminal position. The 7000 yr B.P. and younger surfaces show the land recovering around five semi-independent uplift centers over the Queen Elizabeth Islands, Southampton Island, southeastern Keewatin, western Quebec and northern Baffin Island. These uplift centers correspond to late-glacial centers of retreat.

The change in geometry of the uplift surface over northern Baffin Island is analyzed in detail and is found to be a function of the rate of decay of the Northern Baffin Island Ice Cap. Graphs derived from manipulations of the isochrone and isobase maps describe this functional relationship.

Finally the duration and amount of residual rebound at the regional center of uplift over Southampton Island is predicted on the basis of the measured displacement of isobases between 8000 yr and 1000 yr B.P. It is suggested that rebound in that area will be complete within the next 4000 yr resulting in a further 48 m or less of uplift.

ACKNOWLEDGEMENTS

My interest in this topic was generated largely through discussions with Dr. J. T. Andrews, whose continued encouragement and assistance contributed greatly to this thesis. I thank Drs. J. D. Ives and J. C. Harrison for the careful reading and constructive criticism. This work was supported in part by grant number 20883 from the National Science Foundation. Funds for computer programming were provided by the Graduate School of the University of Colorado. Very frequent discussions with fellow graduate students, R. S. Bradley, J. A. Clark, J. H. England and G. H. Miller proved most beneficial.

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CHAPTER 1

INTRODUCTION

AREA AND OUTLINE OF STUDY

The primary aim of this thesis is to produce information on changing ice loads from the time of the late Wisconsin maximum to the present and the accompanying responses of the earth's crust to these changes. This is done through investigation of published reports on the glacial history and Quaternary deposits of the study area. It is anticipated that this information will be subsequently used in geophysical analysis of crustal and mantle rheology. Because the most valid test of a geophysical uplift model lies in the degree to which it predicts or reproduces an actual situation, it is necessary to provide a comprehensive (though interim) review and analysis of the field data. The northeastern sector of the area formerly covered by the Laurentide Ice Sheet was selected for this analysis because a greater amount of information, in the form of radiocarbon dates on ice marginal positions and relative sea-level stands (440 out of 835 dates for Laurentide North America), exists for this area than for any other.

A map of isochrones on the retreat of late Wisconsin ice is provided for the area between latitudes 62° and 74° N and longitudes 60° and 90° W. A series of isobase maps depicting the amounts of uplift accomplished since 8000 yr B.P. and successive 1000 yr intervals thereafter portray crustal response to the process and pattern of deglaciation. The boundaries of these maps are 58° and 76° N latitude and 60° and 100° W longitude. Methods of construction of the isochrone and isobase maps are discussed in chapters 2 and 3

respectively.

Further treatment of the data is largely restricted to the effect of late ice over Baffin Island on glacio-isostatic movements toward isostatic equilibrium in that area and to the prediction of the amount of residual recovery at an uplift center over Southampton Island.

PREVIOUS STUDIES

Isochrones have been drawn on the disintegration of late Wisconsin ice in North America by Bryson et al. (1969) and Prest (1969). Andrews (1970b, 1973a,b) and Walcott (1972) published maps of postglacial rebound for Laurentide North America but their small scale is not commensurate with the detailed treatment in this thesis. Otherwise only local isobases have been drawn (Løken 1962; Andrews, 1970a; Blake, 1970; Andrews et al., 1970). For the sake of brevity reference to individual field studies will be reserved for later relevant sections.

RADIOCARBON DATA BANK

This study was made feasible and greatly expedited by the existence at the University of Colorado Computing Center of a data bank containing information on dated material relating to changes in sea level in North America. It was compiled by J. T. Andrews, G. H. Miller and A. Street in 1971/72 as part of a project funded by the National Science Foundation (Grant 20883).

The radiocarbon data bank is one "set" of data within the TAXIR system and currently consists of 1044 items, 835 of which pertain to Laurentide North America. An estimated 150 dates have to be

added because of the continuing publication of dates in the issues of Radiocarbon (no. for 1973). Because the general TAXIR system has been described elsewhere (Brill, 1971) only a brief description of the radiocarbon data bank is given below.

The primary function of the data bank is to facilitate rapid information retrieval by the writing of simple programs, the body of which consists of a query or series of queries written in the "near English" language of boolean algebra. Data "items" within the bank are divided into 16 subsets, each defined by a "descriptor". Each descriptor, in turn, consists of a number of elements or "descriptor states". Table 1-I lists the 16 descriptors currently applied to the data bank and examples of corresponding descriptor states. The list, in fact, comprises one item in the response to the following query:

```
PRINT:(SAMPLE NO.,LAB NO.,LATITUDE,LONGITUDE),
(C14 DATE,STANDARD ERROR,MATERIAL DATED,INFERRED RELIABILITY OF DATE),
(UPLIFT,ELEVATION OF SAMPLE,INFERRED RELATIVE SEA LEVEL),
(ELEVATION OF LOCAL MARINE LIMIT,DATE ASSOCIATED WITH MARINE LIMIT),
(RELATION.TO ICE MARGIN,DISTANCE.FROM MAXIMUM ICE MARGIN),
(REFERENCE)
```

```
FOR ITEMS WITH LATITUDE, FROM 6200 TO 7400 AND LONGITUDE, FROM
6000 TO 9200*
```

The above query was designed to produce a printout of all information in the data bank relating to the area between the geographical coordinates specified (see Appendix for date list). The location of parentheses within the PRINT statement specifies the number of lines used in the response for each item. For example

TABLE 1-I

Descriptors of the Radiocarbon Data Bank with corresponding examples of descriptor states.

	Descriptor	Descriptor State
a	Sample no.	26
b	Lab no.	GSC1638 (Geological Survey of Canada)
c	Latitude	6756 (67°51' N. latitude)
d	Longitude	6603 (66° 03' W. longitude)
e	C14 date	8410 (14C yr B.P.)
f	Standard error	340 (yr.)
g	Material dated	SH (Shell)
h	Inferred reliability of date	Fairly Sure
i	Uplift	58 meters (=k+eustatic correction)
j	Elevation of sample	34 meters
k	Inferred relative sea level	37 meters
l	Elevation of local Marine limit	43 meters
m	Date associated with marine limit	No
n	Relation to ice margin	Ice nearby
o	Distance from maximum ice margin	---
p	Reference	Miller Up (unpublished)

the first line appearing in the response to the above query will read:

```
26      GSC1638      6756      6603
```

Following lines of descriptor states of the same item will be successively justified 4 columns to the left.

The first descriptor appearing in the query determines the ordering of the response. The response generated by the above statement will list items in order of increasing "sample no." size. Thus, if the items in the response are to be printed in order of increasing age, the descriptor, C14 DATE, must be placed first.

The last line of the query beginning with the boolean expression FOR is the limiting statement on the size of the response. However, when used in conjunction with the expression SAME, the response can be expanded and divided into useful classes. An example would be:

```
PRINT:(C14 DATE,LATITUDE,LONGITUDE),
```

```
(REFERENCE)
```

```
FOR ITEMS WITH C14 DATE FROM 7750 TO 8250
```

```
PRINT:SAME FOR ITEMS WITH C14 DATE FROM 7250 TO 7750
```

```
PRINT:SAME....
```

```
PRINT:SAME FOR ITEMS WITH C14 DATE FROM 250 TO 750*
```

This query will generate a printout of the descriptor states of the specified descriptors listed in order of the magnitude of the radio-carbon date for successively younger 500 year intervals.

The preceding description of the TAXIR radiocarbon data bank is by no means complete and is merely meant to illustrate the application of the system to the problem at hand and its potential as a time saving tool.

CHAPTER 2

CHRONOLOGY OF GROWTH AND DECAY OF THE ICE SHEET
OVER FOXE BASIN AND BAFFIN ISLAND

LAST INTERGLACIAL AND EARLY WISCONSIN EVENTS

Known deposits ante-dating the late Wisconsin are predominantly restricted to the outer coastal areas of Baffin Island and Northern Labrador (Fig. 2-1). These deposits and their relation to the late Wisconsin ice margin are discussed in the following section and elsewhere (Andrews and Ives, 1972; Andrews et al., 1972a,b); Miller and Dyke, 1974).

Terasmae et al. (1966) provide the only detailed description of pre-Wisconsin terrestrial interglacial deposits from north-central Baffin Island. Along the upper reaches of the Isortoq River, near the northern end of the Barnes Ice Cap, occur several outcrops of plant-bearing beds overlain by a thin deposit of till. Organic materials from these beds have been radiocarbon-dated > 40,000 yr B.P. and the contained fossil assemblage indicates that prior to their deposition the climate was warmer than present, probably similar to the present climate of southern Baffin Island. The beds have, therefore, been assigned an interglacial age, to which has been applied the name Flitaway Interglacial (Andrews, 1968). It has not yet been possible to definitely correlate them with deposits from the southern margins of the Ice Sheet.

Nappe structures within the beds have a westward orientation suggesting that they were overridden by ice from the east. On this basis, it has been argued that the Laurentide Ice Sheet in this area grew somewhere to the east of these outcrops and flowed toward

Foxe Basin where it coalesced with ice flowing in the opposite direction from Melville Peninsula and northeastern Keewatin. Subsequently, an ice dome became established over Foxe Basin with accompanying radial outflow (Ives and Andrews, 1963). That ice flowed eastward and northeastward over Baffin Island and westward over Melville Peninsula is supported by an analysis of the carbonate content of the fine fractions of 36 till samples from these areas (Andrews and Sim, 1964).

Shell fragments have been collected from till in central-northern Baffin Island (Andrews, 1968) and along the southeastern coast (Blake, 1966). These shells were transported by ice flowing from Foxe Basin and through Hudson Strait, respectively. Reliable dating of these shelly tills may prove very valuable in establishing a chronology of the initial growth of the Laurentide Ice Sheet about which very little is presently known. It is worthy of note, however, that most attempts to develop geophysical models of late- and post-glacial rebound proceed on the assumption that equilibrium depression had been achieved prior to the onset of unloading (cf. Brothie and Sylvester, 1969). Before this assumption can be confirmed or rejected it is necessary that the early history of the ice sheet be elucidated.

LATE WISCONSIN MAXIMUM

As early as the late 1800's geomorphologists were outlining evidence for a restricted late Wisconsin ice cover on western Greenland (Chamberlin, 1895) and northern Labrador (Bell, 1895; Low, 1896; Daly, 1902; Coleman, 1920). More recently a three-fold

vertical weathering and glacial-morphological sequence has been outlined for northern Labrador (Ives, 1957; Løken, 1962).

Mercer (1956) was the first to present evidence of the existence of ice-free areas on Baffin Island during the late Wisconsin maximum. Based on differential degrees of weathering of landforms and the distribution of felsenmeer and partially submerged cirques over Meta Incognita Peninsula (referred to by Mercer as Kingait Peninsula), he concluded that much of the northeastern side of the peninsula and the entire southeast coast of Frobisher Bay seaward of Cape Lawrence remained outside the limits of late Wisconsin ice during the last glacial maximum (Mercer, 1956, p. 557).

This was followed by Løken's (1966) description of old tills and glaciomarine deposits dated > 54,000 yr B.P. from the Cape Aston and Cape Christian forelands. These deposits relate to a sea level about 60 m higher than the postglacial marine limit.

Subsequent dating of "old" shells from Henry Kater peninsula (King, 1969) and Remote Peninsula (Ives and Buckley, 1969) and plant material from the Bruce Mountains near Cape Adair (Harrison, 1964) led to the placing of the outermost portions of 5 forelands along the eastern Baffin coast outside the late Wisconsin ice margin on the Glacial Map of Canada (Prest et al., 1967) and Prest's (1969) isochrone map (see also Prest, 1970, p. 697).

Since publication of these isochrone maps, intensive research on Northern Cumberland Peninsula (Pheasant and Andrews, 1972, 1973; Andrews et al., 1972; England and Andrews, 1973; Boyer and Pheasant, 1974) has led to the formulation of a three-fold weathering zonation very similar to that of northern Labrador with

which it has been tentatively correlated (Pheasant and Andrews, 1972). During the late Wisconsin maximum ice advanced only to the vicinity of the fiord heads in northern Cumberland Peninsula (with the exception of Okoa Bay) and remained at its maximum along the local equivalents of the Cockburn Moraines (Andrews and Ives, 1972) until about 8000 yr B.P. Such evidence, along with (1) the increasing number of "old" radiometric dates on marine shells collected from their growth positions (See Fig. 2-1), and (2) the near synchronicity of the marine limits along the 70 to 80 km lengths of eastern Baffin fiords, led Andrews and Ives (1972) to suggest that the Cockburn Moraines may mark the maximum late Wisconsin ice stand over all of eastern Baffin Island. Miller and Dyke (1974) in a more expanded discussion arrive at the same conclusion and present a detailed outline of the late Wisconsin ice border. That outline appears here as the 8000 - 8500 yr B.P. isochrone with the area expanded to include northern Baffin Island. The outermost isochrone is not considered to mark the maximum ice stand south of Bernier Bay, although ^{14}C dates > 23,000 yr B.P. on shells from Boothia Peninsula and Somerset Island may well indicate that much of the Gulf of Boothia was free from glacier ice during the last maximum.

Compton (1964) found shells in marine deposits at 270 m (880 ft) above sea level on northern Brodeur Peninsula and speculated that these high shorelines may relate to a period of isostatic depression older than late Wisconsin, and that, therefore, the degree of preservation of the strandlines meant that the area had not been glaciated during the last glacial maximum. Craig (1965) on the basis of the interlobate appearance of the moraines to the north of

Bernier Bay suggested that Brodeur Peninsula carried an independent ice cap during the closing phases of the late Wisconsin. However, if Compton's (1964) interpretation is correct, ice on the Peninsula did not reach sea level in all areas.

CONSTRUCTION OF THE ISOCHRONE MAP (Fig. 2-1)

Three categories of information were used to pinpoint ice frontal positions in both space and time prior to construction of the isochrone map; namely: (a) location and trend of geomorphological features, (b) lichen diameter isophyses, and (c) finite radiocarbon dates.

Major end moraines and ice-marginal drainage channels along with isophyses on the maximum diameters of the lichen, Rhizocarpon geographicum, appearing in published reports were plotted on a single map at an original scale of 1:1,000,000. The R. geographicum isophyses were assigned dates based on a growth rate curve constructed for the region near the northern end of the Barnes Ice Cap (Andrews and Webber, 1964; Løken and Andrews, 1966). Geomorphological features were dated according to their association with radiocarbon dates on terrestrial organic material or marine bivalves, or by interpolation assuming a roughly uniform rate of retreat.

The radiocarbon dates were divided into three classes according to their description in the TAXIR data bank (Chapter 1): (a) ice contact, (b) ice nearby, and (c) ice far away. Dates described as "ice far away" were used only where there exists a scarcity of "ice contact" dates. A total of 314 dates are available for the area.

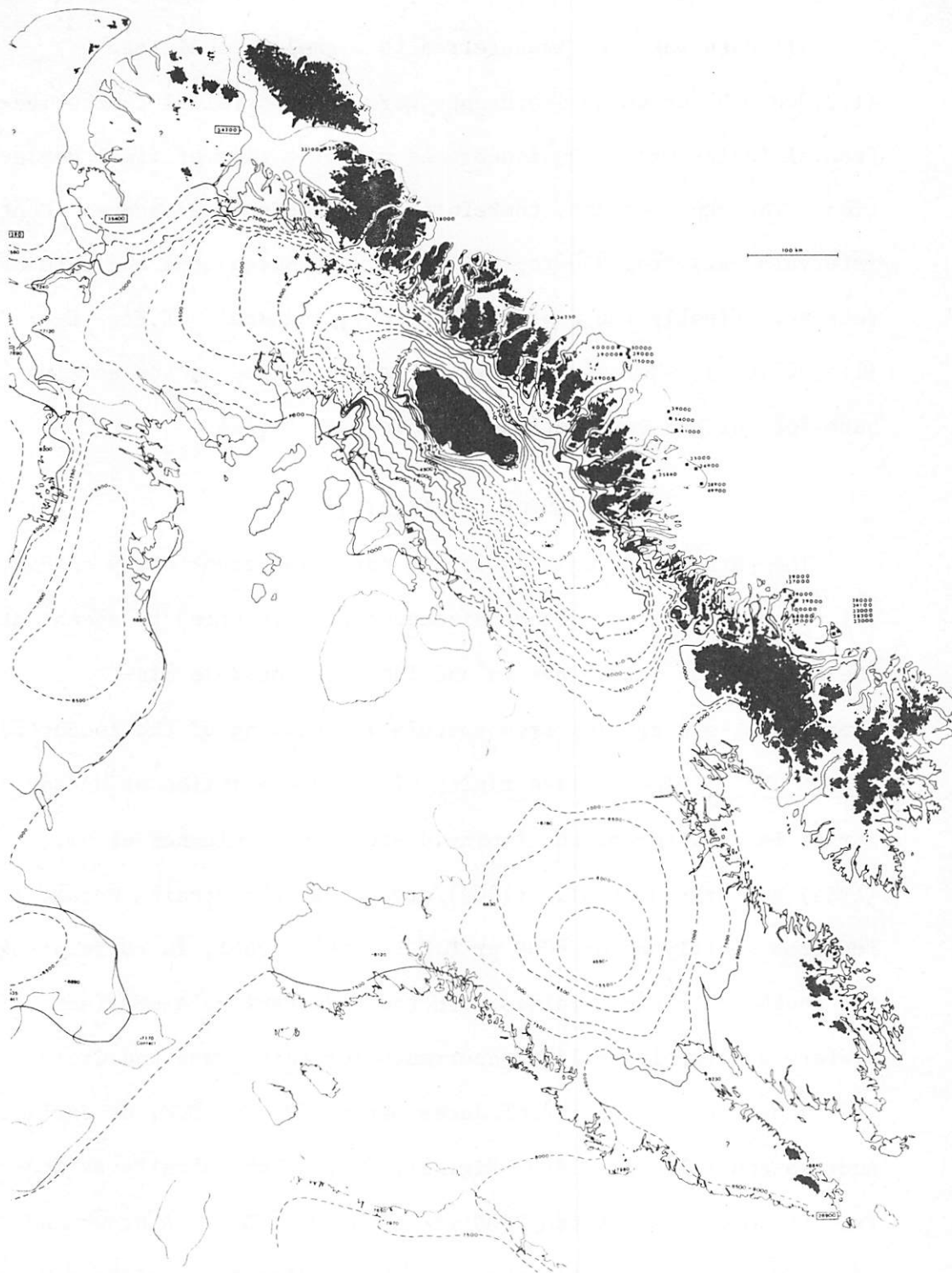


Figure 2-1: Isochrones on the deglaciation of the northeastern sector of the Laurentide Ice Sheet. Contour interval varies from 250 yr to 1000 yr depending upon the rate of retreat and the degree of morphological and chronological control. The locations of ice contact dates from post-glacial deposits (dots) are shown where space permits. The locations along with minimum ages of raised marine deposits which are associated with pre-late Wisconsin ice coverage are shown by solid squares. Dates enclosed by rectangles may or may not be associated with ice transported marine shells.

All data was then transferred to a smaller scale map (1:2,000,000) on which isochrones were drawn parallel to the ice-frontal indicators. The isochrones date the time of final deglaciation. The map does not, therefore, distinguish readvances. Contour intervals vary from 250 to 1000 years, depending upon the rate of retreat. Finally the map was reduced to the scale of Fig. 2-1 (1:5,000,000), whereas the original scale served as the working base for further computations used in this thesis.

PATTERN OF DEGLACIATION

The pattern of deglaciation is portrayed graphically on Fig. 2-1. The following regional discussion is presented to accentuate the main chronological events and identify possible misrepresentations arising from speculative placing of the isochrones.

HUDSON STRAIT The precise timing of the deglaciation of Hudson Strait is a matter of continuing discussion. Falconer et al., (1965) and Bryson et al., (1969) argue that the straits became ice-free shortly after 8000 yr B.P. Blake (1966), in correlating the southern Baffin moraines with the Foxe Peninsula moraines prefers a slightly earlier occurrence for this event and Craig (1965) places the opening of James Bay at 7900 yr B.P. or earlier. Andrews and Falconer (1969, Fig. 11) portray the Straits as ice-free at some time between 7800 yr and 8200 yr B.P. Andrews and Ives (1972) conclude that the available radiometric evidence does not permit definitive statements concerning the exact timing of deglaciation.

For the purposes of this paper the exact timing is not of

vital importance, as the argument is concerned with the interpretation of only a few hundred years. It is sufficient to note that ice recession in outermost Hudson Strait began about 8800 yr B.P. (Blake, 1966), permitting emergence of the York Sound (Frobisher Bay) delta sequence since that time, and that the entire strait and most of Hudson Bay had been deglaciated by 7500 yr B.P. On Fig. 2-1 the 8000 yr B.P. isochrone is somewhat arbitrarily drawn in this region and shows much of the Strait free of ice at that time.

FOXES BASIN Following deglaciation of Hudson Strait the ice retreated rapidly south and southwestward, clearing the whole of the Tyrell Sea (the expanded postglacial Hudson Bay) within a few centuries.

The marine transgression into Foxe Basin was delayed until about 500 to 800 yr later. Dates on the marine limit in Foxe Basin range from 6830 ± 150 yr B.P. (GSC-465) in Bowman Bay, southern Baffin Island, to 6880 ± 180 yr B.P. (GSC-291) on eastern Melville Peninsula and 7170 ± 90 yr B.P. (Gx-1068) on southeastern Southampton Island. These dates indicate that most of Foxe Basin had been inundated by the sea by about 6800 yr to 7000 yr B.P. Perhaps the last area to become ice-free was northern Steensby Inlet where the marine limit has been dated at 6500 yr B.P. from an extrapolated uplift curve (Ives, 1964).

The time of separation of ice over Melville Peninsula from that over northern Baffin Island is problematical as no dates are reported from either side of Fury and Hecla Strait. The strait is, however, shown ice-free at 7000 yr B.P. on the basis of a date of 7120 ± 140 yr B.P. (GSC-307) from the head of Agu Bay.

An ice contact date of 6535 ± 115 (GSC-308) yr B.P. on south-

central Southampton implies the existence of relatively late ice there, probably in the form of a small ice cap landward of the limit of the postglacial marine transgression (Bird, 1972).

SOUTHERN BAFFIN ISLAND Blake (1966) concluded on the basis of the latest ice flow pattern and orientation of eskers that the region near the northeastern shore of Amadjuak Lake had been a center of dispersal during the closing phases of the late Wisconsin. It is proposed here that the Southern Baffin Island Ice Cap became separated from the Northern Baffin Island Ice Cap through development of a calving bay which penetrated from Foxe Basin to the head of Cumberland Sound during the early stages of the marine transgression into Foxe Basin about 7000 yr B.P. This is substantiated by a date of 6760 ± 140 yr B.P. (GSC-466) on shells dating the marine limit southeast of Nettiling Lake.

A ^{14}C date of 4550 ± 220 yr B.P. (GSC-498) on basal peat from a site near the southeastern corner of Amadjuak Lake provides a minimum date for the final disintegration of the ice cap.

CUMBERLAND PENINSULA Cumberland Peninsula today contains the 6000 km^2 Penny Ice Cap and numerous other glaciers and smaller highland ice caps. During the late Wisconsin maximum, Laurentide ice was confluent with Penny ice on the western part of the Peninsula but ice cover at that time was not much more extensive than at present (Andrews and Dugdale, 1971; England and Andrews, 1973; Pheasant and Andrews, 1972, 1973; Miller and Dyke, 1974), and the Penny Ice Cap remained the dominant center of dispersal (Andrews et al., 1970). The glaciers of eastern Cumberland Peninsula were probably only slightly expanded from their present state.

Ice retreat on eastern Baffin Island began about 8000 to 8500 yr B.P. and the Penny Ice Cap probably separated from the northern Baffin Island Ice Cap about 1000 yr later. Following this the Cumberland Peninsula glaciers experienced a number of (Neoglacial) advances and retreats (Miller, 1973).

NORTHERN AND CENTRAL BAFFIN ISLAND If the proposed dates of separation of the Penny, southern Baffin Island and Melville Peninsula ice caps are correct, the northern Baffin Island Ice Cap came into existence as a separate entity about 7000 yr B.P. From that time it retreated in a more or less orderly fashion to the present margins of the Barnes Ice Cap in roughly the manner described by Ives and Andrews (1963). The total recession of the northeastern margin has been only 45 to 50 km, an average rate of retreat of about 6 m yr^{-1} . This calculation is based on the assumption that the Cockburn Moraines which are marked by the 8000 yr - 8500 yr B.P. isochrone (Fig. 2-1) were deposited during the late Wisconsin maximum. On Henry Kater Peninsula King (1969) has described fossiliferous marine deposits which have been dated at $10,210 \pm 180 \text{ yr B.P.}$ (Y-1986), and Andrews et al. (1970) report a date of $9180 \pm 180 \text{ yr B.P.}$ (Y-1832) on shells dating the marine limit at Cape Hooper. These two dates are among the few from all of eastern Baffin Island which are in excess of 8500 yr B.P. (Andrews and Ives, 1972, Fig. 3) and their precise meaning is not clear. Conceivably they may indicate an earlier date of deglaciation in the Home Bay area but the data is not sufficient to permit construction of isochrones for that time.

The detailed field studies and geomorphological mapping of Ives and Andrews (1963), Sim (1963), Andrews (1963 a and b), Ives (1964 a and b), Andrews and Webber (1964), Andrews (1966), Løken and Andrews (1966), King and Buckley (1967), Andrews (1968, 1970a), and Andrews et al., (1970) provide detailed control on the chronology and pattern of recession of the Northern Baffin Island Ice Cap along its eastern and western margins and in the regions near the Barnes Ice Cap. There is little dating control in large tracts of central Baffin Island and the isochrones there have been drawn largely by interpolation between moraine segments. The pattern outlined by Fig. 2-1 will be radically wrong, however, only if the glacier retreated to more than one center. There is no evidence to date to suggest that this was the case.

COMPARISON WITH OTHER ISOCHRONE MAPS

The major differences between Fig. 2-1 and the isochrone maps of Prest (1969) and Bryson et al., (1969) are enumerated below:

(1) Both Prest and Bryson et al., have drawn ice marginal positions for times prior to 8500 yr B.P., and thereby suggest that late Wisconsin Laurentide ice extended beyond the outer coast in all but a few regions. Figure 2-1 carries the implication that areas beyond the Cockburn Moraines, with the exception of those highlands which may have carried local glacier ice, were ice free during the late Wisconsin maximum. The location of marine deposits along with their ages which lie outside the Cockburn Moraines and which are associated with mid or early Wisconsin glacial episodes are shown in support of this hypothesis.

(2) The retreat of ice over north-central Baffin Island is shown in a very generalized manner by Prest and Bryson et al. Here an attempt is made to more precisely portray the chronology and pattern of deglaciation and the rate of unloading.

(3) Prest's map shows ice retreating northward over Southampton Island, whereas on the maps of Bryson et al., and this thesis that Island is considered to have been a center of retreat during the late glacial (ca. 6500 yr to 6000 yr B.P.).

TEMPORAL CHANGES IN GLACIATED AREA AND ICE VOLUMES

Several writers have used various methods to calculate the volumes of ice and water contained in Quaternary ice sheets. Paterson (1972) reviewed earlier attempts and used the "speculative" isochrone map of Prest (1969) to calculate changes in area and volume of the Laurentide and Innuitian ice sheets from 18,000 yr to 6000 yr B.P.

This present thesis presents area measurements and volume calculations for the Northern Baffin Island Ice Cap from 7000 yr B.P. to the present at 1000 yr time intervals. Although the first two results overlap in time with those of Paterson (1972, Table 5), they are significantly different due to the discrepancies between the two base isochrone maps.

CALCULATION METHODS

Using the isochrone map described earlier in this chapter (scale 1:2,000,000) and a polar planimeter the areas of the Northern Baffin Island Ice Cap were measured for 1000 yr intervals from 7000 yr B.P. to the present. Each area measurement was performed twice

and the mean of the two values was chosen as final. The small discrepancies between the two measurements indicate that the final values are accurate within the confines of the map scale (estimated precision $\pm 3\%$).

The ice volumes were then calculated as functions of area using the following formulae:

$$V = 7.854 (R^2 (R^{\frac{1}{2}})) \quad , \quad (2-1)$$

where R is the radius (in m) of the ice sheet and V is volume, and

$$\log_{10} V = 1.23 (\log_{10} S - 1) \quad , \quad (2-2)$$

where V is volume and S is area.

Equation 2-1 results from integrating Hollin's (1962) approximation of an ice cap profile (Antarctica) for circular ice caps. Andrews (1970b) pointed out that use of this formula for an elongated ice cap probably results in an over-estimation of volume. The value of R used is the radius of a circle whose area is equal to the area of the ice cap as determined by planimetry.

Equation 2-2 is the regression line of a log-log plot of volume against area for six existing ice sheets (including Antarctica, Greenland and the Barnes ice caps). The standard error of deviation from the line is 0.1 and the volume calculations are probably accurate to within 12% (Paterson, 1972).

RESULTS

Table 2-I and Fig. 2-2 summarize the results of the area measurements and volume calculations described above. It will be noted that the volumes calculated by equation 2-1 are consistently higher than those calculated by equation 2-2, but that the difference

TABLE 2-I

Change in area and volume of the Northern Baffin Island
Ice Cap from 7000 yr B.P. to the present.

A	B	C	D	E
Time (Yr.B.P.)	Area 10^3 km^2	Volume $(v=7.854(R^2 R^{\frac{1}{2}}))$ (10^3 km^3)	Volume $(\log S=1.23(\log S-1))$ (10^3 km^3)	Diff- erence (%)
7000	140.5	161.80	133.15	17.7
6000	81.0	82.20	62.40	24.1
5000	57.0	52.14	42.15	20.2
4000	37.96	31.95	26.30	17.7
3000	20.88	18.35	12.89	29.8
2000	12.04	7.315	6.53	10.8
1000	7.44	3.921	3.65	6.9
0	6.0	3.170	2.799	11.7

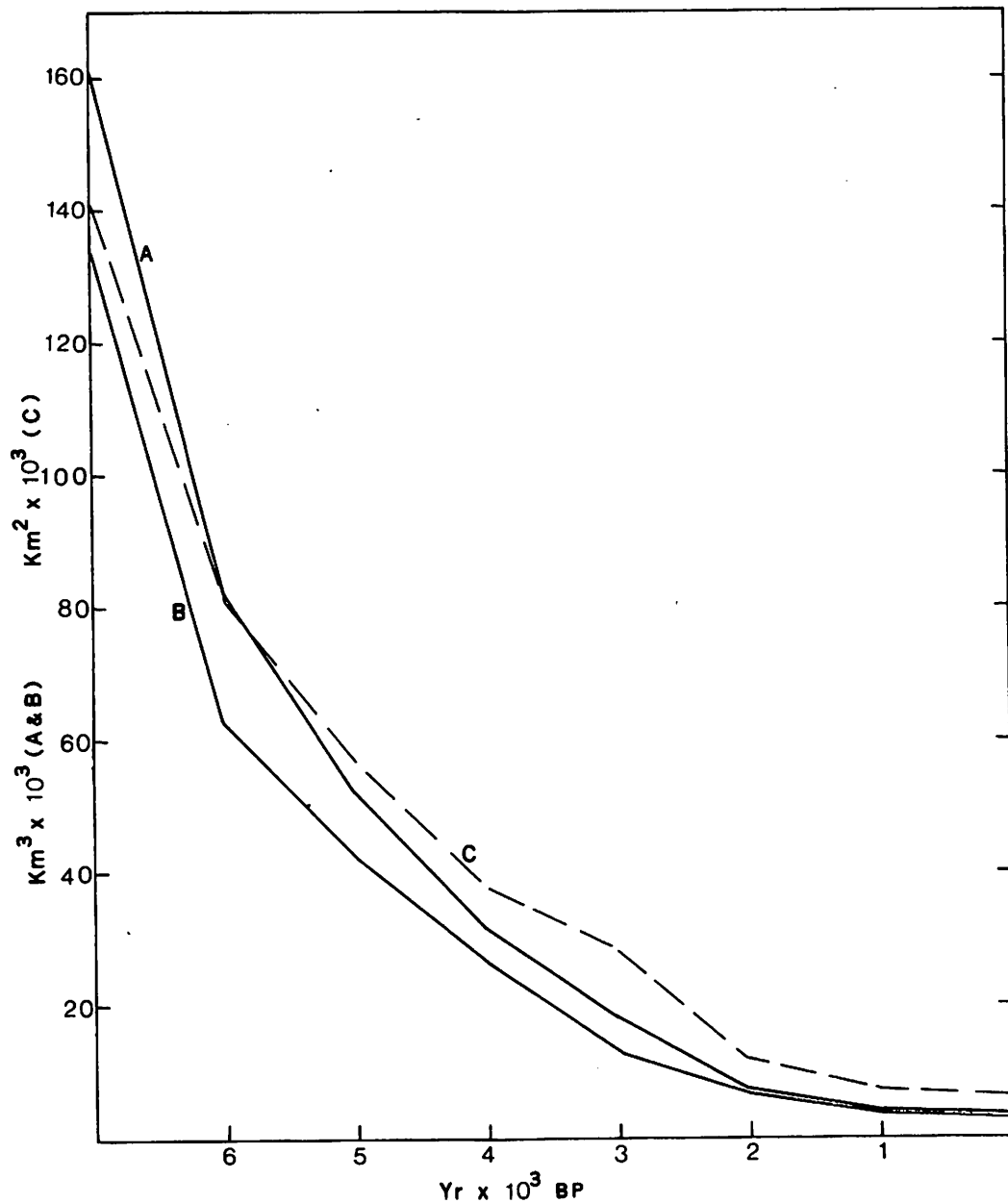


Figure 2-2: Area (by planimeter) and volume (by the two methods discussed in the text) of the Northern Baffin Island Ice Cap from 7000 yr B.P. to the present.

generally diminishes as does the area. Equation 2-1, thus, probably results in an over-estimation of the volumes of elongated ice caps as mentioned previously.

To determine the relative reliability of the two sets of results, the volume of the northern Baffin Island Ice Cap at 5000 yr B.P. was calculated by a different method. The irregular outline of the 5000 yr B.P. isochrone was smoothed to approximate an ellipse. The transverse profile of an elliptically shaped ice cap can be described by the following equation:

$$h = (2h_o x)^{\frac{1}{2}} \quad (2-3)$$

where h is the height at distance x from the ice edge and $h_o = k(pg)^{-1}$, where k is basal shear stress, p is the density of ice, and g is acceleration due to gravity. If $k = 1$ bar, $h_o = 11.3$ m, and

$$h = (23x)^{\frac{1}{2}} \quad (2-4)$$

(Nye, 1952).

From the profile equation the cross-sectional areas under a number of transverse profiles across the ice cap were calculated and the volume derived by integrating. The volume thus calculated for the ice cap at 5000 yr B.P. is $41 \text{ km}^3 \times 10^3$, which differs from the value in column D by only 2.6%. This suggests that the values in column D are the more reliable.

Table 2-II lists the contribution to world wide sea level rise which resulted from melting of the Northern Baffin Island Ice Sheet. The total melting that has occurred since 7000 yr B.P. contributed less than 40 cm to eustatic sea level. This leads to the realiza-

TABLE 2-II

Contribution to world sea level rise by the Northern Baffin
Island Ice Cap (7000 yr B.P. to present)

A	B($\times 10^{14} \text{ m}^3$)	C
Time (Yr.B.P.)	Water content of Northern Baffin Ice Sheet range of values from Col. C & D or Table 2-I	Contribution (in meters) to world sea level rise since time stated in Column A
7000	1.456 - 1.1984	0.3954 - 0.3250
6000	0.7398 - 0.5616	0.1970 - 0.1486
5000	0.469 - 0.3794	0.1220 - 0.0981
4000	0.2876 - 0.2367	0.0718 - 0.0586
3000	0.1652 - 0.1160	0.0379 - 0.0252
2000	0.0658 - 0.0588	0.0103 - 0.0093
1000	0.0353 - 0.03285	0.0019 - 0.0021
0	0.0285 - 0.0252	0 - 0

B = Ice volume \times 0.9

C = Water content at time, t, - Water content of Barnes
Ice Cap \div Area of Ocean ($3.61 \times 10^{14} \text{ m}^2$)

tion that an incipient Laurentide Ice Sheet may reach considerable proportions before its growth would be readily recognizable in the world's marine geological record.

CHAPTER 3

LATE-GLACIAL, POSTGLACIAL AND RESIDUAL REBOUND AND THE RELATION OF REBOUND TO CHANGES IN LOAD

GENERAL STATEMENT

In this chapter, eight isobase maps depicting the amounts of rebound accomplished since 8000 yr B.P. and successive 1000 yr intervals thereafter are presented for the area between latitudes 58° and 76° N and longitudes 60° and 100° W. The uplift pattern as it changes through time is largely controlled by the pattern of deglaciation from the late Wisconsin maximum (see Chapter 2). This relationship is discussed in some detail with particular attention paid to the effect of late ice over Baffin Island in the form of the slowly retreating Northern Baffin Island Ice Cap. Following that, an attempt is made to estimate the amount and duration of residual rebound at the Southampton Island uplift center based on the rate of displacement of the various isobases on the maps toward this center.

CONSTRUCTION OF THE ISOBASE MAPS

For the map area the radiocarbon data bank contains 325 finite dates between 250 yr and 8750 yr B.P. associated with raised marine deposits. These dates were divided into 500 yr class intervals and then subdivided according to the response to the query, "inferred reliability of date", into the following categories:*

* A nominal classification of the probability that the quoted date accurately places in time a strandline of known elevation.

(a) definite, (b) likely, fairly sure, (c) possibly and (d) small chance.

Table 3-I shows the number of dates in each category for each of the 500 yr class intervals. Dates whose inferred reliability was "small chance" were ignored. However, as can be seen from the table, this did not seriously deplete the amount of dating control available.

Technical problems associated with radiocarbon dating of fossil marine organisms are potentially important in a study such as this. However, no adjustments to the data have been made here as no widely accepted conventions regarding such corrections have yet become established. These problems are discussed in detail by Olsson and Blake (1962), Stuiver and Suess (1966) and Mangerud (1972).

The amount of uplift accomplished since the formation of a dated strandline is

$$U_t = E_s + E_t \quad (3-1)$$

where U_t is the amount of uplift that has occurred since time t , E_s is the elevation of a strandline to which a dated sample refers, and E_t is the amount of eustatic sea level rise since time t (considered positive). The value of E_t used here is that supplied by the sea level curve of Mörner (1969).

The values of uplift were plotted on maps with a scale of 1:2,000,000 to allow accurate placing of the data points and careful interpretation of local patterns. Each map utilized data from the nearest of the three class intervals (Table 3-I) in order to extend control to as much of the study area as possible and to provide upper and lower limits on the true values of uplift as well as nearest approximations. This is justified because the class

TABLE 3-I. Inferred reliability of radiocarbon dates for 500 yr class intervals used in construction of the isobase caps.

Date class intervals (yr B.P.x10 ³)	Definite		Fairly sure likely		Possibly		Small chance		Total in class I(Σ_{class})	$\frac{\Sigma_{\text{class}}}{\Sigma_{\text{Dates}}}$ J
	No.	%	No.	%	No.	%	No.	%		
	A	B	C	D	E	F	G	H		
8.25-8.75	6	23.1	14	53.8	1	3.8	5	19.2	26	8.0
7.75-8.25*	16	42.1	17	44.7	3	7.9	2	5.3	38	11.7
7.25-7.75	9	36.0	13	52.0	3	12.0			25	7.7
6.75-7.25*	13	30.9	24	57.1	3	7.1	2	4.7	42	12.9
6.25-6.75	9	32.1	15	53.6	3	10.7	1	3.5	28	8.6
5.75-6.25*	7	36.8	11	57.9	1	5.3			19	5.8
5.25-5.75	5	26.3	13	68.4	1	5.3			19	5.8
4.75-5.25*	7	30.4	13	56.5	3	13.0			23	7.1
4.25-4.75	2	20.1	5	50.0	3	30.0			10	3.1
3.75-4.25*	4	25.0	11	68.8			1	6.3	16	4.9
3.25-3.75	4	20.0	15	75.0			1	5.0	20	6.2
2.75-3.25*	4	28.6	9	64.3			1	7.1	14	4.3
2.25-2.75	2	25.0	6	75.0					8	2.5
1.75-2.25*	2	18.2	7	63.6	1	9.1	1	9.1	11	3.4
1.25-1.75	3	60.0	1	20.0	1	20.0			5	1.5
0.75-1.25*	3	17.6	14	82.4					17	5.2
0.25-0.75	1	25.0	3	75.0					4	1.2
Total in Category and Mean Percentile	97	29.8	191	58.8	23	7.1	14	4.3	325	
	Σ	\bar{X}	Σ	\bar{X}	Σ	\bar{X}	Σ	\bar{X}	Σ_{Dates}	

*Periods for which isobases are provided.

interval of 500 yr is approximately equal to the 2 standard deviation error range on either side of the mean quoted value (± 250 yr) of most samples (Appendix). The isobases were then drawn by interpolation so that the upper and lower limiting values fell near isobases of lesser and greater magnitudes, respectively, and the data providing the nearest approximations of the true value of uplift (i.e., data from the median class interval) exerted the greatest amount of control. Figure 3-1 is presented to illustrate this aspect of the construction method.

If isobases for a given time were extended over areas which had not yet been deglaciated (were still ice-covered), it was assumed that no major breaks in slope occurred on the uplift surface. Furthermore it was assumed that a profile along such a surface was similar to but more subdued than a theoretical ice sheet profile. In such cases useful data were supplied from the youngest of the three class intervals for areas that became deglaciated within 500 yr after the point in time represented by the map.

The patterns presented by this series of maps are derived solely from the interpretation of sound field data rather than upon empirically or theoretically derived models. Indeed these are the patterns which geophysical models must explain.

The 1:2,000,000 scale served as the base for further calculations and measurements. The maps were then reduced to a smaller scale for convenience of presentation. For ease of interpretation the land masses appear as featureless blocked plateaus and the isobases as superimposed topographies. The locations of the data points are shown on each map by differing symbols according to the

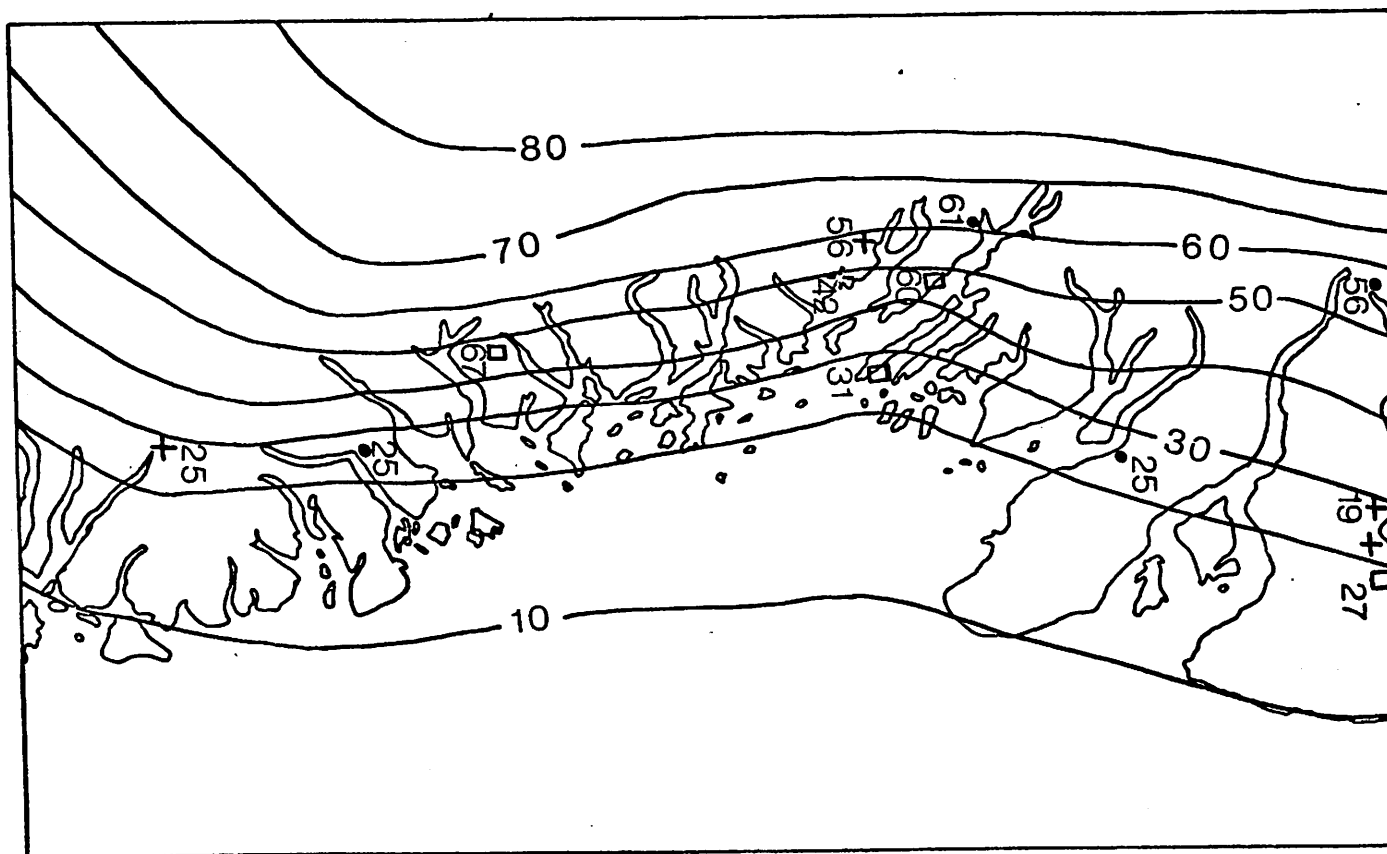


Figure 3-1: A portion of the 7000 yr B. P. uplift surface (Home Bay Area) showing the distribution of data and the method of isobase construction. Solid dots represent locations of strandlines dated between 6750 yr and 7250 yr B.P. Open squares represent locations of strandlines dated between 7250 yr B.P. and 7750 yr B.P. Uplift accomplished since 7000 yr B.P. is, therefore, less than the stated values (in meters). Crosses mark locations of strandlines dated between 6250 yr and 6750 yr B.P. Uplift accomplished since 7000 B.P. is more than the stated value. Scale 1:2,000,000. Contour interval: 10 meters.

class interval from which they were chosen. Due to scale reduction these points are not discernable where the isobases are closely spaced.

It is worth noting that although the study area contains the densest distribution of dated samples of any area of comparable size in North America, the sample points are not randomly distributed. For significant areas, especially Ungava Bay, Hall Peninsula, southern and eastern Cumberland Peninsula, Bylot Island, Borden Peninsula and Brodeur Peninsula, as well as much of Keewatin no dating control has yet been provided. Therefore, considering the large scale of the original maps, future field research and new dates may require that corrections of greater or lesser significance be applied on these areas of the map sheets. Nevertheless, the writer feels that the time is ripe for, and a sufficiently strong body of data exists upon which can be based, a useful analysis. This analysis itself will hopefully point to future research problems.

TEMPORAL CHANGES IN THE FORM OF THE UPLIFT SURFACE

THE 8000 YR B.P. SURFACE (Fig 3-2)

The first in this series of maps depicts the amount of rebound that has been accomplished since 8000 yr B.P., which in several recent papers (Chapter 2) is the proposed date of the late Wisconsin maximum for a large part of the study area.

Because the general outline of the margin of the ice sheet should be reflected, though smoothed, by the isobases on uplift accomplished since onset of deglaciation, and because the outline of the depressed region should continue to generally reflect this limit until several thousand years later, a comparison of the

8000

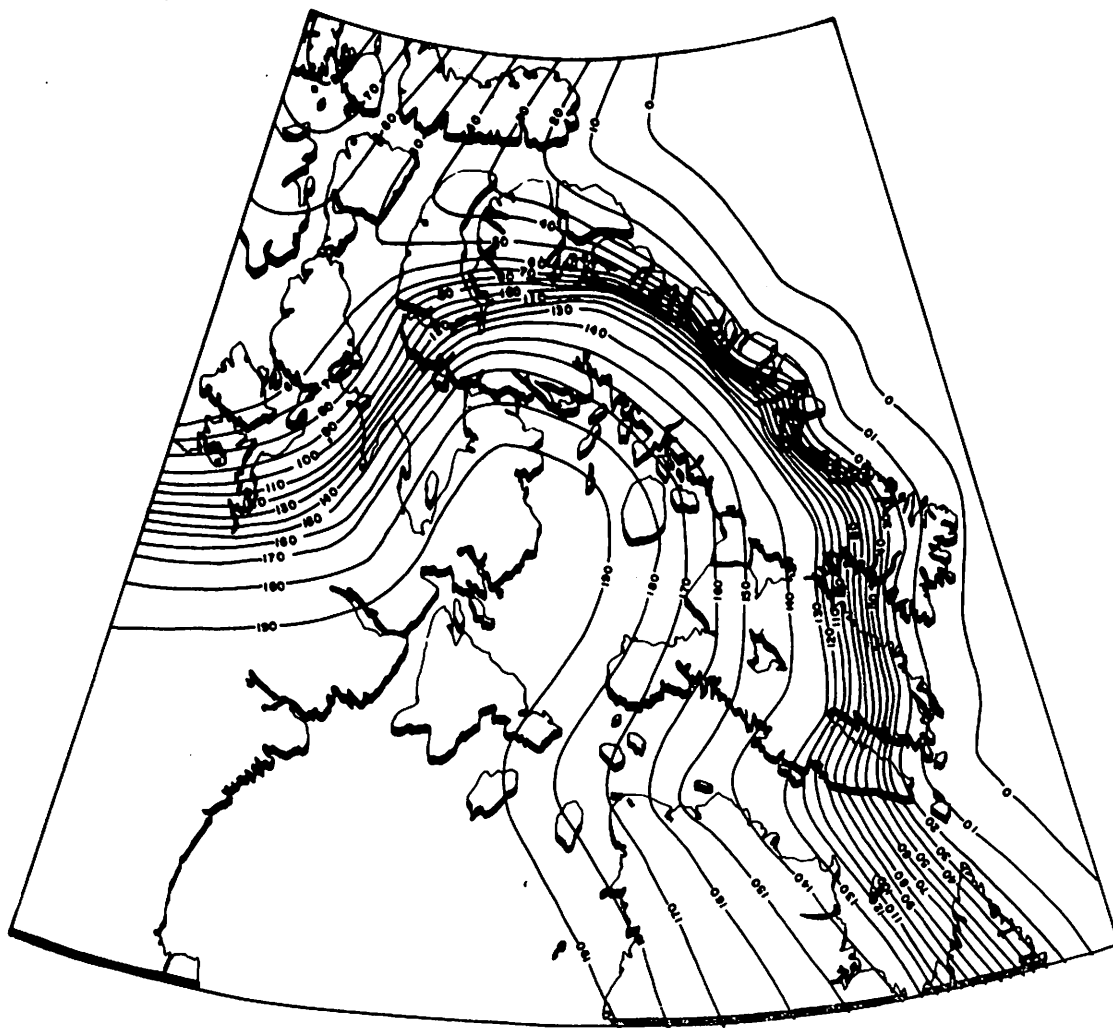


Figure 3-2: Isobases on the amount of uplift accomplished since 8000 yr B.P. Contour interval is 10 meters. Dots mark shoreline locations dated between 7750 yr and 8250 yr B.P. Open squares mark shoreline locations dated between 8250 yr and 8750 yr B.P. Crosses mark shoreline locations dated between 7250 yr and 7750 yr B.P. Some locations are obscured by the contour density.

proposed maximum extent of ice and the isobase pattern on Fig. 3-2 is warranted.

In general the 8000 yr B.P. isochrone (Fig. 2-1) roughly parallels the outermost isobases and is nearly coincident with the 90m and 100m lines. The area over Cumberland Peninsula provides an exception to this which may be an indication that the isobases are incorrectly placed there. Alternately, the low values of uplift and the non-parallelism of the isobases and the ice margin over this peninsula reflects the small amount of glacier retreat (and thus, change in load) that the area has experienced since the last glacial maximum. The peninsula, therefore, is responding to the much greater change in load over areas farther west.

The pronounced embayment of the isobases over Meta Incognita Peninsula and the even more pronounced embayment over northern Baffin Island and Lancaster Sound argue strongly for the hypothesis that the 8000-8500 yr B.P. isochrone marks the maximum ice frontal position. If ice had extended over all or most of Baffin Island to the margins of the continental shelf and then rapidly melted back to the position marked by the 8000 yr B.P. isochrone, one would expect that a great deal of rebound would have remained unaccomplished over the Continental Shelf and that the outline of this extensive ice sheet would be reflected in the isobase pattern. Such is not the case.

The trend of the isobase trough extending southwestward from outer Lancaster Sound roughly follows the 200m isobath as far as southern Somerset Island. During the eustatic minimum sea level was depressed about 120m (Steinen et al., 1973). However, the land

surface was also considerably depressed, so that the potential depth of water in Lancaster Sound and Prince Rupert inlet during the late Wisconsin maximum was still nearly 200m (500m in eastern Lancaster Sound as far west as Northern Brodeur Peninsula). This great depth of water likely provided a very effective calving environment which tended to restrict the coverage of glacier ice. Such was also likely the case in Cumberland Sound and Frobisher Bay.

It is proposed that the major depression in the isobases across the northern part of the map sheet roughly divides the areas of dominant influence of the Laurentide and Innuitian (Blake, 1970) ice sheets. This carries with it the implication that Laurentide ice did not extend nearly as far north in northeastern Keewatin as was previously assumed and the Innuitian Ice Sheet extended its influence over at least Somerset and Prince of Wales islands, considerably south of Barrow Strait. The lower gradient on the isobase surface on the northwestern portion of the map suggests that the Innuitian Ice Sheet was much thinner than its southern counterpart. This latter conclusion was also reached by Moran and Bryson (1969) from an independent approach.

One final notable area of coincidence between the isobase trends and the 8000 yr B.P. isochrone is the region between Ikerbi-lung and Sam Ford fiords where a minor seaward protrusion of the ice front produces a fairly noticeable wave in the isobases.

The pattern of contours on Smoothed Free Air gravity anomalies (Walcott, 1970) extending from north-eastern Labrador to Devon Island and Boothia Peninsula is strikingly similar to the isobase

trends on Fig. 3-2. The +20 milligal contour is very nearly coincident with the 8000 yr B.P. 20m isobase and the zero milligal contour lies very close to the zero isobase at 1000 yr B.P. (Fig. 3-3).

In consideration of both lines of evidence it seems most unlikely that the peninsulas of eastern and northern Baffin Island were overrun by Laurentide ice during the last glacial maximum.

THE 7000 YR TO 1000 YR B.P. SURFACES

The isobases on the amount of uplift accomplished since 7000 yr B.P. (Fig. 3-4) present a significantly different pattern from that on Fig. 3-2, while thereafter (Figs. 3-5 to 3-10) the pattern changes very little. Although the isobases near the margin of the depressed region maintain trends similar to that on the 8000 yr B.P. surface (that continue to reflect the shape of the maximum ice stand), by 7000 yr B.P. the land is recovering around five semi-independent centers over the Queen Elizabeth Islands, northern Baffin Island, Southampton Island, southeastern Keewatin and western Quebec. This detail is different from the smooth patterns shown by Andrews (1970b) and Walcott (1972) but shows fair agreement with the uplift centers determined by Andrews and Barnett (1972) based on the locations of interceptions of orthoginals to local isobases.

The change in the pattern of uplift from 8000 yr to 7000 yr B.P. is very closely related to the major deglaciation events that occurred between these dates. The embayment along Hudson Strait is more pronounced, extending well into Hudson Bay by 7000 yr B.P. The uplift centers all relate to centers of retreat, providing

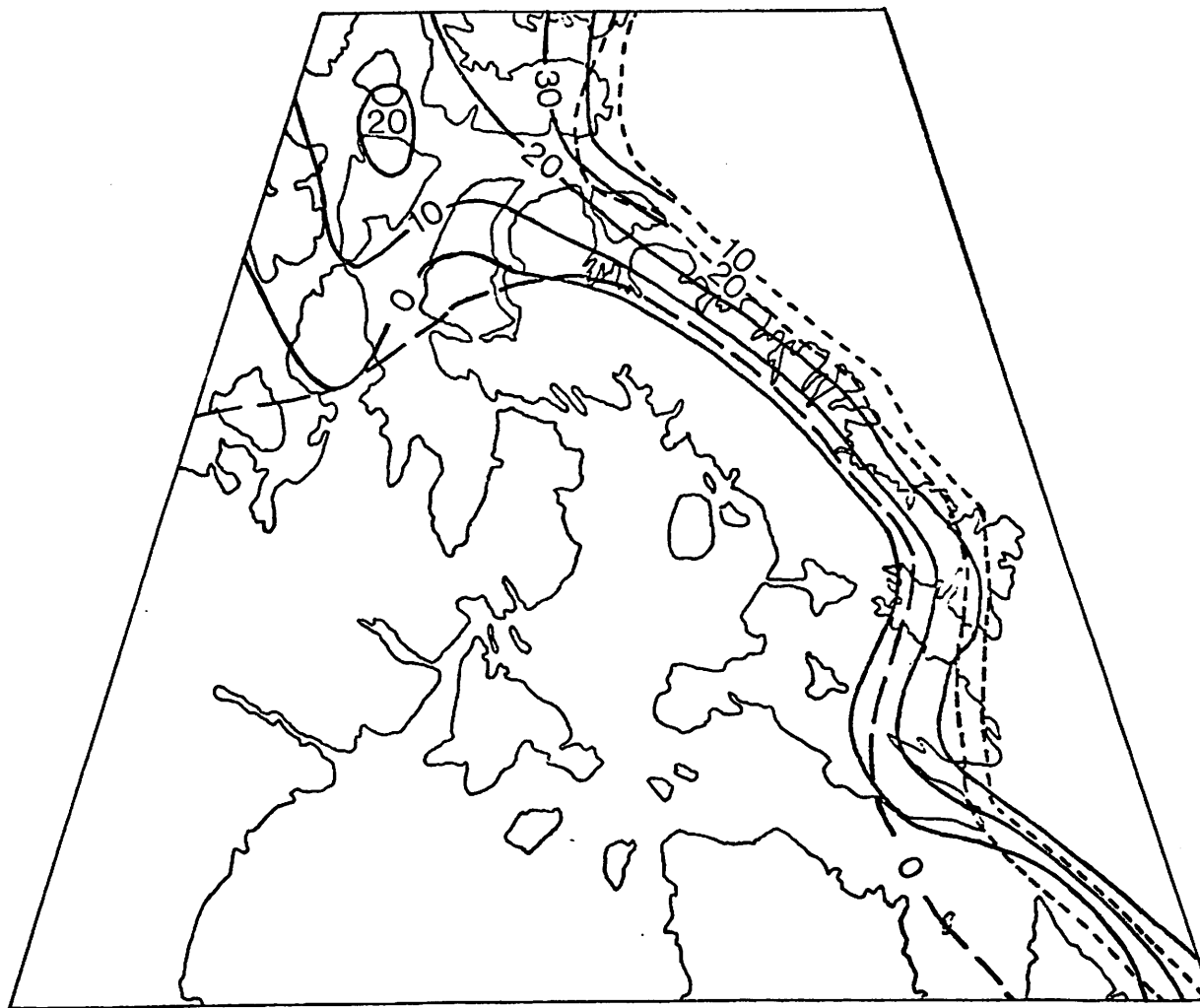


Figure 3-3: A comparison between isobase patterns and the pattern of contours on smoothed Free Air gravity anomalies. Solid lines are gravity anomaly contours in milligals. The short-dashed lines are the 10 m and 20 m isobases from the 8000 yr B.P. surface. The long-dashed line is the 0m isobase from the 1000 yr B.P. surface.

7000

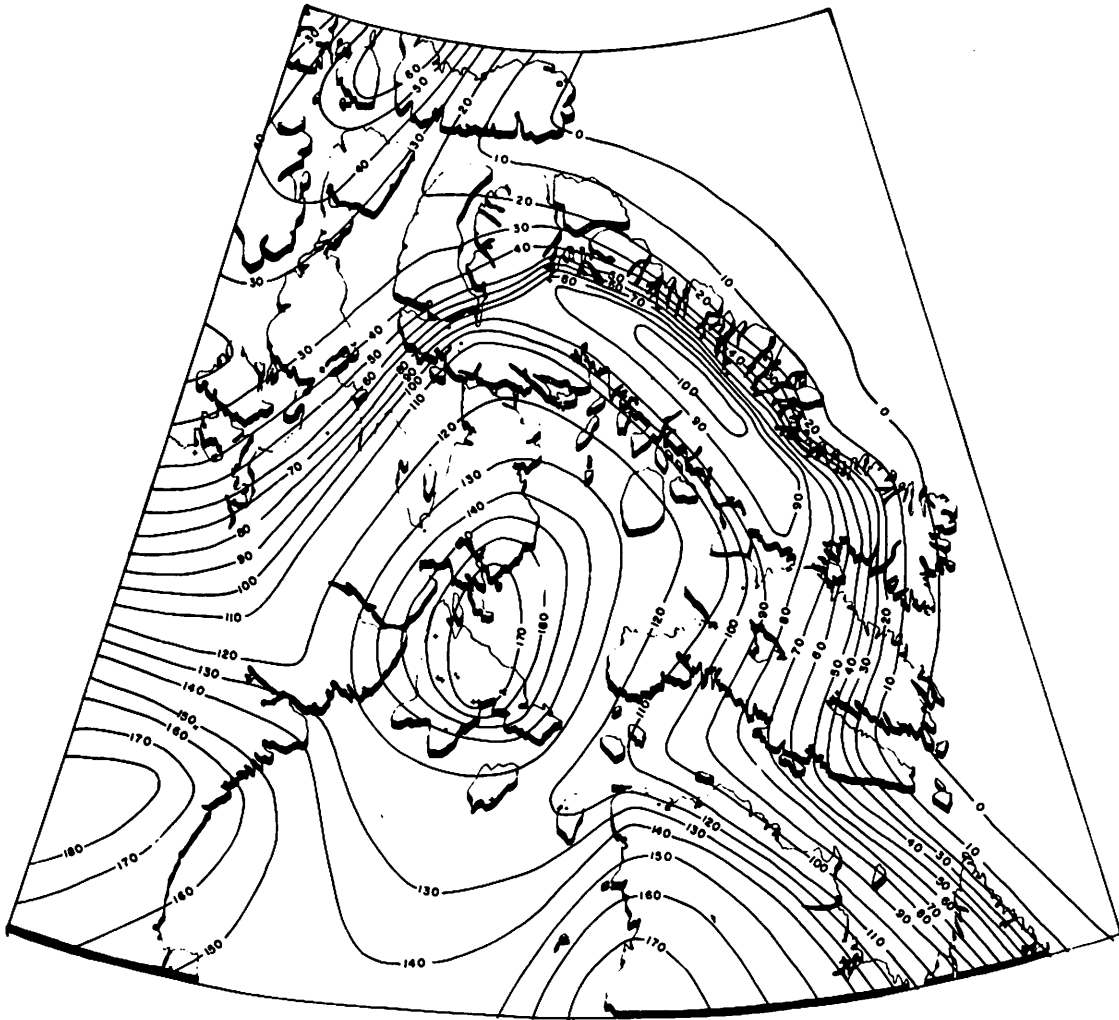


Figure 3-4: Isobases on the amount of uplift accomplished since 7000 yr B.P. Contour interval is 10 m. Dots mark shorelines dated between 6750 yr and 7250 yr B.P. Open squares mark shorelines dated between 7250 yr and 7750 yr B.P. Crosses mark shorelines dated between 6250 yr and 6750 yr B.P.

6000

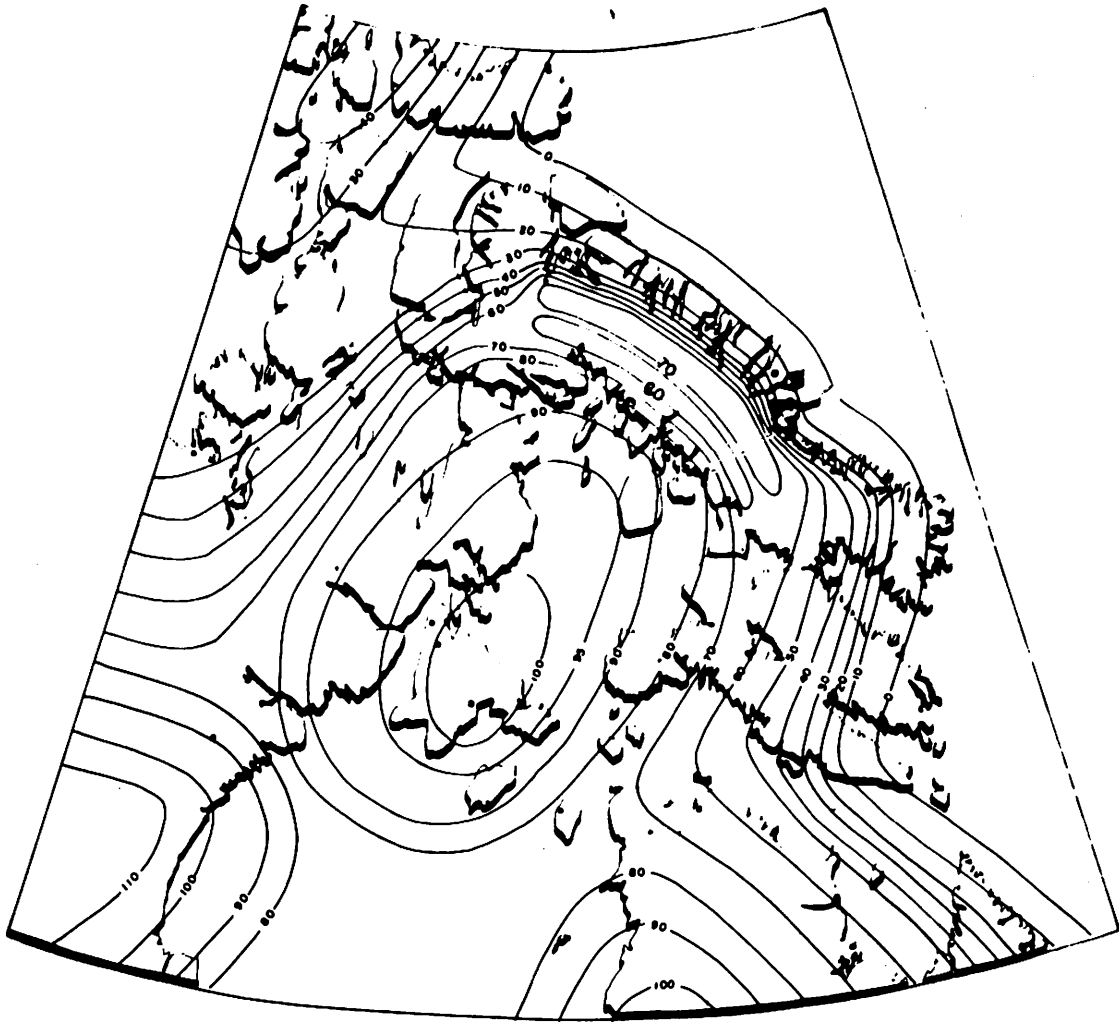


Figure 3-5: Isobases on the amount of uplift accomplished since 6000 yr B.P. Contour interval is 5 meters and 10 meters. Dots mark shorelines dated between 5750 yr and 6250 yr B.P. Open squares mark shorelines dated between 6250 yr and 6750 yr B.P. Crosses mark shorelines dated between 5250 yr and 5750 yr B.P.

5000

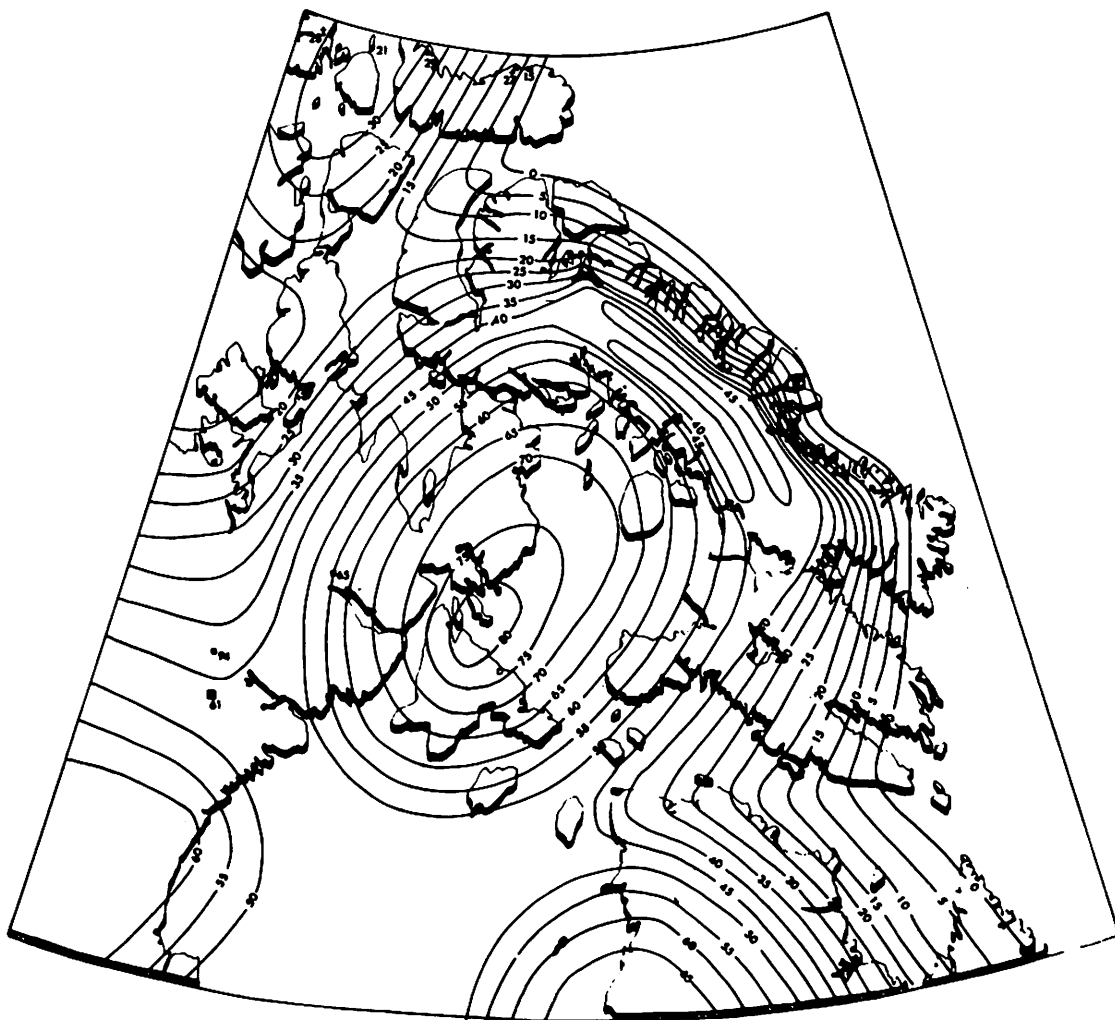


Figure 3-6: Isobases on the amount of uplift accomplished since 5000 yr B.P. Contour interval is 5 meters. Dots mark shorelines dated between 4750 yr and 5250 yr B.P. Open squares mark shorelines dated between 5250 yr and 5750 yr B.P. Crosses mark shoreline dated between 4250 yr and 4750 yr B.P.

4000

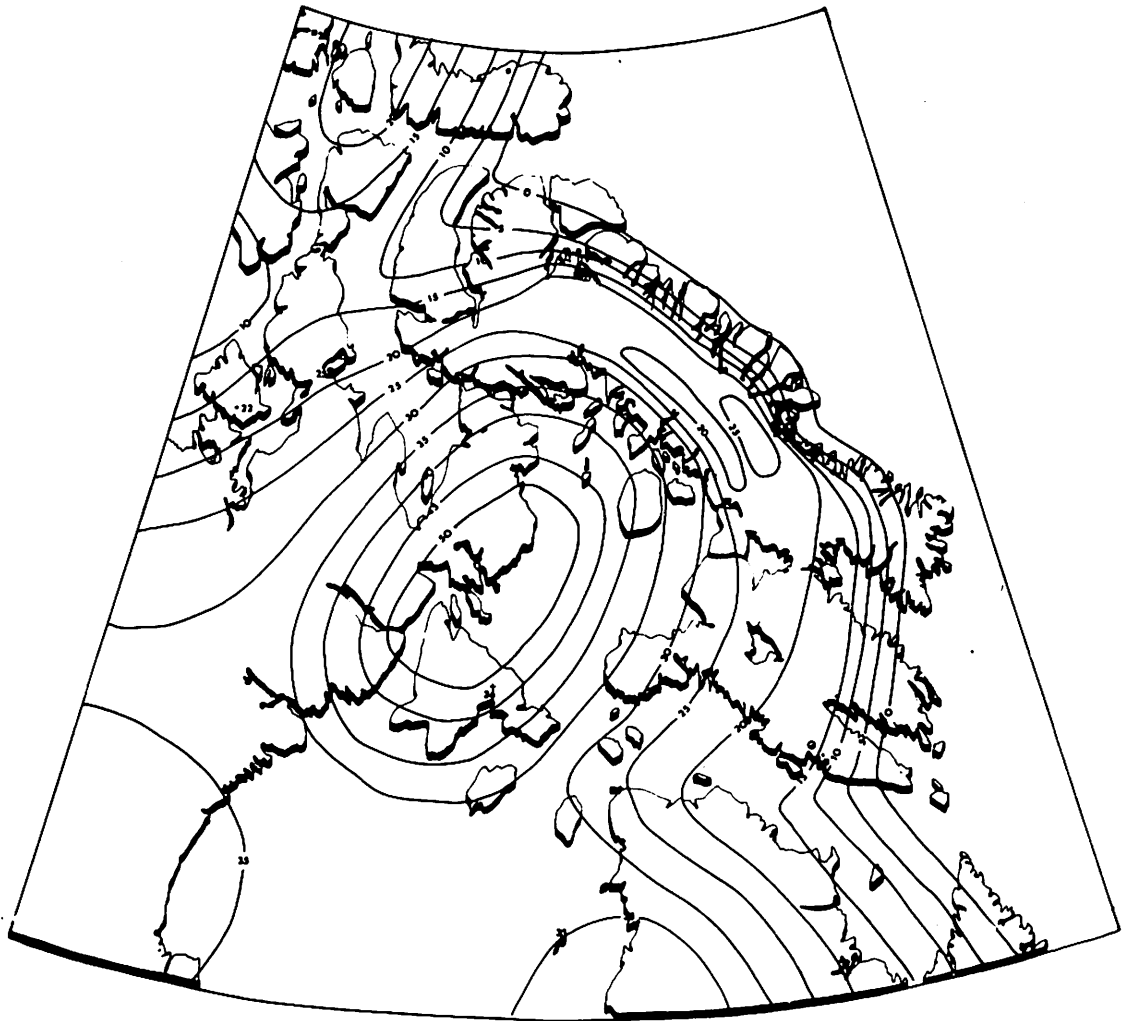


Figure 3-7: Isobases on the amount of uplift accomplished since 4000 yr B.P. Contour interval is 5 meters. Dots mark shorelines dated between 3750 yr and 4250 yr B.P. Open squares mark shorelines dated between 4250 yr and 4750 yr B.P. Crosses mark shorelines dated between 3250 yr and 3750 yr B.P.

3000

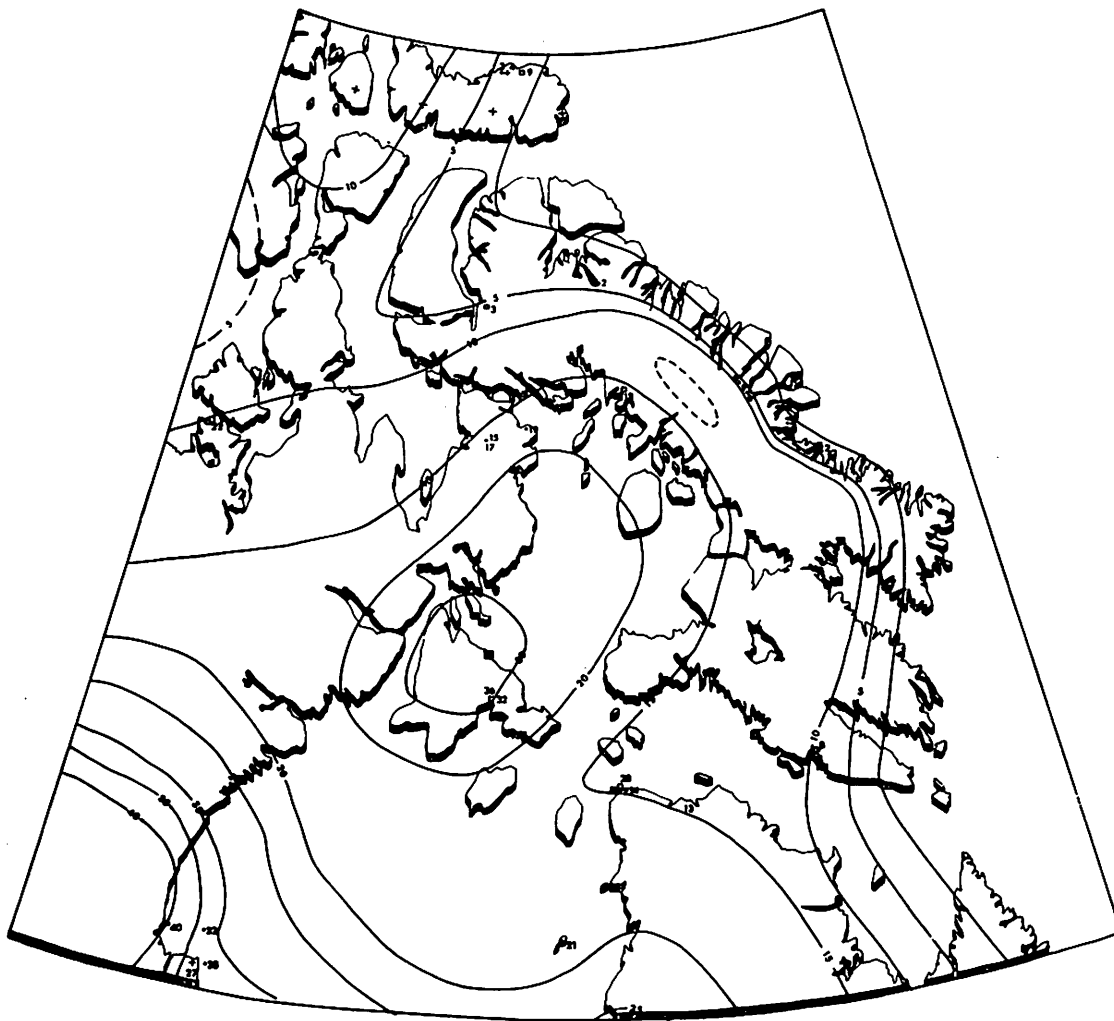


Figure 3-8: Isobases on the amount of uplift accomplished since 3000 yr B.P. Contour interval is 5 meters. Dots mark shorelines dated between 2750 yr and 3250 yr B. P. Open squares mark shorelines dated between 3250 yr and 3750 yr B.P. Crosses mark shorelines dated between 2250 yr and 2750 yr B.P.

2000

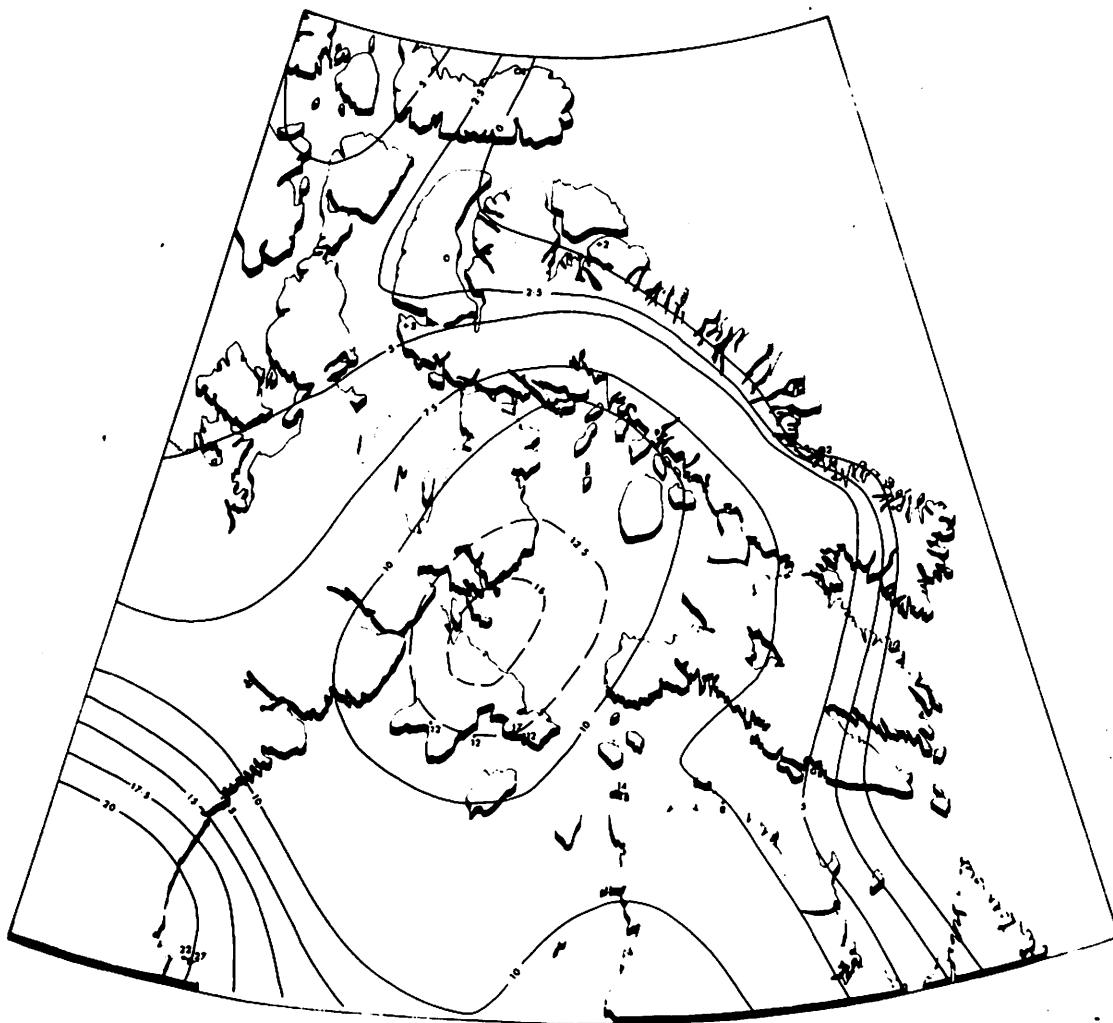


Figure 3-9: Isobases on the amount of uplift accomplished since 2000 yr B.P. Contour interval is 2.5 meters. Dots mark shorelines dated between 1750 yr and 2250 yr B.P. Open squares mark shorelines dated between 2250 yr and 2750 yr B.P. Crosses mark shorelines dated between 1250 yr and 1750 yr B.P.

1000

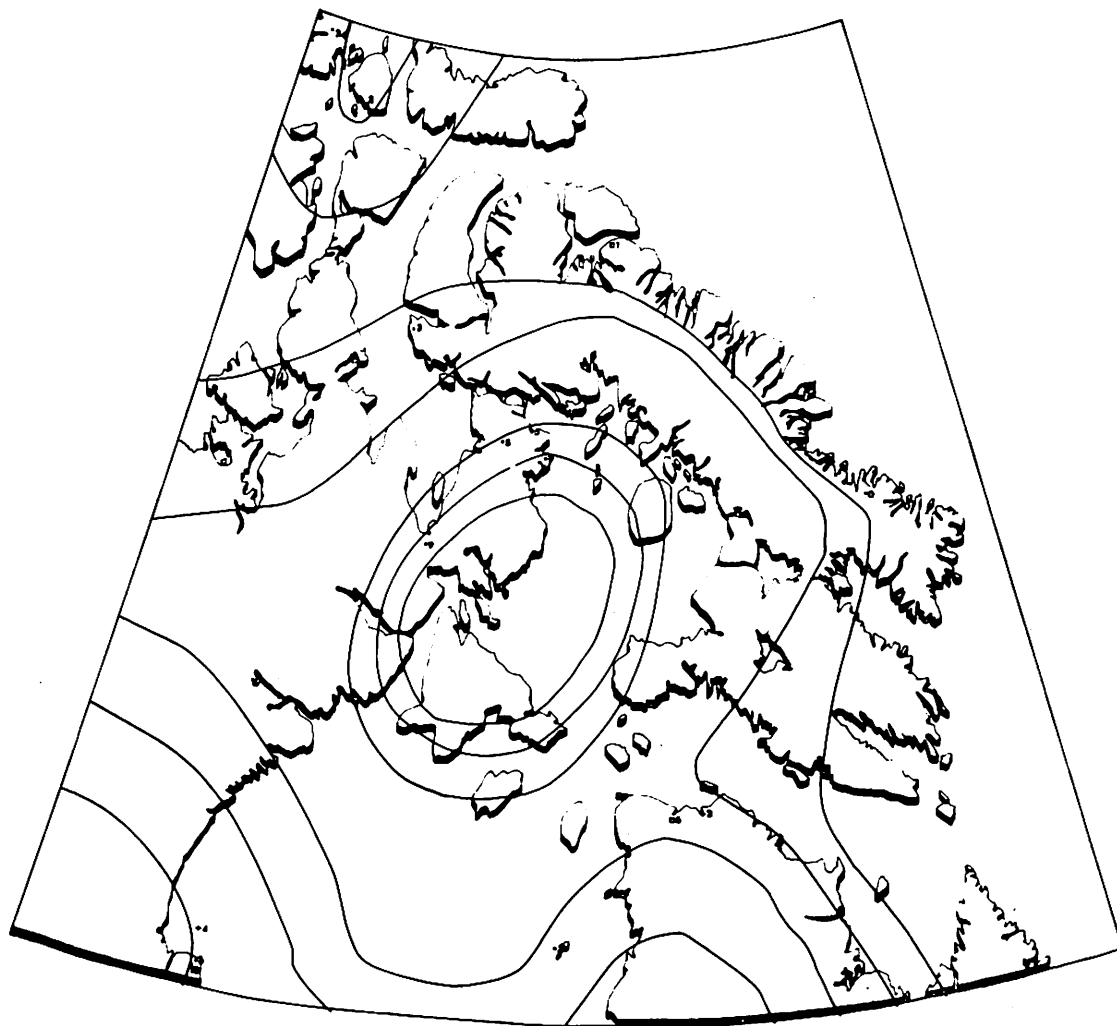


Figure 3-10: Isobases on the amount of uplift accomplished since 1000 yr B.P. Contour interval is 2.5 meters. Dots mark shorelines dated between 750 yr and 1250 yr B.P. Open squares mark shorelines dated between 1250 yr and 1750 yr B.P. Crosses mark shorelines dated between 250 yr and 750 yr B.P.

evidence that the major remnants into which the Laurentide Ice Sheet broke following the catastrophic marine transgressions into Hudson Bay and Foxe Basin acted effectively in restraining the rates of rebound of the underlying mantle. On the other hand, the extremely rapid unloading of great thicknesses of ice from Hudson Bay and Foxe Basin allowed high rates of uplift to proceed unchecked in these areas immediately adjacent to the remnant ice masses. The result of the combination of these two opposing effects upon adjacent areas was to establish and accentuate the various centers of late-glacial and postglacial uplift.

From 7000 yr B.P. to 1000 yr B.P. the amplitudes of the domes on the uplift surface decreases exponentially as the overlying loads of ice decreases.

THE EFFECT OF THE NORTHERN BAFFIN ISLAND ICE CAP UPON THE REBOUND OF BAFFIN ISLAND

EQUIDISTANT DIAGRAMS

Andrews (1970b, Fig. 7-6) demonstrated that synchronous strandlines from the west and east coasts of Baffin Island do not, when projected onto an equidistant diagram, describe a simple curvilinear profile. Figure 3 of Andrews (1973b) documents the same problem. This implies that the Northern Baffin Island Ice Cap exerted a considerable restraining effect upon the area it occupied. Figures 3-4 through 3-10 depict the extent of this effect and show that the shape of the uplift dome closely resembles the shape of the Northern Baffin Island Ice Cap. It is stressed that the isobases over central Baffin Island have been drawn in the only manner that does not

conflict with the data yet reconstructs a physically possible surface. Attempts to fit the data by other patterns resulted in either crossing isolines or creation of cliffs in the statistical surface.

The information contained on the isobase and isochrone maps allows a detailed interpretation of the amount of restrained rebound* imposed by changing ice loads. Figure 3-11 presents equidistant diagrams that graphically portray the influence of the Northern Baffin Island Ice Cap upon strandline geometry and the orderly decline in this influence which accompanied thinning and recession of the ice cap. The northern Baffin Island uplift achieved its greatest amplitude and lasted longest in the region of the present Barnes Ice Cap.

The eastern and western margins of the retreating ice cap are shown on the equidistant diagrams. It is obvious that the relationship between the locations of the crests of the ice cap and the strandlines is not simple. The bases of the troughs in the strandline profiles are not located directly under the western ice margins but are displaced toward the longitudinal axis of the ice cap. This is interpreted as a result of the rigidity of the lithosphere which allowed the strong Southampton Island uplift to extend its influence eastward of the deglaciated area. The bases of the strandline troughs represent the points (or axes when viewed planimetrically; Figs. 3-4 through 3-10) which separate the area influenced dominantly by the strong uplift to the west from that dominated by the more protracted

*Rebound occurring prior to deglaciation of a site.

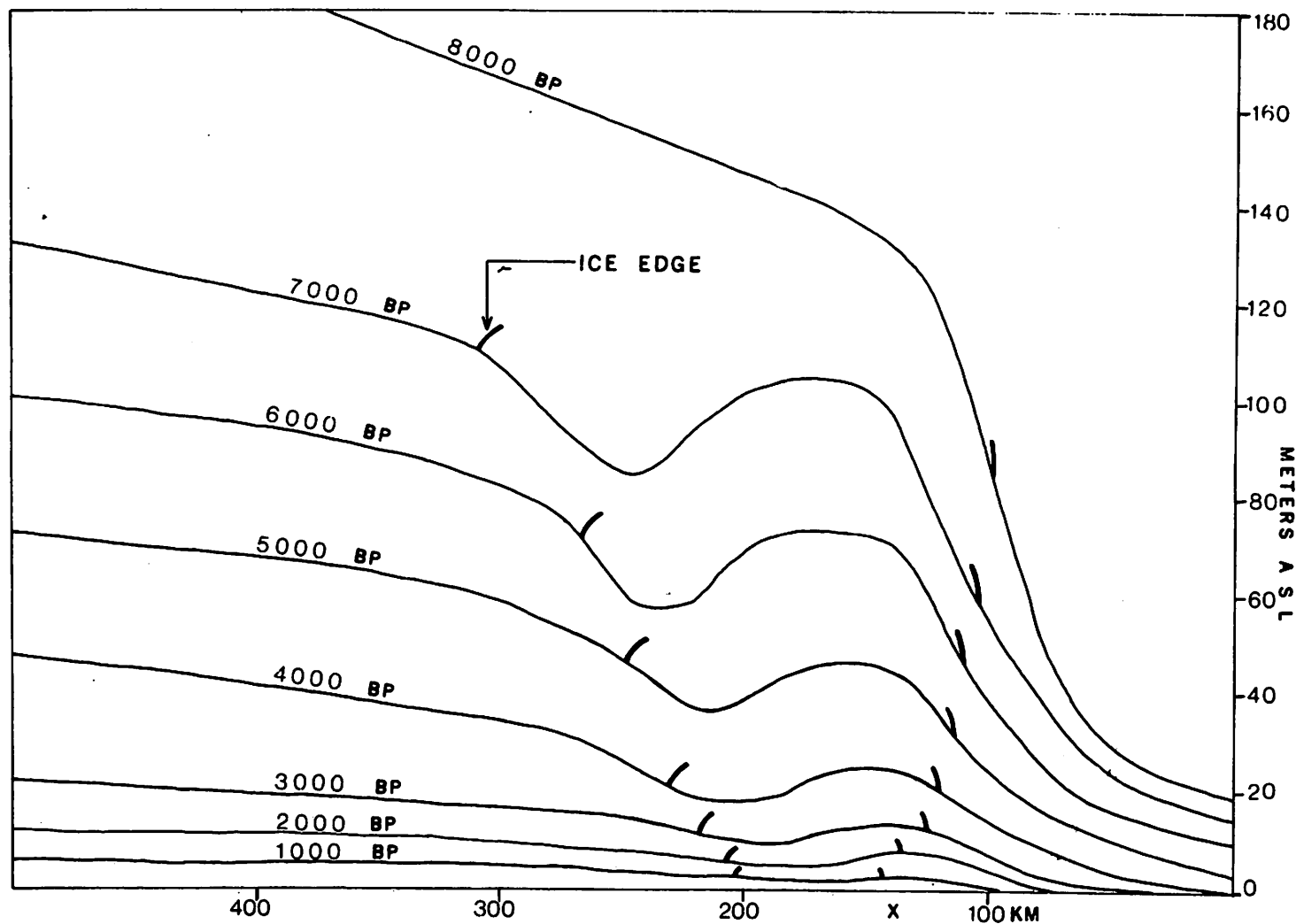


Figure 3-11a: An equidistant diagram constructed from the isobase maps for a profile from Foxe Basin to eastern Baffin Island, running across the Barnes Ice Cap. Vertical exaggeration is 2000 times. See Figure 3-13 for profile location.

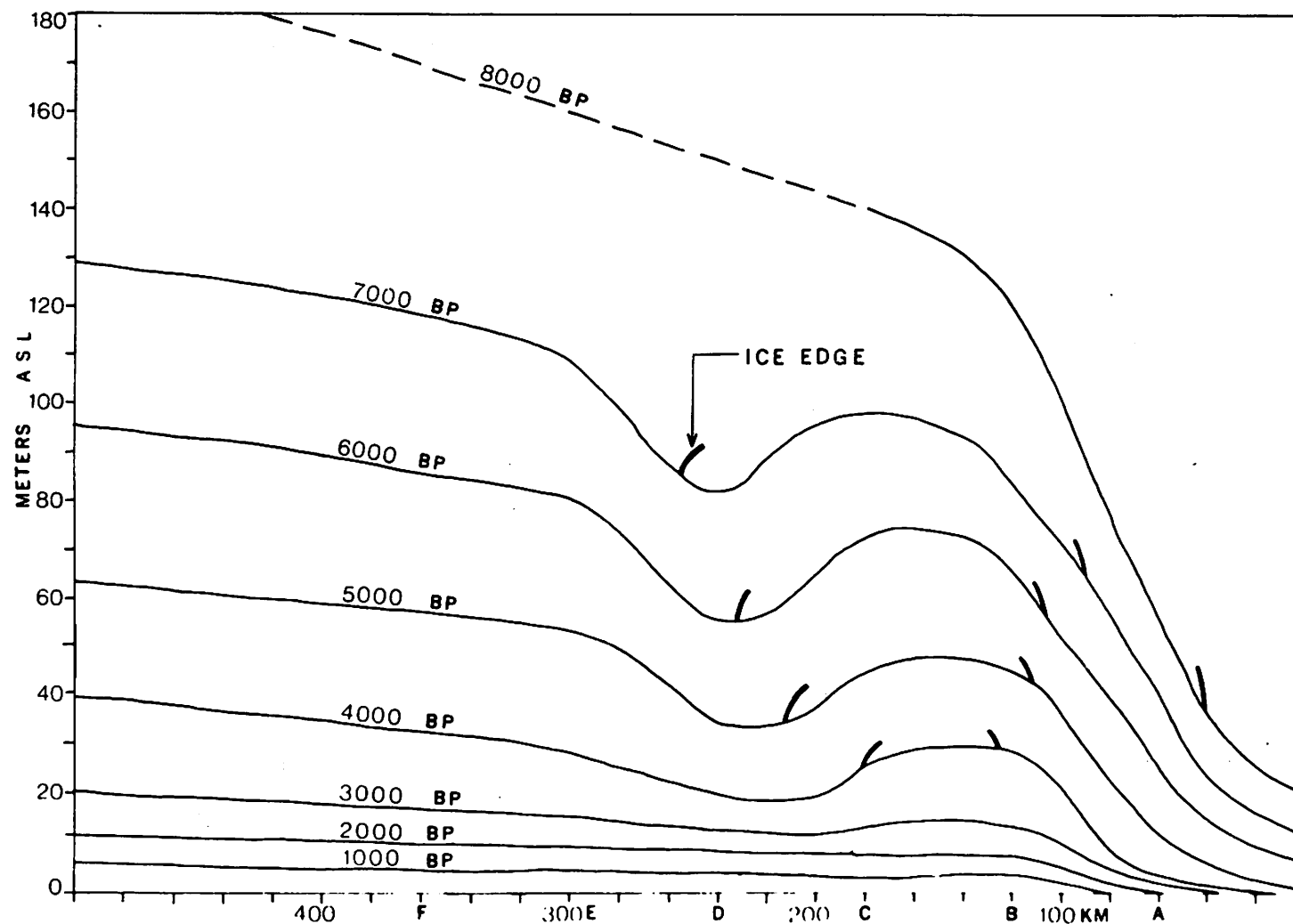


Figure 3-11b: An equidistant diagram constructed from the isobase maps for a profile from Home Bay to Prince Charles Island, Foxe Basin. Vertical exaggeration is 2000 times. See Figure 3-13 for profile location. Points A through F are the locations for which the uplift curves of Figure 3-13 are supplied.

decrease in load to the east. It is evident from the strandline geometries that as the volume of the Northern Baffin Island Ice Cap decreased so did its restraining power, until by 2000 yr B.P. (Fig. 3-11a) almost the entire area under the ice cap came under the dominating influence of the longer wavelength Southampton Island uplift. The northern Baffin Island uplift cannot, therefore, be considered a discrete unit, but rather a regional perturbation within the general Laurentide uplift. This, just as did the other perturbations, resulted from the peculiar, nearly instantaneous, deglaciation of large tracts of central Laurentide North America following the Cockburn-Cochrane phase about 8000 yr B.P. (Falconer et al., 1965).

The amplitude and wavelength of the perturbation were measured for each strandline and plotted on Fig. 3-12. The maximum amplitude was 20m and the half wavelength was 60 km. The rate of decrease in amplitude (curve A) is initially rapid but decreases toward the present, whereas the reverse holds true for decrease in wavelength (curve B). Curve C shows a plot of amplitude against wavelength through time illustrating that the change in geometry of the strandline profiles is, indeed, at first dominated by decrease in amplitude, but since about 3400 yr B.P., by which time the amplitude is only 5 m, rapid shrinkage in the radius of the depression is more important.

The period of 7000 years which was required to eliminate the effect of the Northern Baffin Island Ice Cap upon the uplift surface is a function of the rate of decrease in the volume of that ice cap. It should not, therefore, be confused with the "relaxation time"

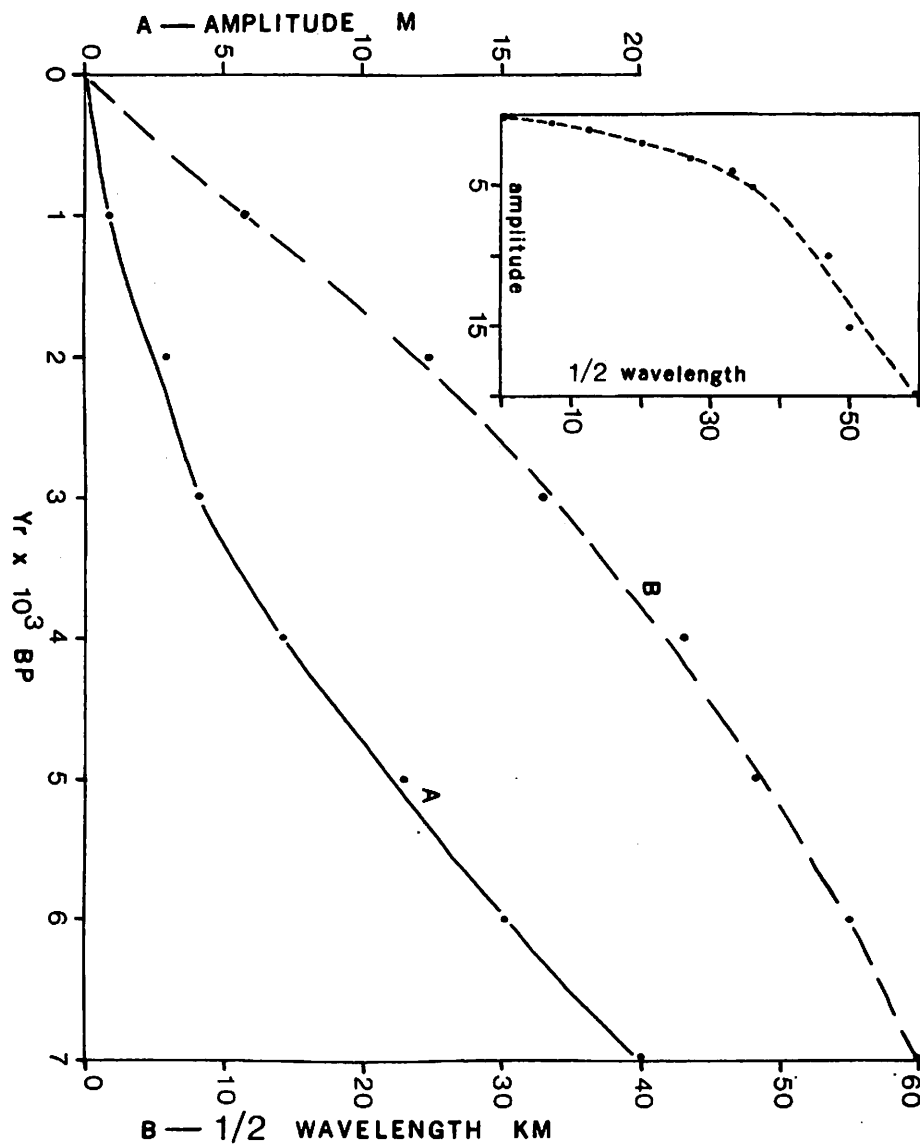


Figure 3-12: Temporal and relative changes in the amplitude and wavelength of the Baffin Island uplift dome. See the inset diagram of Figure 3-14 for the definitions of amplitude and wavelength.

of the mantle to deglaciation, which is controlled by processes operative in the upper mantle, and is a function mainly of the viscosity in the asthenosphere, rather than the rate of unloading.

POSTGLACIAL UPLIFT CURVES

Postglacial uplift curves, almost without exception, show that the rate of rebound decays exponentially as a function of time (Andrews, 1970b). Therefore, the form of the curves generated by an equidistant diagram can provide an independent check upon the accuracy with which the isobases have been placed and, thus, upon the geometry of deformation as interpreted from the field data.

Figure 3-13 contains a series of six uplift curves (semi-log plots) constructed from Fig. 3-11b. The curves show amounts of rebound that have occurred since 8000 yr B.P. for Home Bay and five other points spaced at 60 km intervals along the profile of the equidistant diagram. In all cases the data read from the diagram fall along approximately exponential curves. This is interpreted as an indication of the general accuracy of the isobase maps and the reliability of the elevation measurements, stratigraphic interpretations and radiocarbon age determinations.

The curves, however, depart slightly from a simple exponential function in that they systematically show a slower rate of rebound between 8000 yr and 6000 yr B.P. than is predicted by such a function. This may be accounted for either by: (1) a systematic error in the isobase maps such that uplift is under-estimated on the 8000 yr and 7000 yr B.P. surfaces. The most likely cause of this discrepancy would be the application of too small an eustatic

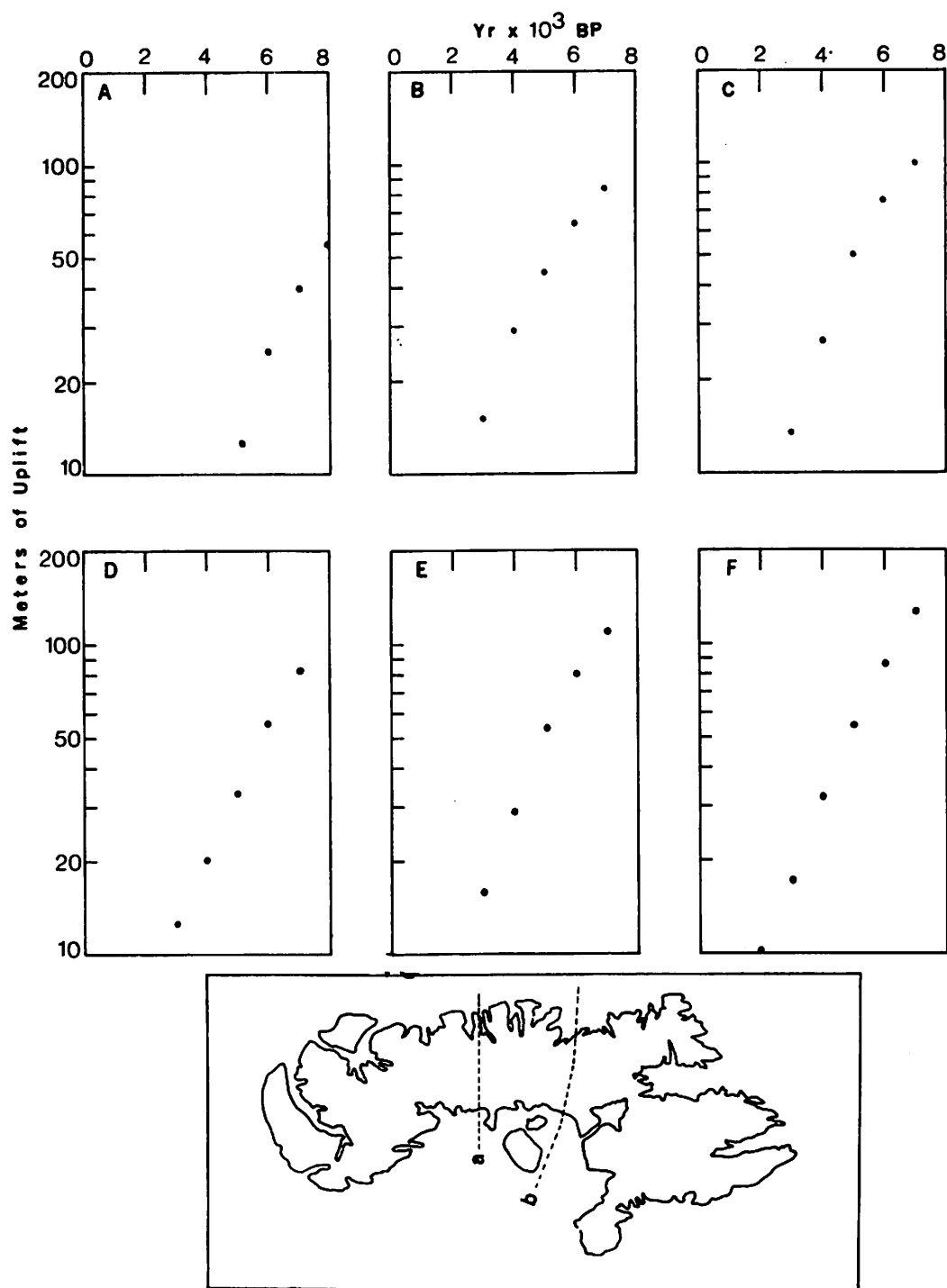


Figure 3-13: Uplift curves derived from the equidistant diagram of Figure 3-11b. Points A through F correspond to the lettered points on the equidistant diagram. The inset map shows the geographic locations of the horizontal axes of the equidistant diagrams.

correction for these times; or (2) the slower rate of rebound prior to 6000 yr B.P. may be real, constituting the restrained rebound portion of the curves (c.f. Andrews 1970b, Fig. 1-6). The more rapid rate of rebound after 6000 yr B.P. is probably a response to the deglaciation of Foxe Basin (Fig. 2-1).

By 2000 yr B.P. the zero isobase lies landward of that point represented by curve A (see insert map, Fig. 3-13). Therefore, if eustatic sea level has continued to rise during the past 2000 years, the outer peninsulas of Home Bay and much of eastern Baffin Island should presently be submerging. Field observations confirm that this is the case (Pheasant, 1971; Schlederman, 1972; Pheasant and Andrews, 1973).

AMPLITUDE OF THE BAFFIN ISLAND UPLIFT DOME AND ICE CAP PARAMETERS

Using the results presented in Table 2-I and equation 2-4, Fig. 3-14 was prepared to illustrate the change in amplitude of the Baffin Island uplift dome as a function of: (a) thickness of ice above the crests of the hypothetical strandlines (curve A), (b) thickness of ice at the center of the ice cap (curve B), and (c) total estimated volume of the ice cap (curves C and C'). The amplitude of the dome was measured from the equidistant diagram (Fig. 3-11a) as the difference in height between the crest of a strandline and the base of the corresponding westward trough (inset, Fig. 3-14).

The curves show the following relationships:

- (a) The amplitude decreases at a decreasing rate as the thickness of ice above the strandline crest decreases. This may be expressed as

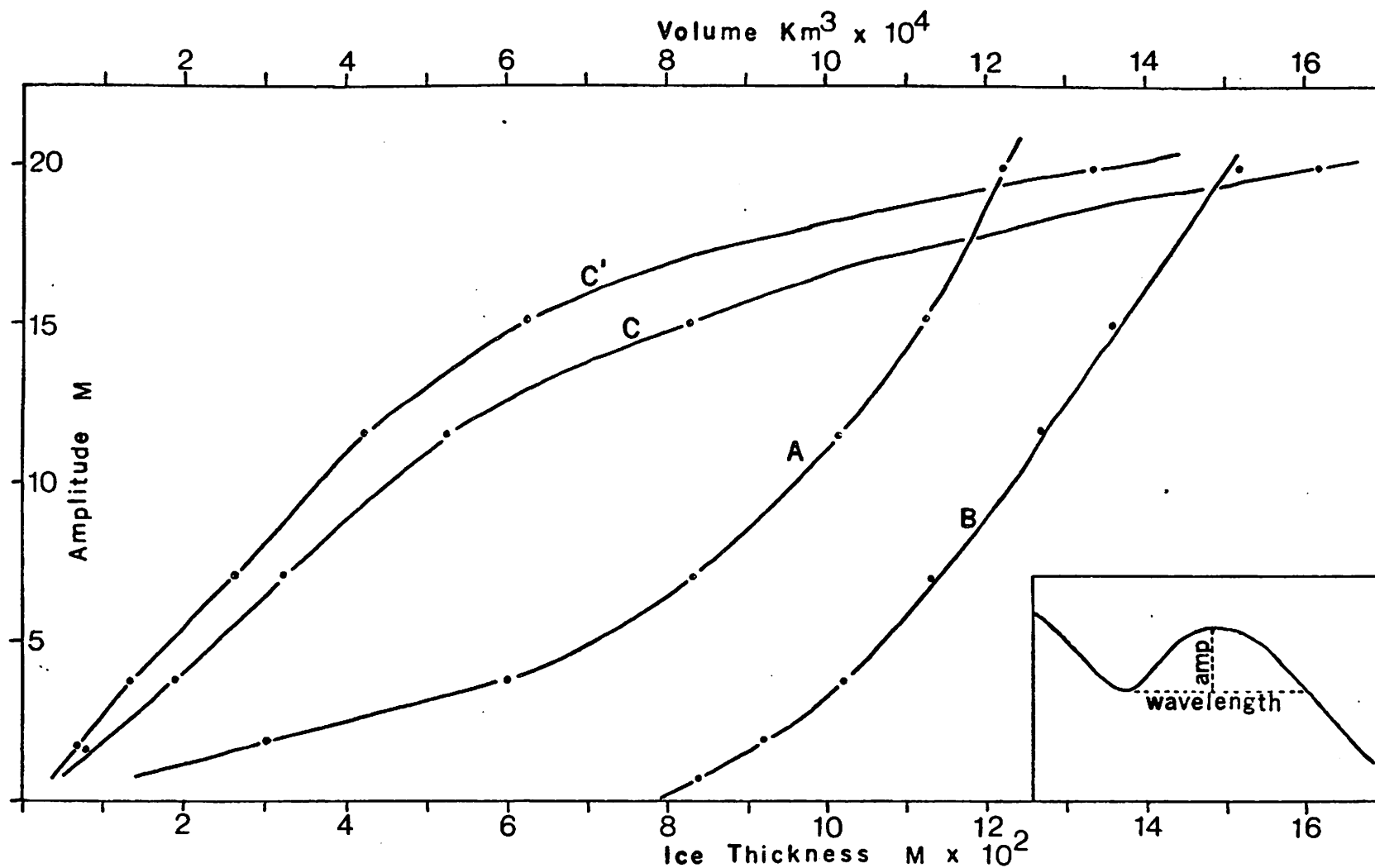


Figure 3-14: The functional relationship between the amplitude of the Baffin Island uplift dome and ice cap parameters (volume and thickness). Curve A: amplitude versus thickness of ice above the crest of the strandline; curve B: amplitude versus thickness of ice at the center of the ice cap; curve C: amplitude versus the volume of the ice cap (from equation 2-1); curve C': amplitude versus the volume of the ice cap (from equation 2-2). The insert graphically defines amplitude and wavelength.

$$A = kT_d^x \quad (3-2)$$

where A is amplitude, T_d is ice thickness above the crest of the strandline and k and x are constants. The amplitude is about 0.5 m when ice thickness is zero. In other words, when the site became deglaciated only about 0.5 m of depression attributable to the load formerly present remained.

(b) The amplitude decreases nearly linearly as the thickness of ice at the center of the ice cap decreases. Extrapolation downward of curve B suggests that a central ice cap thickness of 750 to 800 m may be the critical lower limit on the size of the load necessarily to produce recognizable depression of the crust. This is only slightly less than the present central thickness of the Barnes Ice Cap and a value of 0.5 m seems reasonable for the equilibrium depression of that load.

(c) The amplitude decreases at a decreasing rate as the volume of the Northern Baffin Island Ice Cap decreases. The curves of volume vs. amplitude pass through or nearly through the origin when extrapolated downward, suggesting that a volume of less than $2 \times 10^3 \text{ km}^3$ of ice may produce a small amount of depression or restraint of proceeding rebound.

Figure 3-15 plots volume of the Northern Baffin Island Ice Cap against the thickness of ice at its center. When the central thickness is 750 to 800 m the volume will be about $2 \times 10^3 \text{ km}^3$. For thicknesses exceeding 800 m the volume increases exponentially, whereas lower thicknesses relate to very slowly decreasing volumes. This would seem to support the estimates of the lower critical values of thickness and volume which are necessary to produce

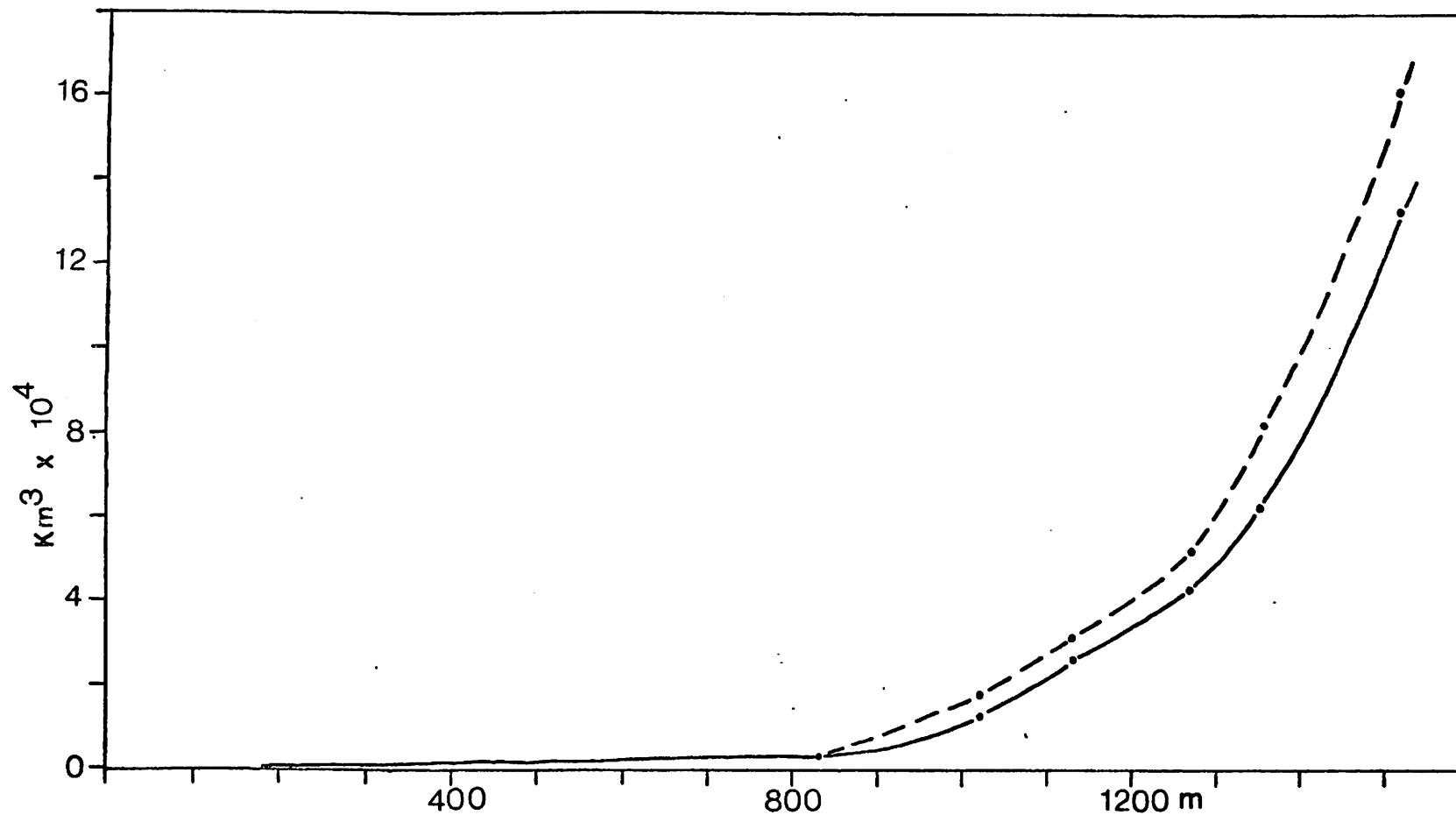


Figure 3-15: Volume of the Northern Baffin Island Ice Cap versus central ice thickness. Broken line is volume according to equation 2-1; solid line in volume according to equation 2-2.

crustal depression.

It is unlikely that depressions of smaller amplitudes and wavelengths than those discussed here can be recognized from age and elevation determinations on postglacial strandlines. Such recognition may be possible when accurate data becomes available through precise levelling and relevening or when purely glacio-isostatic anomalies have been separated from gravimetric data and mapped on large scales.

Barnett (1967) attempted to measure (by levelling) the gradient of the Generator Lake shoreline* which has an age of about 3000 yr. He found that the western end of his traverse was only 22.5 cm higher than the eastern end but concluded that this difference may lie within the error range of the survey method. From Fig. 3-11a one would predict an elevational difference of about 1.0 m over this 20 km distance. However, the great difficulty of identifying in the field exactly synchronous points on a strandline makes it nearly impossible to determine the exact magnitude of such a low gradient. Barnett's field results cannot, therefore, be conclusively interpreted as either enforcing or weakening the validity of the construction of the 3000 yr B.P. strandline on Fig. 3-11a. It is, however, encouraging that the results may at least indicate a slight westward increase in the elevation of the strandline.

Ives (1962) determined that the Isortog Lake shorelines which are dated between 200 yr and 350 yr B.P. are not measurably tilted.

* A strandline formed around a lake dammed by the retreating Northern Baffin Island Ice Cap and located near the present southeastern margin of the Barnes Ice Cap.

This lack of tilting is very probably real as Fig. 3-11a shows that the northern Baffin Island uplift dome has a negligible amplitude since 1000 yr B.P.

Figure 3-16 shows the total uplift that has occurred at a point (X) on the profile of Fig. 3-11a. Curve A is the standard time-elevation curve portraying the roughly exponential decay in the rate of uplift. (Note, however, that this curve shows the same departure from a simple exponential function as those discussed above.) Curve B is a plot of the total amount of uplift achieved since the disappearance of calculated ice thicknesses (based on equation 2-4) from directly above the point. For an area that has experienced unrestrained (true postglacial) rebound for only the past 500 years, this is a more physically meaningful relationship. The points plot with only small scatter from a straight line on semi-log paper and may, therefore, be described by equation 3-2. The scatter of points about the line is probably due mainly to errors in the estimation of ice thicknesses resulting from slight errors in the isochrone map.

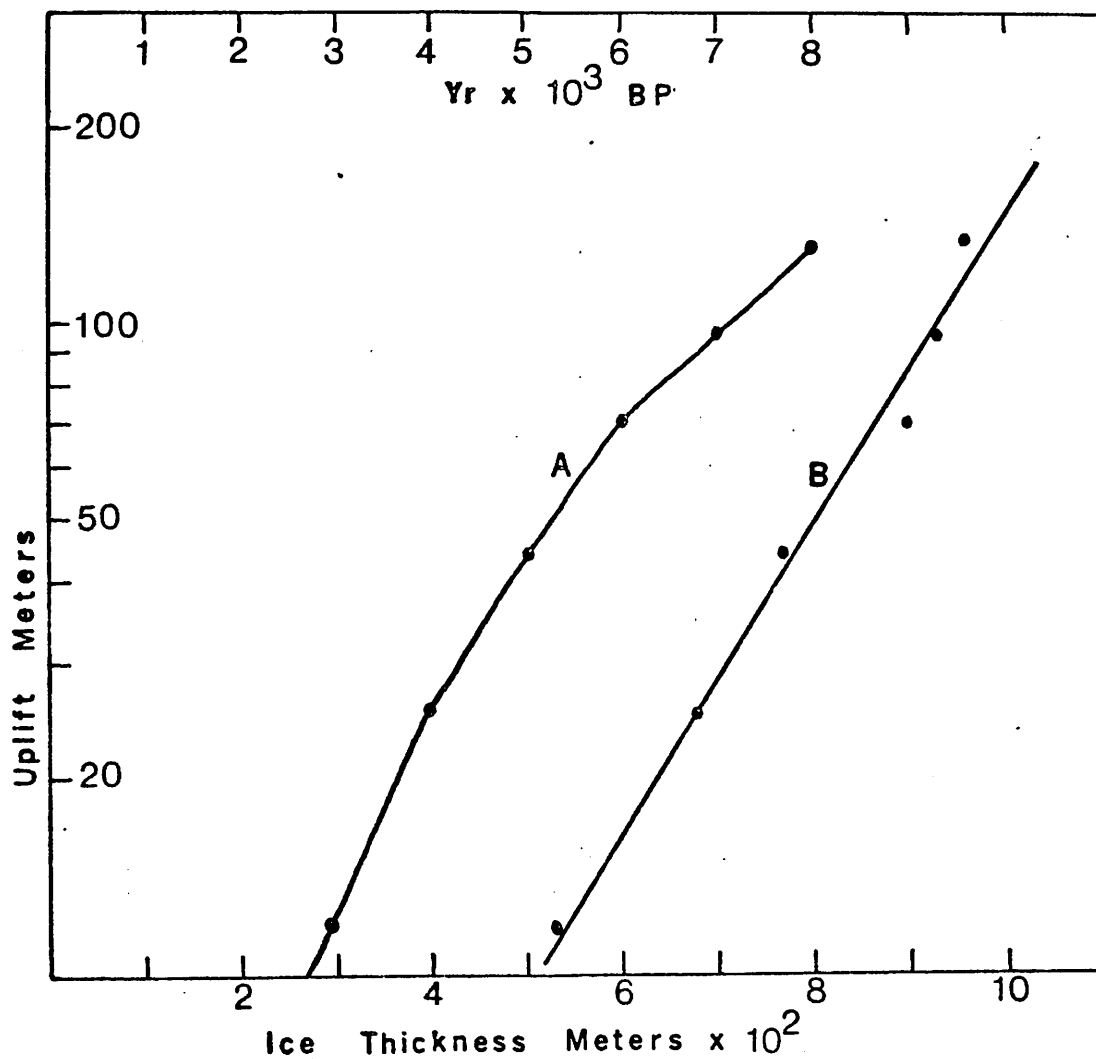


Figure 3-16: Curve A: Uplift as a function of time
 Curve B: Uplift as a function of the rate of unloading...i.e.
 Uplift occurring since the removal of calculated ice thickness
 from immediately above the point represented by the curves (see
 point X on Figure 3-11a).

RESIDUAL REBOUND

Unlike postglacial rebound, residual (remaining) rebound cannot be measured and can be at best only predicted. There is not at present a consensus concerning the correct (or best) method of prediction of this quantity. Estimates of the amount of residual rebound on Southampton Island, for instance, vary from 110 m (Andrews, 1970b) to 140 m (Walcott, 1970). The former estimate is based on the rate of postglacial rebound, while the latter is based on the magnitude of Free Air gravity anomalies, which through a process of smoothing supposedly correlate closely with the amount of mass deficiency in the mantle due to the former ice load. Estimates by these authors are more divergent still for other areas of Hudson Bay.

Here an attempt is made to predict the amount of residual rebound for Southampton Island from the measured apparent "migration" of isobases during postglacial time toward the uplift center over that area and the graphic extrapolation of this migration into the future. The basic assumption of this method is that when the zero isobase reaches the center, rebound will be complete (the lithosphere will be in a state of isostatic equilibrium).

Figure 3-17 presents plots of isobase positions through time relative to the Southampton Island uplift center for three radii (see inset map Fig. 3-18). The lines, therefore, are plots of isobase migration. The last "known" position of any isobase is that portrayed on the map of uplift accomplished since 1000 yr B.P. (Fig. 3-10). Thereafter, isobase positions are speculatively placed by extrapolation of the migration line trends.

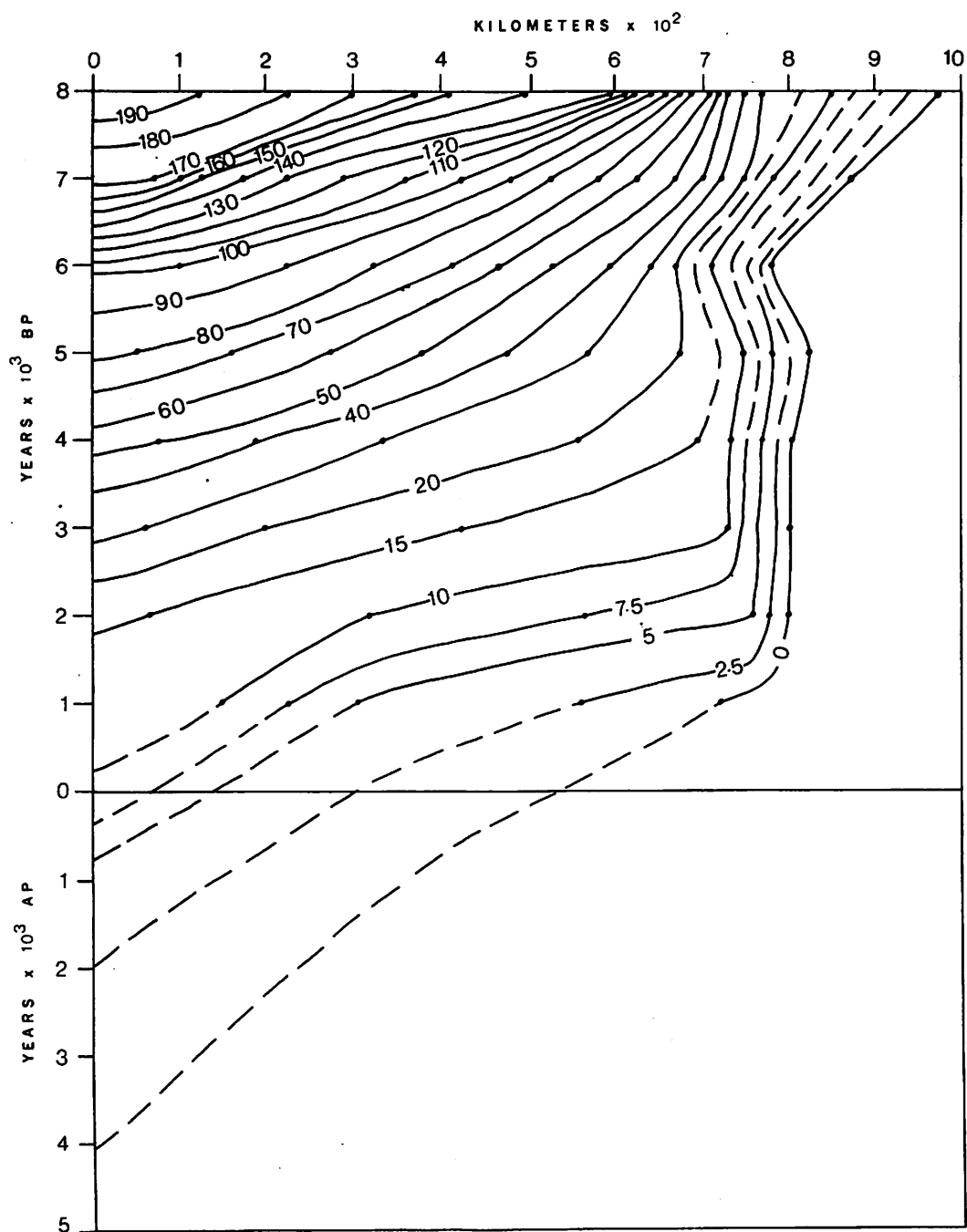


Figure 3-17a: Measured and extrapolated positions of isobases through time along a profile from Southampton Island to the mouth of Frobisher Bay (Profile A, inset map, Figure 3-18).

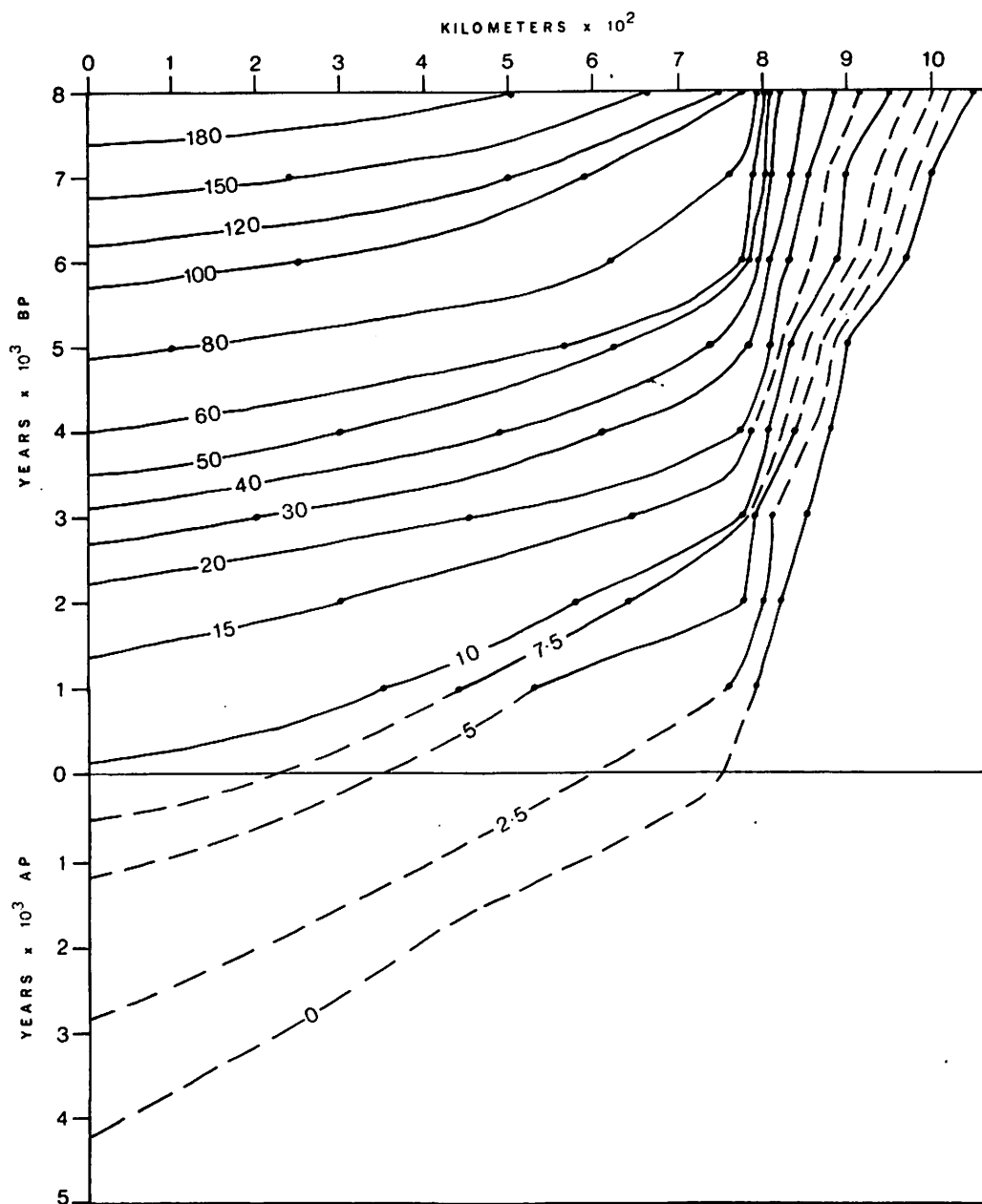


Figure 3-17b: Measured and extrapolated positions of isobases through time along a profile from Southampton Island northeastward across the Barnes Ice Cap (Profile B, inset map, Figure 3-18).

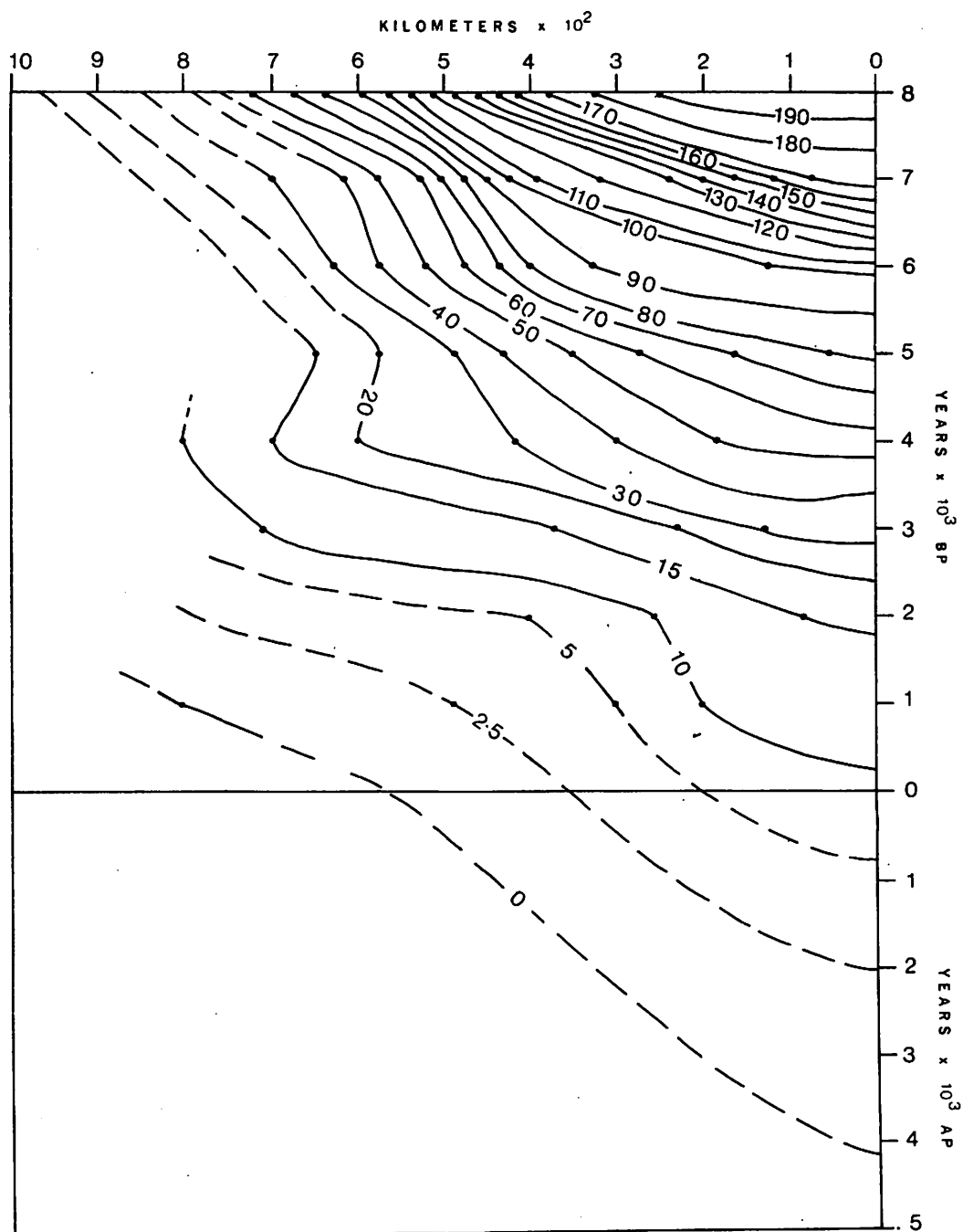


Figure 3-17c: Measured and extrapolated positions of isobases through time along a profile from Southampton Island to King William Island (Profile C, inset map, Figure 3-18).

The extrapolations, if reasonable, suggest that the zero isobase will have migrated to the center, and, therefore, that uplift will be complete in approximately 4000 years from the present. Since 1000 yr B.P. approximately 12 m of uplift have been accomplished at the center. Even if the rate of rebound continues undiminished, the amount of residual rebound of Southampton Island will be only about 48 m.

In the discussion above of the change in geometry of the Baffin Island uplift dome it was shown that this change was at first dominated by decrease in amplitude and later by rapid shrinkage of areal extent. The plots of Fig. 3-17 suggest that the same sequence of events characterizes larger uplift units.

Using the curves of Fig. 3-17, a plot was made of (a) the distance of an isobase from the center of uplift, vs. (b) the length of time required for that isobase to migrate to the center (Fig. 3-18). The three curves are similar, approximately exponential forms, and the difference between them may be attributed to the non-circular form of the Southampton Island uplift dome. It would seem likely that Fig. 3-18 may also be interpreted as a plot of (a) the radius of a depressed region, vs. (b) the time required to complete recovery of that depression. If so the recovery time increases exponentially as the radius of the depressed region increases. The relationship shown by this graph may not necessarily apply strictly to other regions which may have different crustal and/or mantle properties.

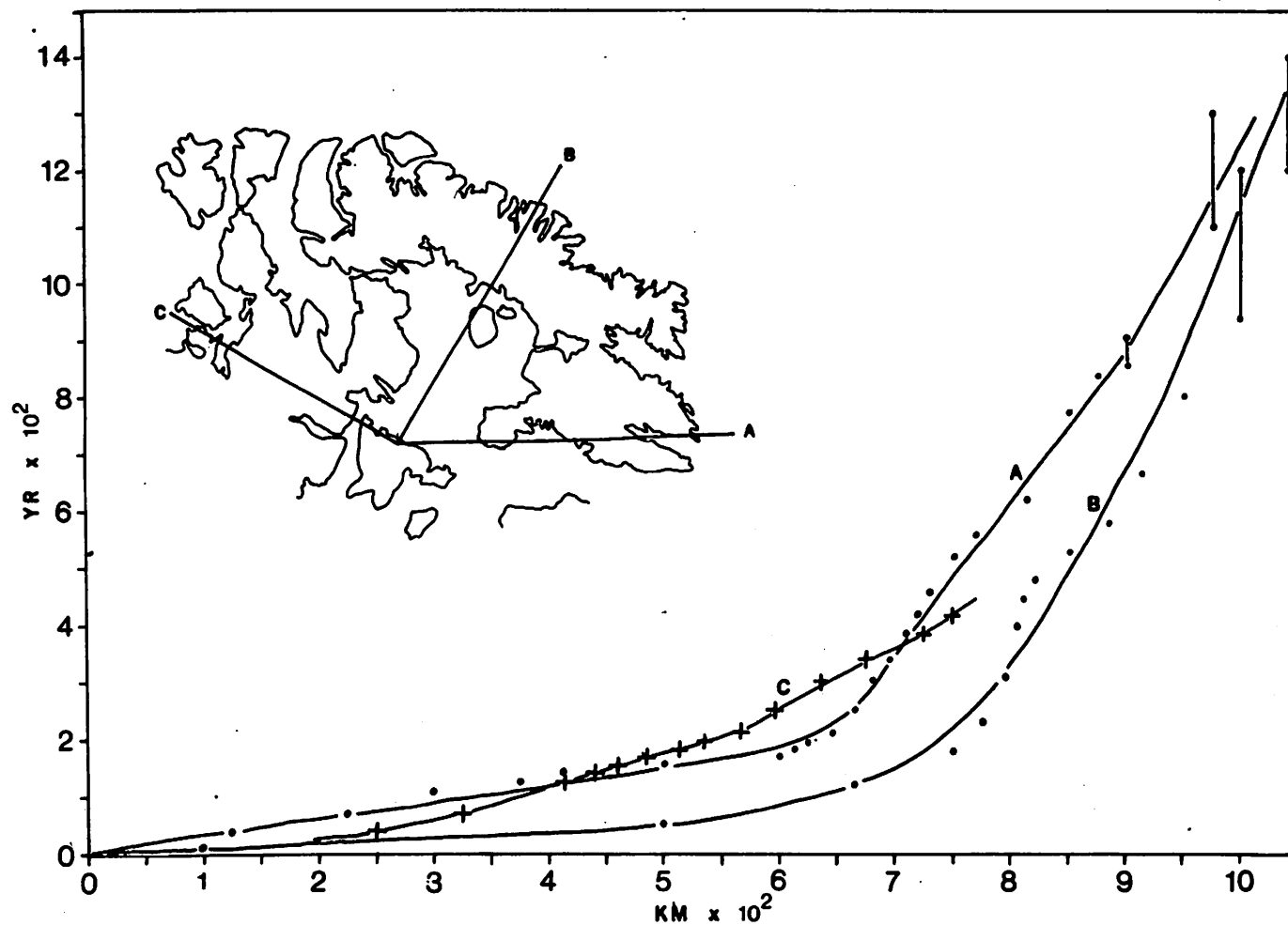


Figure 3-18: Distance of an isobase from the regional center of uplift versus the time required for that isobase to migrate to the center. Data is from Figures 3-17a, b, and c.

CHAPTER 4

CONCLUSIONS AND IMPLICATIONS

Detailed conclusions were presented in the preceding chapters. The following list reiterates the major points of the discussion in a more generalized manner. Following the outline of the conclusions, implications of the results for geomorphologists and geophysicists are discussed.

CONCLUSIONS

1) There is little information on the pattern and chronology of growth of the northeastern sector of the Laurentide Ice Sheet. A good working hypothesis has been proposed by Ives and Andrews (1963) which is supported by the findings of Terrasmae et al. (1966). Accurate absolute age determinations on the Isortoq plant beds and on shells transported from Foxe Basin and deposited in till over western Baffin Island would provide very useful dating control on the initiation and maximum of this sector of the ice sheet. Shells transported from Hudson Strait and deposited in till on southern Baffin could provide equally useful dates. Ives (personal communication, 1974) reports the discovery by Løken of shelly limestone tills in extreme northern Labrador.

2) Late Wisconsin ice over eastern and northern Baffin Island remained at its maximum until approximately 8000 yr B.P. The maximum stand was along the Cockburn Moraines or the time-stratigraphic equivalents of these moraines.

3) The retreat of ice over Baffin Island is described with reasonable accuracy by Fig. 2-1 although future field research will undoubtedly necessitate significant revisions. The isochrone map has been used to produce reasonable estimates of temporal change in ice area, volume and thickness.

4) The 8000 yr B.P. isobase surface reflects with considerable detail the late Wisconsin maximum ice stand. Depression due to the ice load extended about 200 km beyond the margin, while deflection at the margin was about 90 m.

5) The pattern of isobases on the 7000 yr B.P. surface is significantly different from that on the 8000 yr B.P. surface. By 7000 yr B.P. the Laurentide Ice Sheet had fragmented into several major bodies. This pattern of decay is reflected in the isobase surfaces for 7000 yr to 1000 yr B.P.

6) The Northern Baffin Island Ice Cap, due to its considerable size and slow rate of retreat, produced a restraining effect upon the isostatic recovery of central Baffin Island. The maximum dimensions of the central Baffin uplift were 20 m amplitude by 65 km transverse radius by 650 km longitudinal radius. The decrease of these dimensions was controlled by the rate of decrease of the ice load.

7) From the combination of isochrone and isobase maps it is possible to produce uplift curves showing the functional relationship between amounts of uplift and rates of unloading. This information is probably more physically meaningful than traditional curves showing uplift with time.

8) Isobase positions are continuously displaced toward the regional uplift centers as the crust recovers from the depression of former ice loads. The pattern of this "isobase migration" during the postglacial provides a basis for extrapolation of the future migration of the zero isobase. On the basis of such extrapolation the zero isobase is expected to reach the center of uplift on Southampton Island in about 4000 yr time. Because the future rate of uplift is unlikely to exceed the average rate for the past 1000 years (ca. $1.2 \text{ m } 100 \text{ yr}^{-1}$), the amount of residual rebound at the center is at most 48 m.

IMPLICATIONS

A) Geomorphological implications

1) Due to the inherent strength (rigidity) of the lithosphere, an ice sheet will depress the crust for a considerable distance beyond its own margin. In the zone of peripheral depression marine strandlines may form which become elevated subsequent to glacial unloading. The presence of a raised marine feature does not, therefore, necessarily mean that the subjacent area was glacierized immediately prior to the formation of that feature. Likewise, the common practice of using the marine limit age as being synonymous with the date of deglaciation should be applied with a great deal more care.

B) Geophysical implications

1) Small wavelength disturbances on the overall deformation surface produced by an ice sheet can be caused by regional ice bodies during deglaciation. These perturbations have completely

relaxed by now.

2) The crustal flexural parameter can be calculated as follows:

$$\alpha = D (1.108)^{-1} \text{ (Walcott, 1970) } , \quad (4-1)$$

where α is the flexural parameter and D is the distance from the ice edge to the proximal side of the forebulge. The proximal side of the forebulge is defined by the zero isobase. D on eastern Baffin Island is approximately 200 km. Therefore, the flexural parameter in the eastern Baffin Island area is ca 180 km. This is in agreement with the value computed by Walcott for the Caribou Mountains area of Alberta. In that area the minimum value of the flexural parameter lies between 110 and 140 km. Pheasant and Andrews (1973) derived an estimated value of ≥ 135 km for northern Cumberland Peninsula.

3) If the predicted migration of the zero isobase is reasonable, uplift will be complete at the Southampton Island uplift center within 4000 years. This precludes a lengthy relaxation time for glacio-isostatic recovery (cf. Walcott, 1970, p. 5).

C) Implications for geophysical modelling and future field research

1) Attempts to construct geophysical models of glacio-isostatic recovery generally proceed on the assumption that the ice sheet had produced an equilibrium depression prior to the onset of general retreat. To test this assumption it is necessary to firmly establish the pattern and chronology of ice sheet growth (see conclusion 1 above).

2) Because smaller wavelength depressions may have been produced by residual ice caps centered over regional retreat areas, it is important that geophysical models should predict these depressions as well as the general depression field produced by the ice sheet at its maximum. Field work leading to the production of detailed, large scale isochrone and isobase maps of these retreat areas is an essential companion to such modelling.

The Glacial Map of Canada shows that with only minor breaks the periphery of the Labrador-Ungava Peninsula is covered with postglacial marine and lacustrine sediments, and that these sediments extend well inland from the coast in many regions. Careful studies of the deformation of strandlines in these areas should provide the information needed for a detailed analysis of the interaction between changing ice loads and crustal response. Equally good opportunities for extension and refinement of data exist in Keewatin.

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APPENDIX A

ID: RADIOCARBON DATA BANK THESIS QUERY 3 A. DYKE

READ DATA BANK

STORED ON TAPE 03/10/73

THE SIZE OF THE DATA BANK IS MEASURED BY THE FOLLOWING
CRITICAL VALUES.

NO. OF DESCRIPTORS [NODESC] = 16

NO. OF NUMERIC DESCRIPTORS [FTBI] = 7

NO. OF DESCRIPTOR DICTIONARY ENTRIES [NODDE] = 16

NO. OF DESCRIPTOR-STATE SUBDICTIONARY RESERVATIONS
[NODSSE] = 4036

NO. OF OVERFLOW ENTRIES [OVX] = 149

NO. OF STRINGS IN IFILES [NOIF] = 156

NO. OF BUFFERS OF IFILES [IBUF] = 1

LENGTH OF LAST BUFFERLOAD OF IFILES [LWAIF] = 18

NO. OF BITS IN LAST WORD OF LAST BUFFERLOAD [LBAIF] = 24

NO. OF ITEMS = 1044

PRINT: (LATITUDE, LONGITUDE, SAMPLE NO., LAB NO.),
(C14 DATE, STANDARD ERROR, MATERIAL DATED, INFERRED

RELIABILITY OF DATE),

(UPLIFT, ELEVATION OF SAMPLE, INFERRED RELATIVE SEA LEVEL),
(ELEVATION OF LOCAL MARINE LIMIT,

DATE ASSOCIATED WITH MARINE LIMIT),

(RELATION TO ICE MARGIN, DISTANCE FROM MAXIMUM ICE MARGIN,
REFERENCE)

FOR ITEMS WITH LATITUDE. FROM 5800 TO 7600 AND LONGITUDE,
FROM 6000 TO 10000

AND C14 DATE, FROM 8250 TO 8750*

NO. OF ITEMS IN QUERY RESPONSE = 27
 NO. OF ITEMS IN THE DATA BANK = 1044
 PERCENTAGE OF RESPONSE/TOTAL DATA BANK = 2.59

6224 6629 538 GSC 463
 8710 180 SH DEFINITE
 62 METERS 20 METERS 38 METERS
 --- ---
 ICE NEARBY --- GSCRCVI BLAKE

6749 8825 95 GSC288
 8620 140 SH LIKELY
 213 METERS 190 METERS 1 METERS
 215 METERS PROBABLE
 ICE FAR AWAY --- GSC5 RC 1966

6756 6603 26 GSC1638
 8410 340 SH FAIRLY SURE
 58 METERS 34 METERS 37 METERS
 43 METERS ---
 ICE NEARBY --- MILLER UP

6757 6545 537 GAK3092
 8290 170 SH SMALL CHANCE
 50 METERS 23 METERS 30 METERS
 30 METERS ---
 ICE CONTACT --- ---

6805 9009 307 GSC47
 8700 120 SH POSSIBLY
 195 METERS 171 METERS 171 METERS
 190 METERS NO
 ICE FAR AWAY --- GSC1 RC 1962

6812 9034 296 1-179GS
 8370 200 SH LIKELY
 215 METERS 164 METERS 194 METERS
 194 METERS YES
 ICE NEARBY --- I-2 RC 1962

6818 9040 310 GSC40
 8450 110 SH SMALL CHANCE
 29 METERS 8 METERS 0 METERS
 190 METERS ---
 --- --- GSC1 RC 1962

6838 6824 546 GX 0930
 8435 105 SH FAIRLY SURE
 76 METERS 47 METERS 55 METERS
 55 METERS YES
 ICE CONTACT --- ANDREWS 1967

6857 6834 544 GSC 813
 8630 190 SH DEFINITE
 77 METERS 40 METERS 54 METERS
 78 METERS YES
 ICE NEARBY --- ANDREWSETA1970
 6857 6907 542 I 2611
 8300 135 SH FAIRLY SURE
 82 METERS 62 METERS 0 METERS
 87 METERS YES
 ICE CONTACT --- ANDREWS 1967
 6857 6907 543 Y 1830
 8430 140 SH FAIRLY SURE
 107 METERS 70 METERS 86 METERS
 86 METERS YES
 ICE CONTACT --- ANDREWS 1967
 6915 9116 297 I-215GS
 8360 170 SH LIKELY
 214 METERS 175 METERS 194 METERS
 194 METERS YES
 ICE NEARBY --- I-2 RC 1962
 6918 6810 539 I 3236
 8670 140 SH DEFINITE
 72 METERS 48 METERS 0 METERS
 --- YES
 ICE CONTACT --- KING1969
 6928 6752 547 I 3133
 8530 140 SH DEFINITE
 48 METERS 21 METERS 26 METERS
 42 METERS NO
 ICE NEARBY --- KING1969
 7156 7834 530 I 1316
 8250 750 SH DEFINITE
 96 METERS 47 METERS 77 METERS
 77 METERS YES
 ICE NEARBY --- ANDREWS 1967
 7200 7915 529 I 724
 8350 300 SH DEFINITE
 107 METERS 75 METERS 87 METERS
 87 METERS YES
 ICE NEARBY --- ANDREWS 1967
 7442 9448 643 I 1220
 8550 350 SH LIKELY
 137 METERS 114 METERS 114 METERS
 --- PROBABLE
 --- --- BLAKE CJES1970

PRINT: SAME FOR ITEMS WITH LATITUDE. FROM 5800 TO 7600
AND LONGITUDE, FROM
6000 TO 10000 AND C14 DATE, FROM 7750 TO 8250*

7459 9859 125 GSC250
8590 140 SH FAIRLY SURE
132 METERS 106 METERS 109 METERS
109 METERS YES
ICE NEARBY --- GSC6 RC 1967

7512 9804 73 GSC191
8520 150 SH FAIRLY SURE
120 METERS 98 METERS 98 METERS
98 METERS YES
ICE NEARBY --- GSC4 RC 1965

7538 8430 20 S410
8370 115 SH SMALL CHANCE
46 METERS 25 METERS 25 METERS
76 METERS NO
ICE FAR AWAY --- BARR ARCTIC715

7540 8423 142 GSC991
8270 150 BO FAIRLY SURE
57 METERS 38 METERS 38 METERS
73 METERS NO
ICE FAR AWAY --- GSC9 RC 1970

7540 8427 19 GSC991
8270 150 BO FAIRLY SURE
61 METERS 42 METERS 42 METERS
76 METERS NO
ICE FAR AWAY --- BARR ARCTIC715

7540 8435 21 Y1296
8740 120 SH SMALL CHANCE
41 METERS 16 METERS 16 METERS
76 METERS NO
--- --- BARR ARCTIC715

7541 8437 18 Y1295
8250 160 SH SMALL CHANCE
27 METERS 8 METERS 8 METERS
76 METERS NO
ICE FAR AWAY --- BARR ARCTIC715

7545 9825 126 GSC377
8440 150 SH LIKELY
97 METERS 76 METERS 76 METERS
76 METERS YES
ICE NEARBY --- GSC6 RC 1967

7547 8355 367 Y1295
8250 160 BO LIKELY
27 METERS 8 METERS 8 METERS
66 METERS NO
ICE FAR AWAY --- YALE 1X RC1969

7547 8355 368 Y1296
 8740 120 SH LIKELY
 41 METERS 16 METERS 16 METERS
 60 METERS NO
 ICE FAR AWAY --- YALE 1X RC1969

5800 9500 61 GX1063
 8010 95 SH LIKELY
 184 METERS 67 METERS 167 METERS
 150 METERS YES
 ICE CONTACT --- WAGNER 67 \$

5948 6416 347 11322
 8190 710 SH POSSIBLY
 34 METERS 15 METERS 1 METERS
 --- PROBABLE

--- ANDREWS GB1967

5948 6416 536 I 1322
 8190 710 SH FAIRLY SURE
 34 METERS 14 METERS 15 METERS
 --- NO

ICE NEARBY --- ANDREWS67

6212 7538 149 GSC672
 7970 250 SH FAIRLY SURE
 115 METERS 99 METERS 1 METERS
 104 METERS PROBABLE

ICE NEARBY --- GSC7 RC 1968

6212 7538 518 GSC 672
 7970 250 SH FAIRLY SURE
 156 METERS 99 METERS 140 METERS
 140 METERS YES

ICE CONTACT --- MATTHEWS67

6236 7040 513 GSC 425
 7980 220 SH FAIRLY SURE
 127 METERS 75 METERS 110 METERS
 110 METERS YES

ICE CONTACT --- GSCRCVI BLAKE

6253 6950 514 GSC 433
 7880 140 SH FAIRLY SURE
 109 METERS 64 METERS 93 METERS
 93 METERS YES

ICE CONTACT --- GSCRCVI BLAKE

6323 6825 535 GSC 462
 8230 240 SH DEFINITE
 119 METERS 85 METERS 100 METERS
 100 METERS YES

ICE CONTACT --- GSCRCVI BLAKE

6608 6543 516 GAK3093
 7870 150 SH DEFINITE
 51 METERS 27 METERS 35 METERS
 35 METERS YES

ICE CONTACT --- ---

6730	8700	67	MAP			
	8200	200	GE	LIKELY		
		174	METERS	11 METERS	155	METERS
			155	METERS	YES	
			ICE	CONTACT	---	GLMAPCAMADA S
6747	6538	515	GAK2566			
	7950	170	SH	DEFINITE		
		42	METERS	15 METERS	26	METERS
			46	METERS	NO	
			ICE	NEARBY	---	---
6747	6538	534	GAK3090			
	8230	160	SH	FAIRLY SURE		
		45	METERS	15 METERS	26	METERS
			46	METERS	NO	
			ICE	NEARBY	---	---
6805	6625	29	GAK3677			
	7950	140	SH	FAIRLY SURE		
		56	METERS	30 METERS	40	METERS
			40	METERS	YES	
			ICE	NEARBY	---	MEARS UP
6805	6625	510	GAK3677			
	7950	140	SH	DEFINITE		
		56	METERS	30 METERS	40	METERS
			40	METERS	YES	
			ICE	NEARBY	---	---
6805	6625	517	GAK3677			
	7950	140	SH	DEFINITE		
		56	METERS	30 METERS	40	METERS
			40	METERS	YES	
			ICE	NEARBY	---	---
6826	6646	518	Y 1833			
	7960	140	SH	DEFINITE		
		36	METERS	16 METERS	20	METERS
			40	METERS	NO	
			ICE	FAR AWAY	---	ANDREWS 1967
6829	9209	308	GSC46			
	7790	100	SH	LIKELY		
		95	METERS	80 METERS	60	METERS
			190	METERS	NO	
			ICE	FAR AWAY	---	GSCI RC 1962
6832	6810	519	Y 1834			
	7820	140	SH	FAIRLY SURE		
		69	METERS	46 METERS	54	METERS
			77	METERS	YES	
			ICE	CONTACT	---	ANDREWS 1967
6847	6837	509	---			
	7900	150	9E	DEFINITE		
		78	METERS	62 METERS	62	METERS
			62	METERS	YES	
			ICE	CONTACT	---	ANDREWSE7A270

6847 6837 526 ---
 7900 150 GE DEFINITE
 78 METERS 62 METERS 62 METERS
 62 METERS YES
 --- --- ANDREWS ET AL
 6851 9040 295 I-213GS
 7880 150 SH POSSIBLY
 105 METERS 89 METERS 89 METERS
 183 METERS NO
 ICE FAR AWAY --- I-2 RC 1962
 6858 6834 379 ---
 8000 200 GE FAIRLY SURE
 54 METERS 37 METERS 37 METERS
 78 METERS NO
 ICE NEARBY --- ANDREWS ET AL 70
 6858 6903 380 ---
 8000 200 GE FAIRLY SURE
 71 METERS 54 METERS 54 METERS
 88 METERS NO
 ICE NEARBY --- ANDREWS ET AL 70
 6918 6810 520 1 3213
 7970 140 SH DEFINITE
 46 METERS 21 METERS 30 METERS
 50 METERS NO
 ICE NEARBY --- KING 1969
 6928 6752 533 1 3134
 8160 135 SH DEFINITE
 43 METERS 19 METERS 25 METERS
 42 METERS NO
 ICE NEARBY --- KING 1969
 6930 8530 66 MAP
 8200 200 GE LIKELY
 186 METERS 167 METERS 167 METERS
 167 METERS YES
 ICE CONTACT --- GL MAP CANADA S
 6938 7002 521 I 1673
 7970 340 SH FAIRLY SURE
 84 METERS 34 METERS 68 METERS
 68 METERS YES
 ICE CONTACT --- ANDREWS 1967
 6938 7002 522 I 1602
 7900 210 SH FAIRLY SURE
 84 METERS 33 METERS 68 METERS
 68 METERS YES
 ICE CONTACT --- ANDREWS 1967
 6951 6859 532 Y 1705
 8190 120 SH DEFINITE
 46 METERS 22 METERS 27 METERS
 27 METERS YES
 ICE FAR AWAY --- ANDREWS 1967

6952 7028 523 I 1932
 7940 130 SH DEFINITE
 67 METERS 51 METERS 51 METERS
 61 METERS YES
 ICE CONTACT --- ANDREWS 1967

7002 6834 524 I 2831
 7754 135 SH DEFINITE
 27 METERS 7 METERS 12 METERS
 22 METERS NO
 ICE FAR AWAY --- LOKEMUNPUFL

7012 7120 540 GSC 639
 8000 120 SH DEFINITE
 70 METERS 53 METERS 11 METERS
 --- YES
 ICE CONTACT --- ANDREWS 1967

7020 7108 541 I 1933
 8210 130 SH DEFINITE
 38 METERS 19 METERS 6 METERS
 --- YES
 ICE CONTACT --- ANDREWS 1967

7117 7415 528 GSC1060
 8090 140 SH FAIRLY SURE
 108 METERS 77 METERS 90 METERS
 90 METERS YES
 ICE NEARBY --- GSC RC 1971

7121 7253 531 I 1983
 8180 130 SH DEFINITE
 58 METERS 39 METERS 29 METERS
 60 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

7127 7508 527 GSC1064
 7890 160 SH DEFINITE
 90 METERS 74 METERS 74 METERS
 74 METERS YES
 ICE CONTACT --- GSC RC 1971

7156 7834 530 I 1316
 8250 750 SH DEFINITE
 96 METERS 57 METERS 77 METERS
 77 METERS YES
 ICE NEARBY --- ANDREWS 1967

7203 8110 525 I 1246
 7930 300 SH FAIRLY SURE
 77 METERS 61 METERS 61 METERS
 61 METERS YES
 ICE CONTACT --- ANDREWS 1967

7258 9503 120 GSC616
 7750 140 SH FAIRLY SURE
 61 METERS 46 METERS 46 METERS
 --- NO
 --- --- GSC6 RC 1967

7538 8426 17 S434
 8200 140 SH POSSIBLY
 49 METERS 30 METERS 30 METERS
 76 METERS NO
 ICE FAR AWAY --- BARR ARCTIC71s

7541 8437 18 Y1295
 8250 160 SH SMALL CHANCE
 27 METERS 8 METERS 8 METERS
 76 METERS NO
 ICE FAR AWAY --- BARR ARCTIC71s

7547 8355 367 Y1295
 8250 160 BU LIKELY
 27 METERS 8 METERS 8 METERS
 60 METERS NO
 ICE FAR AWAY --- YALE 1X RC1969

7557 9752 158 GSC726
 8090 150 SH SMALL CHANCE
 97 METERS 79 METERS 79 METERS
 95 METERS NO
 --- --- GSC7 RC 1968

PRINT: SAME FOR ITEMS WITH LATITUDE, FROM 5800 TO 7600
 AND LONGITUDE, FROM

6000 TO 10000 AND C14 DATE, FROM 7250 TO 7750*

NO. OF ITEMS IN QUERY RESPONSE = 25

NO. OF ITEMS IN THE DATA BANK = 1044

PERCENTAGE OF RESPONSE/TOTAL DATA BANK = 2.39

5811 9503 622 GSC92
 7270 120 SH FAIRLY SURE
 153 METERS 142 METERS 142 METERS
 145 METERS ---
 ICE NEARBY --- GSC3 RC 1964

5950 8000 148 GSC706
 7430 180 SH POSSIBLY
 168 METERS 139 METERS 156 METERS
 156 METERS YES
 ICE NEARBY --- GSC7 RC 1968

5950 8000 220 GSC 706
 7430 180 SH DEFINITE
 167 METERS 139 METERS 155 METERS
 155 METERS YES
 ICE CONTACT --- ANDREWS67

5950 8000 353 GSC706
 7430 180 SH FAIRLY SURE
 151 METERS 139 METERS 0 METERS
 155 METERS YES
 ICE NEARBY --- ANDREWS GB1967

6214 7542 346 I 1729
 7650 250 SH LIKELY
 125 METERS 111 METERS 1 METERS
 144 METERS PROBABLE
 ICE NEARBY --- ANDREWS GB1967
 6214 7542 595 I 729
 7650 250 SH FAIRLY SURE
 158 METERS 111 METERS 144 METERS
 144 METERS YES
 ICE CONTACT --- ANDREWS67
 6233 7723 98 GSC327
 7350 150 SH LIKELY
 122 METERS 110 METERS 110 METERS
 110 METERS YES
 ICE FAR AWAY --- GSC5 RC 1966
 6233 7723 593 GSC 327
 7350 150 SH FAIRLY SURE
 131 METERS 110 METERS 119 METERS
 140 METERS NO
 ICE NEARBY --- MATTHEWS67
 6253 6951 498 GSC 504
 7490 160 SH FAIRLY SURE
 54 METERS 41 METERS 1 METERS
 93 METERS NO
 ICE NEARBY --- GSCRCVI BLAKE
 6755 6637 30 GUK3678
 7560 130 SH FAIRLY SURE
 67 METERS 45 METERS 54 METERS
 54 METERS YES
 ICE NEARBY --- MEARS . UP
 6755 6637 511 GAK3678
 7560 130 SH DEFINITE
 67 METERS 40 METERS 54 METERS
 54 METERS YES
 ICE NEARBY --- ---
 6843 6750 506 I 3065
 7460 130 SH DEFINITE
 31 METERS 17 METERS 19 METERS
 55 METERS NO
 ICE NEARBY --- ANDREWS 1967
 6844 6839 505 I 3063
 7560 120 SH FAIRLY SURE
 68 METERS 48 METERS 55 METERS
 55 METERS YES
 ICE CONTACT --- ANDREWSETA1970

6847	6837	502	Y	1835			
	7290	120	SH	DEFINITE			
		60 METERS		47 METERS		49 METERS	
		62 METERS		NO			
				ICE NEARBY	---	ANDREWS 1967	
7002	6834	524	I	2831			
	7750	135	SH	DEFINITE			
		27 METERS		7 METERS		12 METERS	
		22 METERS		NO			
				ICE FAR AWAY	---	LOKEMUNPUHL	
7008	9227	116	GSC601				
	7660	150	SH	LIKELY			
		41 METERS		25 METERS		27 METERS	
		---		NO			
				---	---	GSC6 RC 1967	
7009	6855	512	GSC 556				
	7740	140	SH	DEFINITE			
		27 METERS		12 METERS		1 METERS	
		34 METERS		NO			
				ICE FAR AWAY	---	ANDREWS 1967	
7013	7118	503	I	1553			
	7500	200	SH	DEFINITE			
		76 METERS		63 METERS		63 METERS	
		63 METERS		YES			
				ICE CONTACT	---	ANDREWS 1967	
7020	8648	496	GSC 306				
	7690	140	SH	DEFINITE			
		111 METERS		97 METERS		97 METERS	
		97 METERS		YES			
				ICE CONTACT	---	GSCRCV CRAIG	
7117	8743	497	I	1254			
	7560	500	SH	DEFINITE			
		100 METERS		87 METERS		87 METERS	
		87 METERS		YES			
				ICE CONTACT	---	CRAIG 1965	
7258	9503	120	GSC616				
	7750	140	SH	FAIRLY SURE			
		61 METERS		46 METERS		46 METERS	
		---		NO			
				---	---	GSC6 RC 1967	
7442	9459	280	GSC1193				
	7380	140	BO	LIKELY			
		62 METERS		50 METERS		50 METERS	
		---		NO			
				ICE FAR AWAY	---	GSC11 RC 1971	

7541 8437 16 Y1294
 7480 120 SH POSSIBLY
 16 METERS 3 METERS 0 METERS
 76 METERS NO
 ICE FAR AWAY --- BARR ARCTIC715

7541 9848 647 GSC736
 7670 150 SH POSSIBLY
 40 METERS 23 METERS 26 METERS
 --- NO
 ICE FAR AWAY --- BLAKE CJES1970

7547 8355 366 Y1294
 7480 120 SH LIKELY
 16 METERS 3 METERS 3 METERS
 60 METERS NO
 ICE FAR AWAY --- YALE 1X RC1969

PRINT: SAME FOR ITEMS WITH LATITUDE. FROM 5800 TO 7600
 AND LONGITUDE. FROM 6000 TO 10000 AND C14 DATE.
 FROM 6750 TO 7250*
 NO. OF ITEMS IN QUERY RESPONSE = 48
 NO. OF ITEMS IN THE DATA BANK = 1044
 PERCENTAGE OF RESPONSE/TOTAL DATA BANK = 4.60

5950 8000 147 GSC688
 6940 140 SH POSSIBLY
 157 METERS 148 METERS 148 METERS
 156 METERS NO
 --- --- GSC7 RC 1968

5950 8000 352 GSC688
 6940 140 SH LIKELY
 157 METERS 148 METERS . METERS
 155 METERS YES
 ICE NEARBY --- ANDREWS GB1967

6202 7432 362 NPL82
 6760 140 SH LIKELY
 54 METERS 45 METERS . METERS
 45 METERS YES
 --- --- NPL IV RC1966

6207 7438 365 NPL85
 6970 125 SH LIKELY
 100 METERS 90 METERS 90 METERS
 --- PROBABLE
 ICE NEARBY --- NPL IV RC1966

6210 9541 247 1GSC A
 6975 250 SH FAIRLY SURE
 181 METERS 64 METERS 171 METERS
 171 METERS YES
 ICE CONTACT --- LEE59

6210 9541 283 I[GSC]8
 6975 250 SH POSSIBLY
 74 METERS 64 METERS 1 METERS
 171 METERS PROBABLE
 --- --- I-1 RC 1961
 6215 8000 60 GX1070
 7115 100 SH LIKELY
 111 METERS 91 METERS 101 METERS
 --- NO
 ICE FAR AWAY --- WAGNER 67 S
 6216 7602 578 L 765
 7160 195 SH FAIRLY SURE
 130 METERS 113 METERS 119 METERS
 140 METERS NO
 ICE NEARBY --- MATTHEWS67
 6227 7636 345 1726
 7160 195 SH LIKELY
 122 METERS 111 METERS 111 METERS
 136 METERS NO
 ICE FAR AWAY --- ANDREWS GB1967
 6227 7636 596 I 726
 7160 195 SH FAIRLY SURE
 122 METERS 111 METERS 10 METERS
 136 METERS YES
 ICE NEARBY --- ANDREWS67
 6232 7725 99 NPL58
 6900 130 SH FAIRLY SURE
 92 METERS 83 METERS 83 METERS
 110 METERS NO
 ICE FAR AWAY --- NPL3 RC 1966
 6232 7725 360 NPL58
 6900 130 SH LIKELY
 120 METERS 81 METERS 111 METERS
 111 METERS YES
 ICE NEARBY --- NPL111 RC 1965
 6325 7643 119 GSC560
 6920 150 SH SMALL CHANCE
 52 METERS 43 METERS 43 METERS
 150 METERS NO
 --- --- GSC6 RC 1967
 6325 7643 488 GSC 560
 6920 150 SH LIKELY
 52 METERS 43 METERS 4 METERS
 134 METERS NO
 ICE NEARBY --- GSCRCVI BLAKE

6333 8500 81 GSC311
 6950 130 SH FAIRLY SURE
 136 METERS 127 METERS 127 METERS
 127 METERS YES
 ICE NEARBY --- GSC4 RC 1965
 6333 8500 224 I 2443
 6950 130 SH FAIRLY SURE
 136 METERS 127 METERS 0 METERS
 172 METERS YES
 ICE CONTACT --- ANDREWS67 BIRD
 6333 8500 357 12443
 6950 130 SH LIKELY
 136 METERS 127 METERS 127 METERS
 172 METERS NO
 ICE FAR AWAY --- ANDREWS GB1967
 6343 6827 489 GSC 464
 6750 170 SH DEFINITE
 39 METERS 15 METERS 30 METERS
 30 METERS YES
 ICE NEARBY --- GSCRCVI BLAKE
 6350 8100 229 GX 1068
 7170 90 SH FAIRLY SURE
 141 METERS 58 METERS 130 METERS
 130 METERS YES
 ICE CONTACT --- WAGNER67 S
 6417 8257 84 GSC323
 6890 210 SH FAIRLY SURE
 156 METERS 147 METERS 147 METERS
 147 METERS YES
 ICE NEARBY --- GSC4 RC 1965
 6417 8257 202 GSC838
 6890 210 SH LIKELY
 169 METERS 145 METERS 160 METERS
 160 METERS YES
 ICE NEARBY --- GSC11RC 1971
 6417 8257 226 GSC 838
 6890 210 SH FAIRLY SURE
 164 METERS 147 METERS 155 METERS
 155 METERS YES
 ICE CONTACT --- BIRD711
 6419 8829 86 GSC289
 6830 170 SH FAIRLY SURE
 135 METERS 126 METERS 126 METERS
 126 METERS YES
 ICE NEARBY --- GSC4 RC 1965
 6419 8829 94 GSC289
 6830 170 SH FAIRLY SURE
 135 METERS 126 METERS 126 METERS
 149 METERS ---
 ICE FAR AWAY --- GSC5 RC 1966

6443 8446 83 GSC324
 6930 150 SH FAIRLY SURE
 125 METERS 116 METERS 116 METERS
 116 METERS YES
 ICE NEARBY --- GSC4 RC 1965
 6443 8446 211 GSC782
 6930 150 SH LIKELY
 139 METERS 115 METERS 130 METERS
 130 METERS YES
 ICE NEARBY --- GSC11RC 1971
 6443 8446 225 GSC 782
 6930 150 SH FAIRLY SURE
 166 METERS 116 METERS 157 METERS
 157 METERS YES
 ICE CONTACT --- BIRD70
 6519 7309 473 GSC 465
 6830 150 SH DEFINITE
 114 METERS 99 METERS 105 METERS
 105 METERS YES
 ICE CONTACT --- GSCRCVI BLAKE
 6602 7136 485 GSC 466
 6760 140 SH DEFINITE
 107 METERS 91 METERS 98 METERS
 98 METERS YES
 ICE CONTACT --- GSCRCVI BLAKE
 6644 8642 87 GSC286
 6850 140 SH FAIRLY SURE
 130 METERS 121 METERS 121 METERS
 121 METERS YES
 ICE NEARBY --- GSC4 RC 1965
 6644 8642 96 GSC286
 6850 140 SH LIKELY
 130 METERS 121 METERS 0 METERS
 142 METERS PROBABLE
 ICE FAR AWAY --- GSC5 RC 1966
 6747 6538 507 GAK3365
 7100 140 SH DEFINITE
 25 METERS 14 METERS 15 METERS
 46 METERS NO
 ICE FAR AWAY --- ---
 6752 8210 97 GSC291
 6880 180 SH FAIRLY SURE
 143 METERS 134 METERS 134 METERS
 147 METERS NO
 ICE NEARBY --- GSC5 RC 1966
 6752 8210 487 GSC 291
 6880 180 SH DEFINITE
 143 METERS 134 METERS 134 METERS
 134 METERS YES
 ICE CONTACT --- CRAIG1965

6842 9227 293 2-212GS
 7160 160 SH LIKELY
 64 METERS 53 METERS 53 METERS
 183 METERS NO
 ICE FAR AWAY --- I-2 RC 1962
 6844 6856 467 GSC 739
 6930 150 SH LIKELY
 61 METERS 34 METERS 52 METERS
 52 METERS YES
 ICE CONTACT --- ANDREWS 1967
 6918 6810 504 I 3135
 7160 140 SH DEFINITE
 25 METERS 7 METERS 14 METERS
 50 METERS NO
 ICE FAR AWAY --- KING 1969
 6937 7001 499 I 1598
 7200 150 SH DEFINITE
 56 METERS 41 METERS 45 METERS
 68 METERS NO
 ICE NEARBY --- ANDREWS 1967
 6937 7001 500 I 1672
 7080 170 SH DEFINITE
 45 METERS 35 METERS 0 METERS
 68 METERS NO
 ICE NEARBY --- ANDREWS 1967
 6937 7001 501 I 1554
 7030 190 SH DEFINITE
 56 METERS 26 METERS 46 METERS
 68 METERS NO
 ICE NEARBY --- ANDREWS 1967
 7009 6856 493 GSC 599
 7000 150 SH DEFINITE
 31 METERS 21 METERS 21 METERS
 34 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7036 8608 494 GSC 307
 7120 140 SH DEFINITE
 107 METERS 97 METERS 97 METERS
 97 METERS YES
 ICE CONTACT --- GSCRCV CRAIG
 7055 8627 495 GSC 304
 7240 150 SH DEFINITE
 104 METERS 89 METERS 93 METERS
 93 METERS YES
 ICE CONTACT --- GSCRCV CRAIG
 7139 8413 480 GSC 390
 6890 150 SH DEFINITE
 49 METERS 40 METERS 40 METERS
 67 METERS NO
 ICE NEARBY --- GSCRCV CRAIG

7247 9537 248 L 571A
 7150 350 SH LIKELY
 46 METERS 30 METERS 35 METERS
 122 METERS NO
 ICE FAR AWAY --- FARRAND62BIRDS

7247 9537 314 L571A
 7150 350 SH POSSIBLY
 41 METERS 30 METERS 0 METERS
 --- PROBABLE
 --- --- LAMONTV11 RC61

7442 9459 644 GSC1193
 7210 140 BO LIKELY
 61 METERS 50 METERS 50 METERS
 --- NO
 ICE FAR AWAY --- BLAKE CJES1970

7538 8426 15 S428
 6900 115 PT SMALL CHANCE
 66 METERS 57 METERS 57 METERS
 76 METERS NO
 ICE FAR AWAY --- BARR ARCTIC71S

PRINT: SAME FOR ITEMS WITH LATITUDE, FROM 5800 TO 7600
 AND LONGITUDE, FROM

6000 TO 10000 AND C14 DATE, FROM 6250 TO 6750*

NO. OF ITEMS IN QUERY RESPONSE = 33

NO. OF ITEMS IN THE DATA BANK = 1044

PERCENTAGE OF RESPONSE/TOTAL DATA BANK = 3.16

5950 8000 218 I 2415
 6590 125 SH DEFINITE
 108 METERS 94 METERS 100 METERS
 155 METERS NO
 ICE FAR AWAY --- ANDREWS67

5950 8000 350 I2415
 6590 125 SH FAIRLY SURE
 106 METERS 94 METERS 100 METERS
 --- NO
 --- --- ANDREWS GB1967

6207 7438 363 NPL83
 6660 125 SN LIKELY
 83 METERS 75 METERS 0 METERS
 --- NO
 --- --- NPL 1V RC1966

6207 7438 364 NPL84
 6660 130 SH LIKELY
 82 METERS 74 METERS 1 METERS
 --- NO
 --- --- NPL 1V RC1966

6207 7438 591 NPL 84
 6660 130 SH FAIRLY SURE
 83 METERS 74 METERS 75 METERS
 140 METERS NO
 ICE FAR AWAY --- MATTHEWS67
 6215 8030 58 GX1071
 6335 85 SH LIKELY
 83 METERS 66 METERS 76 METERS
 --- NO
 ICE FAR AWAY --- WAGNER 67 \$
 6215 8030 59 GX1069
 6395 90 SH LIKELY
 98 METERS 81 METERS 91 METERS
 --- NO
 ICE FAR AWAY --- WAGNER 67 \$
 6249 9448 181 GSC1016
 6570 140 SH LIKELY
 130 METERS 122 METERS 1 METERS
 168 METERS PROBABLE
 --- --- GSC9 RC 1970
 6250 7254 348 12444
 6580 125 SH LIKELY
 52 METERS 44 METERS 44 METERS
 --- NO
 ICE FAR AWAY --- ANDREWS GB1967
 6250 7254 594 I 2444
 6580 125 SH FAIRLY SURE
 52 METERS 44 METERS 1 METERS
 --- NO
 ICE FAR AWAY --- ANDREWS67
 6343 6827 489 GSC 464
 6750 170 SH DEFINITE
 39 METERS 15 METERS 30 METERS
 30 METERS YES
 ICE NEARBY --- GSCRCVI BLAKE
 6345 6832 491 GSC 533
 6440 160 SH SMALL CHANCE
 37 METERS 4 METERS 30 METERS
 30 METERS YES
 ICE NEARBY --- GSCRCVI MATTH
 6406 8405 82 GSC308
 6535 115 SH POSSIBLY
 174 METERS 166 METERS 166 METERS
 166 METERS YES
 ICE NEARBY --- GSC4 RC 1965

6406 8405 223 GX09986
 6535 115 SH FAIRLY SURE
 168 METERS 160 METERS 160 METERS
 160 METERS YES
 ICE CONTACT --- ANDREWS67 BIRD
 6425 8530 85 GSC337
 6580 125 SH FAIRLY SURE
 96 METERS 88 METERS 88 METERS
 88 METERS YES
 ICE NEARBY --- GSC4 RC 1965
 6425 8530 227 1 2963
 6580 125 SH FAIRLY SURE
 108 METERS 88 METERS 100 METERS
 157 METERS NO
 ICE NEARBY --- BIRD70
 6513 8532 80 GSC309
 6610 125 SH FAIRLY SURE
 120 METERS 112 METERS 112 METERS
 112 METERS YES
 ICE NEARBY --- GSC4 RC 1965
 6513 8532 222 1 2432
 6610 125 SH FAIRLY SURE
 120 METERS 112 METERS 1 METERS
 145 METERS YES
 ICE CONTACT --- ANDREWS67 BIRD
 6513 8532 355 12432
 6610 125 SH LIKELY
 120 METERS 112 METERS 1 METERS
 --- PROBABLE
 ICE FAR AWAY --- ANDREWS GB1967
 6519 7309 474 GSC 553
 6590 140 SH DEFINITE
 98 METERS 88 METERS 90 METERS
 105 METERS NO
 ICE NEARBY --- GSCRCVI BLAKE
 6606 8405 356 GX09986
 6535 114 SH POSSIBLY
 174 METERS 166 METERS 1 METERS
 166 METERS YES
 ICE NEARBY --- ANDREWS GB1967
 6723 6500 492 GAK3685
 6350 110 SH DEFINITE
 25 METERS 15 METERS 18 METERS
 18 METERS YES
 ICE CONTACT --- ---
 6740 9827 69 GSC235
 6740 150 SH LIKELY
 28 METERS 18 METERS 20 METERS
 --- NO
 ICE FAR AWAY --- GSC4 RC 1965

6832 6810 481 I 2695
 6560 125 SH DEFINITE
 42 METERS 26 METERS 34 METERS
 79 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 6832 6823 468 GSC 827
 6650 180 SH FAIRLY SURE
 56 METERS 37 METERS 48 METERS
 59 METERS NO
 ICE NEARBY --- ANDREWS ET AL 1970
 6910 7528 471 I 406
 6725 250 SH DEFINITE
 112 METERS 89 METERS 104 METERS
 104 METERS YES
 ICE CONTACT --- ANDREWS 1967
 6921 7549 484 I 2410
 6270 210 SH FAIRLY SURE
 91 METERS 75 METERS 84 METERS
 94 METERS NO
 ICE NEARBY --- ANDREWS 1967
 6955 6843 476 I 2962
 6520 150 GR DEFINITE
 19 METERS 9 METERS 11 METERS
 21 METERS NO
 ICE FAR AWAY --- LOKEMUNPURL
 6956 6840 475 I 1934
 6560 125 SH DEFINITE
 19 METERS 11 METERS 1 METERS
 21 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7003 7129 486 GSC 633
 6270 150 SH DEFINITE
 57 METERS 50 METERS 0 METERS
 --- YES
 ICE CONTACT --- ANDREWS 1967
 7115 7458 483 GSC 1094
 6330 140 SH POSSIBLY
 39 METERS 32 METERS 32 METERS
 91 METERS YES
 --- GSC RC 1971
 7153 8055 466 GSC 328
 6410 150 SH DEFINITE
 61 METERS 54 METERS 54 METERS
 54 METERS YES
 ICE CONTACT --- ANDREWS 1967
 7153 8055 469 GSC 328
 6410 150 SH DEFINITE
 64 METERS 45 METERS 57 METERS
 57 METERS YES
 ICE CONTACT --- GSC RV I

PRINT:SAME FOR ITEMS WITH LATITUDE, FROM 5800 TO 7600
AND LONGITUDE, FROM

6000 TO 10000 AND C14 DATE, FROM 5750 TO 6250*

NO. OF ITEMS IN QUERY RESPONSE = 20

NO. OF ITEMS IN THE DATA BANK = 1044

PERCENTAGE OF RESPONSE/TOTAL DATA BANK = 1.92

6208 7441 152 GSC801

6080 160 SH FAIRLY SURE

61 METERS 55 METERS 55 METERS

--- NO

ICE FAR AWAY ---

GSC7 RC 1968

6210 7558 590 N 285

6070 140 SH FAIRLY SURE

36 METERS 21 METERS 30 METERS

140 METERS NO

ICE FAR AWAY ---

MATTHEWS67

6215 8325 57 Gx1066

5815 90 SH LIKELY

95 METERS 15 METERS 90 METERS

--- NO

ICE FAR AWAY ---

WAGNER 67 \$

6343 6827 490 GSC 503

6140 170 SH DEFINITE

20 METERS 14 METERS 1 METERS

30 METERS NO

ICE NEARBY ---

GSCRCVI MATTH

6346 9540 93 GSC439

5900 130 SH FAIRLY SURE

81 METERS 70 METERS 76 METERS

137 METERS ---

ICE FAR AWAY ---

GSC5 RC 1966

6447 7552 470 GSC 561

6120 140 SH LIKELY

82 METERS 76 METERS 1 METERS

150 METERS NO

ICE NEARBY ---

GSCRCVI

6608 9618 146 GSC693

6140 150 SH FAIRLY SURE

55 METERS 49 METERS 49 METERS

152 METERS NO

ICE FAR AWAY ---

GSC7 RC 1968

6842 6921 452 I 2412

5900 130 SH FAIRLY SURE

59 METERS 33 METERS 54 METERS

54 METERS YES

ICE CONTACT ---

ANDREWS 1967

6852 6925 449 I 3062

5840 150 SH DEFINITE

48 METERS 39 METERS 43 METERS

48 METERS YES

				ICE CONTACT	---	ANDREWSETA1970
6853	6902	465	1	3065		
	6190	120	SH	FAIRLY SURE		
		36 METERS	30 METERS	0 METERS		
		40 METERS	YES			
				ICE CONTACT	---	ANDREWSETA1971
6857	6903	482	I	2583		
	6130	120	SH	DEFINITE		
		30 METERS	16 METERS	24 METERS		
		87 METERS	NO			
				ICE FAR AWAY	---	ANDREWS 1967
6912	7443	582	F1			
	6000	200	GE	FAIRLY SURE		
		76 METERS	70 METERS	70 METERS		
		70 METERS	YES			
				ICE CONTACT	---	ANDREWS70AAR
6914	7425	581	F2			
	5800	200	GE	FAIRLY SURE		
		65 METERS	60 METERS	60 METERS		
		60 METERS	YES			
				ICE CONTACT	---	ANDREWS70AAR
6918	7537	472	I	405		
	6050	250	SH	DEFINITE		
		82 METERS	73 METERS	76 METERS		
		104 METERS	NO			
				ICE NEARBY	---	ANDREWS 1967
6937	7001	478	I	1596		
	6150	170	SH	DEFINITE		
		43 METERS	31 METERS	37 METERS		
		68 METERS	NO			
				ICE FAR AWAY	---	ANDREWS 1967
6950	7020	479	GSC	631		
	6220	140	SH	DEFINITE		
		37 METERS	31 METERS	31 METERS		
		61 METERS	NO			
				ICE NEARBY	---	ANDREWS 1967
7000	7137	477	I	1556		
	6240	140	SH	DEFINITE		
		72 METERS	46 METERS	66 METERS		
		66 METERS	YES			
				ICE CONTACT	---	ANDREWS 1967
7009	7735	463	I	486		
	5750	250	SH	POSSIBLY		
		60 METERS	55 METERS	55 METERS		
		95 METERS	NO			
				ICE FAR AWAY	---	ANDREWS 1967
7236	9520	121	GSC	652		
	5960	140	SH	FAIRLY SURE		
		27 METERS	12 METERS	21 METERS		
		---	NO			
				---	---	GSC6 RC 1967

7540 8437 14 5432
 6100 125 80 FAIRLY SURE
 17 METERS 11 METERS 11 METERS
 76 METERS NO
 ICE FAR AWAY --- BARR ARCTIC715

PRINT: SAME FOR ITEMS WITH LATITUDE, FROM 5800 TO 7600
 AND LONGITUDE, FROM
 6000 TO 10000 AND C14 DATE, FROM 5250 TO 5750*

NO. OF ITEMS IN QUERY RESPONSE = 20
 NO. OF ITEMS IN THE DATA BANK = 1044
 PERCENTAGE OF RESPONSE/TOTAL DATA BANK = 1.92

6412 8150 56 GX1067
 5440 360 SH LIKELY
 84 METERS 76 METERS 80 METERS
 --- NO
 ICE FAR AWAY --- WAGNER 67 \$

6431 9603 92 GSC299
 5480 150 SH FAIRLY SURE
 94 METERS 90 METERS 90 METERS
 130 METERS NO
 ICE FAR AWAY --- GSC5 RC 1960

6448 8259 228 S 13
 5600 300 SH FAIRLY SURE
 61 METERS 52 METERS 56 METERS
 155 METERS NO
 ICE FAR AWAY --- BIRD70

6448 8259 337 S13
 5600 300 SH LIKELY
 65 METERS 52 METERS 60 METERS
 --- NO
 ICE FAR AWAY --- ANDREWS GB1967

6610 9014 311 GSC41
 5470 140 SH LIKELY
 65 METERS 56 METERS 61 METERS
 122 METERS NO
 ICE FAR AWAY --- GSC1 RC 1962

6723 6455 455 GX 1824
 5330 450 SH LIKELY
 22 METERS 16 METERS 18 METERS
 18 METERS YES
 ICE CONTACT --- ---

6836 6850 464 I 2411
 5380 185 SH DEFINITE
 38 METERS 31 METERS 34 METERS
 34 METERS YES
 ICE CONTACT --- ANDREWS 1967

6858 6834 466 I 2548
 5580 130 SH DEFINITE
 22 METERS 18 METERS 18 METERS
 78 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 6907 7502 462 I 1831
 5570 130 SH DEFINITE
 68 METERS 64 METERS 64 METERS
 83 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 6916 7415 580 F3
 5500 200 GE FAIRLY SURE
 54 METERS 50 METERS 50 METERS
 50 METERS YES
 ICE CONTACT --- ANDREWS 70AAR
 6928 7531 453 I 1833
 5270 140 SH DEFINITE
 71 METERS 51 METERS 67 METERS
 14 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7001 7130 456 I 1321
 5390 150 SH LIKELY
 41 METERS 37 METERS 1 METERS
 66 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7009 7735 463 I 486
 5750 250 SH POSSIBLY
 60 METERS 55 METERS 55 METERS
 96 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7016 7143 451 I 1243
 5560 250 SH LIKELY
 31 METERS 27 METERS 0 METERS
 57 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7111 7503 448 GSC1163
 5500 180 SH LIKELY
 35 METERS 29 METERS 31 METERS
 44 METERS NO
 ICE NEARBY --- GSC RC 1971
 7159 7915 458 L 7620
 5400 200 SH DEFINITE
 32 METERS 4 METERS 28 METERS
 77 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7203 8107 454 I 1319
 5710 200 SH FAIRLY SURE
 66 METERS 17 METERS 61 METERS
 17 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

7540 8436 13 S431
 5280 100 DW FAIRLY SURE
 15 METERS 11 METERS 11 METERS
 76 METERS NO
 ICE FAR AWAY --- BARR ARCTIC718

7540 8436 641 S431
 5280 100 DW FAIRLY SURE
 15 METERS 11 METERS 11 METERS
 --- NO
 ICE FAR AWAY --- BLAKE CJES1970

7558 8958 203 GSC1072
 5250 130 DR FAIRLY SURE
 29 METERS 26 METERS 26 METERS
 --- NO
 ICE FAR AWAY --- GSC11RC 1971

PRINT: SAME FOR ITEMS WITH LATITUDE, FROM 5800 TO 7600
 AND LONGITUDE, FROM
 6000 TO 10000 AND C14 DATE, FROM 4750 TO 5250*

NO. OF ITEMS IN QUERY RESPONSE = 23
 NO. OF ITEMS IN THE DATA BANK = 1044
 PERCENTAGE OF RESPONSE/TOTAL DATA BANK = 2.20

5950 8000 221 I 2547
 4960 130 SH DEFINITE
 58 METERS 50 METERS 55 METERS
 155 METERS NO
 ICE FAR AWAY --- ANDREWS67

5950 8000 354 12547
 4960 130 SH LIKELY
 58 METERS 50 METERS 55 METERS
 155 METERS NO
 ICE FAR AWAY --- ANDREWS GB1967

6210 7558 589 N 284
 5230 130 SH FAIRLY SURE
 18 METERS 13 METERS 15 METERS
 140 METERS NO
 ICE FAR AWAY --- MATTHEWS67

6213 7538 150 GSC812
 4770 140 OR FAIRLY SURE
 30 METERS 24 METERS 28 METERS
 104 METERS NO
 ICE FAR AWAY --- GSC7 RC 1968

6735 6519 432 GAK3091
 4950 140 SH FAIRLY SURE
 12 METERS 5 METERS 9 METERS
 17 METERS NO
 ICE FAR AWAY --- ---

6749 6611 28 GAK3724
 4810 110 SH LIKELY
 10 METERS 2 METERS 8 METERS
 49 METERS NO
 ICE FAR AWAY --- MILLER UP
 6832 6808 461 I 2549
 5100 120 SH DEFINITE
 18 METERS 14 METERS 15 METERS
 75 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 6847 6627 27 GAK3723
 5200 100 SH LIKELY
 25 METERS 17 METERS 22 METERS
 --- NO
 ICE FAR AWAY --- MILLER UP
 6848 6924 446 I 3066
 4850 120 SH FAIRLY SURE
 39 METERS 13 METERS 37 METERS
 37 METERS YES
 ICE CONTACT --- ANDREWS ETA 1970
 6852 6927 429 I 2442
 4990 175 SH FAIRLY SURE
 43 METERS 9 METERS 40 METERS
 40 METERS YES
 ICE CONTACT --- ANDREWS 1967
 6917 7410 579 F4
 5200 200 GE FAIRLY SURE
 48 METERS 45 METERS 45 METERS
 45 METERS YES
 ICE CONTACT --- ANDREWS 70AAR
 6920 7359 577 F5
 5000 200 GE FAIRLY SURE
 47 METERS 44 METERS 44 METERS
 44 METERS YES
 ICE CONTACT --- ANDREWS 70AAR
 6928 7003 450 I 2096
 5190 120 SH DEFINITE
 43 METERS 29 METERS 40 METERS
 40 METERS YES
 ICE NEARBY --- ANDREWS 1967
 6955 6843 434 I 2961
 4830 120 OR DEFINITE
 11 METERS 4 METERS 9 METERS
 21 METERS NO
 ICE FAR AWAY --- LOKEMUNPUBL

6955 7654 459 I 1244
 5070 450 SH FAIRLY SURE
 46 METERS 43 METERS 43 METERS
 73 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7001 7133 435 I 1670
 4770 140 SH POSSIBLY
 22 METERS 16 METERS 20 METERS
 66 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7001 7133 436 I 1931
 4920 180 SH DEFINITE
 23 METERS 16 METERS 20 METERS
 66 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7001 7133 438 I 1669
 4770 140 SH DEFINITE
 24 METERS 18 METERS 22 METERS
 66 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7121 7253 457 I 1238
 5070 200 SH LIKELY
 20 METERS 17 METERS 17 METERS
 66 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7156 7835 440 I 1245
 4875 350 SH DEFINITE
 40 METERS 38 METERS 38 METERS
 78 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7442 9450 645 W127
 5050 350 SH POSSIBLY
 21 METERS 18 METERS 1 METERS
 --- NO
 ICE FAR AWAY --- BLAKE CJES1970
 7541 9848 646 GSC783
 4750 140 SH POSSIBLY
 28 METERS 24 METERS 26 METERS
 --- NO
 ICE FAR AWAY --- BLAKE CJES1970
 7558 8958 203 GSC1072
 5250 130 DR FAIRLY SURE
 29 METERS 26 METERS 26 METERS
 --- NO
 ICE FAR AWAY --- GSC11RC 1971

PRINT: SAME FOR ITEMS WITH LATITUDE, FROM 5800 TO 7600
 AND LONGITUDE, FROM
 6000 TO 10000 AND C14 DATE, FROM 4250 TO 4750*

NO. OF ITEMS IN QUERY RESPONSE = 10
 NO. OF ITEMS IN THE DATA BANK = 1044
 PERCENTAGE OF RESPONSE/TOTAL DATA BANK = .96

6239 6939 635 GAK1281
 4460 100 OR LIKELY
 15 METERS 14 METERS 0 METERS
 --- NO
 ICE FAR AWAY --- BLAKE CJES1970

6836 6855 445 I 2582
 4590 115 SH DEFINITE
 25 METERS 10 METERS 23 METERS
 34 METERS NO
 ICE NEARBY --- ANDREWS 1967

6852 6927 430 I 2584
 4430 110 SH FAIRLY SURE
 27 METERS 19 METERS 26 METERS
 40 METERS NO
 ICE NEARBY --- ANDREWS 1967

6853 6901 428 I 2413
 4420 110 SH FAIRLY SURE
 22 METERS 18 METERS 21 METERS
 40 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

6909 9359 309 GSC45
 4460 80 SH LIKELY
 25 METERS 24 METERS 24 METERS
 190 METERS NO
 ICE FAR AWAY --- GSC1 RC 1962

7001 7133 437 I 1671
 4270 140 SH DEFINITE
 25 METERS 18 METERS 24 METERS
 66 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

7009 7735 443 I 487
 4700 210 SH FAIRLY SURE
 23 METERS 21 METERS 21 METERS
 96 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

7152 7833 441 I 1318
 4400 490 SH POSSIBLY
 17 METERS 14 METERS 16 METERS
 88 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

7538 8428 12 S430
 4300 95 PT POSSIBLY
 27 METERS 26 METERS 26 METERS
 76 METERS NO
 ICE FAR AWAY --- BARR ARCTIC715
 7541 9848 646 GSC783
 4750 140 SH POSSIBLY
 28 METERS 24 METERS 26 METERS
 --- NO
 ICE FAR AWAY --- BLAKE CJES1970

PRINT: SAME FOR ITEMS WITH LATITUDE, FROM 5800 TO 7600
 AND LONGITUDE, FROM
 6000 TO 10000 AND C14 DATE, FROM 3750 TO 4250*

NO. OF ITEMS IN QUERY RESPONSE = 16
 NO. OF ITEMS IN THE DATA BANK = 1044
 PERCENTAGE OF RESPONSE/TOTAL DATA BANK = 1.53

6207 7438 361 NPL71
 3900 125 SH LIKELY
 12 METERS 10 METERS 12 METERS
 --- NO
 ICE FAR AWAY --- NPL 111 RC1965
 6207 7438 592 NPL 71
 3900 125 SH FAIRLY SURE
 12 METERS 10 METERS 12 METERS
 140 METERS NO
 ICE FAR AWAY --- MATTHEWS67
 6239 6939 444 P 707
 4067 73 --- FAIRLY SURE
 11 METERS 12 METERS 10 METERS
 93 METERS NO
 ICE FAR AWAY --- ESK SITE ARC71
 6239 6939 634 P707
 4067 73 OR LIKELY
 16 METERS 15 METERS 0 METERS
 --- NO
 ICE FAR AWAY --- BLAKE CJES1970
 6244 6941 423 P 708
 3814 69 --- FAIRLY SURE
 16 METERS 18 METERS 16 METERS
 93 METERS NO
 ICE FAR AWAY --- ESK SITE ARC71
 6254 6951 422 GSC 596
 3750 140 SH SMALL CHANCE
 3 METERS 3 METERS 0 METERS
 93 METERS NO
 ICE FAR AWAY --- GSCRCVI BLAKE

6345 6832 447 GSC 849
 4140 130 --- DEFINITE
 14 METERS 15 METERS 13 METERS
 30 METERS NO
 ICE FAR AWAY --- ESK SITE ARC71
 6832 6801 419 I 2585
 3850 105 SH FAIRLY SURE
 5 METERS 2 METERS 5 METERS
 79 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 6843 6910 433 I 2546
 4050 130 SH FAIRLY SURE
 21 METERS 15 METERS 20 METERS
 20 METERS YES
 ICE CONTACT --- ANDREWS 1967
 6847 6837 415 I 2586
 3890 107 SH LIKELY
 8 METERS 3 METERS 8 METERS
 62 METERS NO
 ICE NEARBY --- ANDREWS 1967
 6906 7448 442 GSC 557
 4000 140 SH DEFINITE
 27 METERS 3 METERS 26 METERS
 83 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 6910 8359 250 P 207
 3958 168 BO LIKELY
 47 METERS 50 METERS 46 METERS
 137 METERS NO
 ICE FAR AWAY --- FARRAND62ISOTI
 6910 8359 427 P 207
 3958 168 --- FAIRLY SURE
 24 METERS 51 METERS 23 METERS
 134 METERS NO
 ICE FAR AWAY --- ESK SITE ARC71
 6937 7001 439 I 1597
 4090 150 SH DEFINITE
 21 METERS 14 METERS 20 METERS
 68 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7001 7133 412 I 1668
 3830 140 SH DEFINITE
 19 METERS 17 METERS 19 METERS
 66 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7153 8055 431 I 1320
 4010 440 SH LIKELY
 22 METERS 18 METERS 21 METERS
 54 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

PRINT: SAME FOR ITEMS WITH LATITUDE, FROM 5800 TO 7600
 AND LONGITUDE, FROM
 6000 TO 10000 AND C14 DATE, FROM 3250 TO 3750*

NO. OF ITEMS IN QUERY RESPONSE = 20
 NO. OF ITEMS IN THE DATA BANK = 1044
 PERCENTAGE OF RESPONSE/TOTAL DATA BANK = 1.92

5950 8000 219 1 2417
 3530 110 SH DEFINITE
 21 METERS 10 METERS 21 METERS
 155 METERS NO
 ICE FAR AWAY --- ANDREWS67

5950 8000 351 12417
 3530 110 SH FAIRLY SURE
 21 METERS 10 METERS 21 METERS
 --- NO
 ICE FAR AWAY --- ANDREWS GB1967

6230 7740 587 MEEVS 6
 3500 0 EK LIKELY
 28 METERS 32 METERS 28 METERS
 140 METERS NO
 ICE FAR AWAY --- ANDREWSEA71 \$

6239 6937 425 M 1531
 3480 200 --- FAIRLY SURE
 9 METERS 11 METERS 9 METERS
 93 METERS NO
 ICE FAR AWAY --- ESK SITE ARC71

6239 6937 633 GSC1051
 3390 210 OR LIKELY
 18 METERS 18 METERS 0 METERS
 --- NO
 ICE FAR AWAY --- BLAKE CJES1970

6239 6939 424 P 710
 3577 69 --- FAIRLY SURE
 12 METERS 14 METERS 12 METERS
 93 METERS NO
 ICE FAR AWAY --- ESK SITE ARC71

6254 6951 422 GSC 596
 3750 140 SH SMALL CHANCE
 3 METERS 3 METERS 0 METERS
 93 METERS NO
 ICE FAR AWAY --- GSCRCVI BLAKE

6411 8311 216 S 12
 3670 270 SH FAIRLY SURE
 36 METERS 32 METERS 36 METERS
 155 METERS NO
 ICE FAR AWAY --- LEE60

6411 8311 336 S12
 3670 270 SH LIKELY
 32 METERS 32 METERS 32 METERS
 --- NO
 ICE FAR AWAY --- ANDREWS GB1967
 6735 6519 411 GSC1507
 3570 140 PT FAIRLY SURE
 5 METERS 7 METERS 5 METERS
 22 METERS NO
 ICE FAR AWAY --- ---
 6822 9747 294 I-178GS
 3690 120 SH LIKELY
 22 METERS 22 METERS 22 METERS
 174 METERS NO
 ICE FAR AWAY --- I-2 RC 1962
 6858 6903 421 Y 1831
 3580 120 SH DEFINITE
 5 METERS 1 METERS 5 METERS
 87 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 6917 7415 410 I 2830
 3585 140 SH FAIRLY SURE
 19 METERS 14 METERS 19 METERS
 52 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 6937 7001 413 I 1601
 3530 130 SH LIKELY
 17 METERS 13 METERS 17 METERS
 68 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 6937 7001 414 I 1600
 3520 230 SH FAIRLY SURE
 15 METERS 11 METERS 15 METERS
 68 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 6950 7020 416 GSC 584
 3450 170 PT FAIRLY SURE
 6 METERS 6 METERS 6 METERS
 61 METERS NO
 ICE NEARBY --- ANDREWS 1967
 6956 7702 418 I 1247
 3550 200 WO DEFINITE
 18 METERS 18 METERS 18 METERS
 70 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7131 8433 417 GSC 190
 3480 130 SH LIKELY
 3 METERS 3 METERS 3 METERS
 67 METERS NO
 ICE FAR AWAY --- GSCRCIVCR:IG

7152 7833 420 I 1317
 3600 480 SH DEFINITE
 16 METERS 15 METERS 16 METERS
 88 METERS NO
 ICE FAR AWAY --- ANDREWS 1967
 7550 8330 739 INFER
 3500 500 CU LIKELY
 9 METERS 12 METERS 9 METERS
 --- NO
 ICE FAR AWAY --- ANDREWSETAL71S

PRINT: SAME FOR ITEMS WITH LATITUDE. FROM 5800 TO 7600
 AND LONGITUDE, FROM
 6000 TO 10000 AND C14 DATE, FROM 2750 TO 3250*

NO. OF ITEMS IN QUERY RESPONSE = 15
 NO. OF ITEMS IN THE DATA BANK = 1044
 PERCENTAGE OF RESPONSE/TOTAL DATA BANK = 1.44

5837 9349 194 GSC685
 3180 140 SH FAIRLY SURE
 38 METERS 38 METERS 38 METERS
 --- NO
 ICE FAR AWAY --- GSC11RC 1971
 5844 9405 195 GSC261
 3040 130 SH SMALL CHANCE
 23 METERS 23 METERS 23 METERS
 --- NO
 ICE FAR AWAY --- GSC11RC 1971
 5850 9355 63 GX1072
 3190 80 SH FAIRLY SURE
 32 METERS 30 METERS 32 METERS
 150 METERS YES
 ICE FAR AWAY --- WAGNER 67 \$
 5900 9400 62 GX1065
 2800 110 SH FAIRLY SURE
 40 METERS 38 METERS 40 METERS
 150 METERS YES
 ICE FAR AWAY --- WAGNER 67 \$
 6213 7539 151 GSC818
 2840 160 OR FAIRLY SURE
 6 METERS 6 METERS 6 METERS
 104 METERS NO
 ICE FAR AWAY --- GSC7 RC 1968
 6244 6941 426 P 699
 3043 63 --- DEFINITE
 8 METERS 10 METERS 8 METERS
 93 METERS NO
 ICE FAR AWAY --- ESK SITE ARC71

6910 8359 251 P 213
 2910 129 BO FAIRLY SURE
 17 METERS 21 METERS 17 METERS
 137 METERS NO
 ICE FAR AWAY --- FARRAND62ISOTI

6910 8359 408 P 213
 2910 129 --- FAIRLY SURE
 15 METERS 22 METERS 15 METERS
 134 METERS NO
 ICE FAR AWAY --- ESK SITE ARC71

6920 8220 749 P213
 2910 129 --- LIKELY
 19 METERS 22 METERS 19 METERS
 --- NO
 ICE FAR AWAY --- ANDREWSETAL71s

6922 7354 409 GSC 564
 3100 150 SH LIKELY
 16 METERS 13 METERS 16 METERS
 42 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

6937 7001 398 I 1555
 2800 140 SH DEFINITE
 8 METERS 5 METERS 8 METERS
 48 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

6937 7001 405 1 1599
 2990 140 SH DEFINITE
 9 METERS 8 METERS 9 METERS
 48 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

6951 7029 406 GSC 583
 2770 140 SH DEFINITE
 6 METERS 6 METERS 6 METERS
 61 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

7153 7820 401 GSC 654
 2780 140 SH LIKELY
 2 METERS 2 METERS 2 METERS
 --- NO
 ICE FAR AWAY --- ANDREWS 1967

7541 8439 11 S433
 2900 85 BO FAIRLY SURE
 3 METERS 3 METERS 3 METERS
 76 METERS NO
 ICE FAR AWAY --- BARR ARCTIC71s

PRINT: SAME FOR ITEMS WITH LATITUDE, FROM 5800 TO 7600
AND LONGITUDE, FROM
6000 TO 10000 AND C14 DATE, FROM 2250 TO 2750*

NO. OF ITEMS IN QUERY RESPONSE = 9
NO. OF ITEMS IN THE DATA BANK = 1044
PERCENTAGE OF RESPONSE/TOTAL DATA BANK = .86

5842 9351 196 GSC683
2320 130 SH FAIRLY SURE
27 METERS 27 METERS 27 METERS
--- NO
ICE FAR AWAY --- GSC11RC 1971

6215 7525 586 GSC 701
2670 0 --- FAIRLY SURE
13 METERS 17 METERS 13 METERS
140 METERS NO
ICE FAR AWAY --- GSCRC \$

6230 7740 585 OHITUK6
2650 250 EK LIKELY
14 METERS 18 METERS 14 METERS
140 METERS NO
ICE FAR AWAY --- ANDREWSEA71 \$

6239 6937 399 M 1535
2410 120 --- DEFINITE
10 METERS 12 METERS 10 METERS
93 METERS NO
ICE FAR AWAY --- ESK SITE ARC71

6239 6939 407 P 698
2608 50 --- DEFINITE
5 METERS 7 METERS 5 METERS
93 METERS NO
ICE FAR AWAY --- ESK SITE ARC71

6335 8200 230 P 76
2632 128 80 FAIRLY SURE
17 METERS 21 METERS 17 METERS
155 METERS NO
ICE FAR AWAY --- ANDREWSEA71 \$

6825 6648 402 GAK1992
2400 90 80 LIKELY
2 METERS 3 METERS 2 METERS
35 METERS NO
ICE NEARBY --- ANDREWS 1467

6827 6652 753 GAK1992
2400 90 80 LIKELY
1 METERS 3 METERS 1 METERS
--- NO
ICE FAR AWAY --- ANDREWSETAL71\$

7540 8437 10 13231
 2450 90 PT FAIRLY SURE
 4 METERS 4 METERS 4 METERS
 76 METERS NO
 ICE FAR AWAY --- BARR ARCTIC71\$

PRINT: SAME FOR ITEMS WITH LATITUDE. FROM 5800 TO 7600
 AND LONGITUDE, FROM
 6000 TO 10000 AND C14 DATE, FROM 1750 TO 2250*

NO. OF ITEMS IN QUERY RESPONSE = 11
 NO. OF ITEMS IN THE DATA BANK = 1044
 PERCENTAGE OF RESPONSE/TOTAL DATA BANK = 1.05

5845 9359 197 GSC723
 2120 130 SM FAIRLY SURE
 22 METERS 22 METERS 22 METERS
 --- NO
 ICE FAR AWAY --- GSC11RC 1971

6212 7310 584 NIAQUN4
 2200 100 EK LIKELY
 8 METERS 12 METERS 6 METERS
 140 METERS NO
 ICE FAR AWAY --- ANDREWSEA71 \$

6230 7740 583 EETEEV6
 1800 200 EK LIKELY
 8 METERS 12 METERS 8 METERS
 140 METERS NO
 ICE FAR AWAY --- ANDREWSEA71 \$

6239 6937 400 M 1534
 2200 120 --- DEFINITE
 6 METERS 8 METERS 6 METERS
 93 METERS NO
 ICE FAR AWAY --- ESK SITE ARC71

6320 8330 232 WALIS20
 2200 100 EK FAIRLY SURE
 12 METERS 15 METERS 12 METERS
 --- NO
 ICE FAR AWAY --- ANDREWSEA71 \$

6333 8539 88 M1085
 2225 200 OR FAIRLY SURE
 12 METERS 12 METERS 12 METERS
 127 METERS NO
 ICE FAR AWAY --- CRANE+GRIFFEN

6335 8200 231 P 77
 2191 120 BO FAIRLY SURE
 12 METERS 14 METERS 12 METERS
 155 METERS NO
 ICE FAR AWAY --- ANDREWSEA71 \$

6418 9546 312 GSC23
 1800 60 OR POSSIBLY
 9 METERS 9 METERS 9 METERS
 --- NO
 ICE FAR AWAY --- GSC1 PC 1962

6758 6517 403 GAK2771
 2090 100 OR SMALL CHANCE
 3 METERS 1 METERS 3 METERS
 30 METERS NO
 ICE FAR AWAY --- ---

6902 7502 404 I 489
 2050 170 SH FAIRLY SURE
 9 METERS 9 METERS 9 METERS
 83 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

6906 7447 392 I 1830
 1950 100 SH DEFINITE
 6 METERS 2 METERS 6 METERS
 83 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

PRINT: SAME FOR ITEMS WITH LATITUDE, FROM 5800 TO 7600
 AND LONGITUDE, FROM
 6000 TO 10000 AND C14 DATE, FROM 1250 TO 1750*

NO. OF ITEMS IN QUERY RESPONSE = 5
 NO. OF ITEMS IN THE DATA BANK = 1044
 PERCENTAGE OF RESPONSE/TOTAL DATA BANK = .48

6210 7548 588 GSC 537
 1600 140 PT LIKELY
 6 METERS 8 METERS 6 METERS
 140 METERS NO
 ICE FAR AWAY --- MATTHEWS67

6237 6932 394 M 1529
 1470 110 --- DEFINITE
 2 METERS 4 METERS 2 METERS
 93 METERS NO
 ICE FAR AWAY --- ESK SITE ARC71

6258 6940 393 M 1533
 1670 150 --- DEFINITE
 2 METERS 4 METERS 2 METERS
 93 METERS NO
 ICE FAR AWAY --- ESK SITE ARC71

6340 7710 764 INFER
 1300 300 CU POSSIBLY
 15 METERS 18 METERS 15 METERS
 --- NO
 ICE FAR AWAY --- ANDREWS ET AL 71s

7242 7750 395 S 47F
 1380 95 --- DEFINITE
 1 METERS 3 METERS 1 METERS
 --- NO
 ICE FAR AWAY --- ESK SITE ARC71

PRINT: SAME FOR ITEMS WITH LATITUDE, FROM 5800 TO 7600
 AND LONGITUDE, FROM
 6000 TO 10000 AND C14 DATE, FROM 750 TO 1250*

NO. OF ITEMS IN QUERY RESPONSE = 17
 NO. OF ITEMS IN THE DATA BANK = 1644
 PERCENTAGE OF RESPONSE/TOTAL DATA BANK = 1.63

5845 9350 198 GSC682
 1240 130 SH FAIRLY SURE
 10 METERS 10 METERS 10 METERS
 --- NO
 ICE FAR AWAY --- GSC11RC 1971

5846 9357 199 GSC684
 1020 140 SH FAIRLY SURE
 6 METERS 6 METERS 6 METERS
 --- NO
 ICE FAR AWAY --- GSC11RC 1971

5950 8000 217 1 2418
 1150 150 SH FAIRLY SURE
 5 METERS 5 METERS 5 METERS
 155 METERS NO
 ICE FAR AWAY --- ANDREWS67

5950 8000 358 12418
 1150 100 SH FAIRLY SURE
 5 METERS 5 METERS 5 METERS
 --- NO
 ICE FAR AWAY --- ANDREWS GB1967

6515 8500 233 SITE 21
 800 0 EK FAIRLY SURE
 4 METERS 6 METERS 4 METERS
 --- NO
 ICE FAR AWAY --- ANDREWSEA71 \$

6667 8645 765 INFER
 900 100 CU LIKELY
 9 METERS 12 METERS 9 METERS
 --- NO
 ICE FAR AWAY --- ANDREWSETAL71\$

6713 6340 390 GAK3096
 930 100 OR DEFINITE
 0 METERS 0 METERS 0 METERS
 4 METERS NO
 ICE FAR AWAY --- ---

6735 6518 389 GAK3639
 750 140 OR DEFINITE
 0 METERS 0 METERS 0 METERS
 17 METERS NO
 ICE FAR AWAY --- ---

6845 8130 359 GSC691
 1020 130 SH FAIRLY SURE
 7 METERS 7 METERS 7 METERS
 --- NO
 ICE FAR AWAY --- ANDREWS GB1967

6845 8130 396 GSC 691
 1020 130 SH DEFINITE
 7 METERS 7 METERS 7 METERS
 134 METERS NO
 ICE FAR AWAY --- ANDREWS 1967

6847 8114 143 GSC691
 1020 130 SH LIKELY
 7 METERS 7 METERS 7 METERS
 --- NO
 ICE FAR AWAY --- GSC7 RC 1968

6910 8359 397 ---
 1100 0 --- FAIRLY SURE
 5 METERS 8 METERS 5 METERS
 134 METERS NO
 ICE FAR AWAY --- ESK SITE ARC71

6920 8220 766 INFER
 1100 0 CU LIKELY
 5 METERS 8 METERS 5 METERS
 --- NO
 ICE FAR AWAY --- ANDREWSETAL715

7055 8911 79 GSC239
 940 130 OR LIKELY
 3 METERS 3 METERS 3 METERS
 --- NO
 ICE FAR AWAY --- GSC4 RC 1965

7055 8911 391 GSC 239
 940 130 WO FAIRLY SURE
 3 METERS 3 METERS 3 METERS
 --- NO
 ICE FAR AWAY --- GSCRCIVCRAIG

7440 9510 762 INFER
 1200 400 CU LIKELY
 3 METERS 6 METERS 3 METERS
 --- NO
 ICE FAR AWAY --- ANDREWSETAL715

7540 9910 763 GSC148
 1150 160 --- LIKELY
 2 METERS 2 METERS 0 METERS
 --- NO
 ICE FAR AWAY --- ANDREWSETAL71S

PRINT: SAME FOR ITEMS WITH LATITUDE, FROM 5800 TO 7600
 AND LONGITUDE, FROM
 6000 TO 10000 AND C14 DATE, FROM 250 TO 750*

NO. OF ITEMS IN QUERY RESPONSE = 4
 NO. OF ITEMS IN THE DATA BANK = 1044
 PERCENTAGE OF RESPONSE/TOTAL DATA BANK = .38

5900 9353 64 GX1073
 385 80 SH FAIRLY SURE
 4 METERS 4 METERS 4 METERS
 150 METERS YES
 ICE FAR AWAY --- WAGNER 67 S
 6206 7433 372 Y1717
 330 100 OR FAIRLY SURE
 3 METERS 3 METERS 3 METERS
 --- NO
 ICE FAR AWAY --- YALE 1X RC1969
 6735 6518 389 GAK3639
 750 140 OR DEFINITE
 0 METERS 0 METERS 0 METERS
 17 METERS NO
 ICE FAR AWAY --- ---
 6910 8359 767 K 504
 600 150 BO FAIRLY SURE
 4 METERS 8 METERS 4 METERS
 137 METERS NO
 ICE FAR AWAY --- FARRAND62ISOTI

END

TOTAL RUN TIME IN SECONDS: 31.415

Occasional Papers

INSTITUTE OF ARCTIC AND ALPINE RESEARCH

- Occasional Paper No. 1: The Taxir Primer, R.C. Brill, 1971.
- Occasional Paper No. 2: Present and Paleo-Climatic Influences on the Glacierization and Deglaciation of Cumberland Peninsula, Baffin Island, J. T. Andrews and R. G. Barry, and others, 1972.
- Occasional Paper No. 3: Climatic Environment of the East Slope of the Colorado Front Range, R. G. Barry, 1972.
- Occasional Paper No. 4: Short-Term Air-Sea Interactions and Surface Effects in the Baffin Bay—Davis Strait Region from Satellite Observations, J. D. Jacobs, R. G. Barry, B. Stankov, and J. Williams, 1972.
- Occasional Paper No. 5: Simulation of the Climate at the Last Glacial Maximum Using the NCAR Global Circulation Model, Jill Williams, R. G. Barry, and W. M. Washington, 1973.
- Occasional Paper No. 6: Guide to the Mooses of Colorado, William A. Weber, 1973.
- Occasional Paper No. 7: A Climatological Study of Strong Downslope Winds in the Boulder Area, W. A. R. Brinkmann, 1973.
- Occasional Paper No. 8: Environmental Inventory and Land Use Recommendations for Boulder County, Colorado, Richard F. Madole, 1973.
- Occasional Paper No. 9: Studies of Climate and Ice Conditions in Eastern Baffin Island, 1971-73, J. D. Jacobs, R. G. Barry, R. S. Bradley, and R. L. Weaver.
- Occasional Paper No. 10: Simulation of the Atmospheric Circulation Using the NCAR Global Circulation Model with Present-Day and Glacial Period Boundary Conditions, Jill Williams, 1974.
- Occasional Paper No. 11: Solar and Atmospheric Radiation Data for Broughton Island, Eastern Baffin Island, Canada, 1971-1973, John D. Jacobs, 1974.
- Occasional Paper No. 12: Deglacial Chronology and Uplift History: Northeastern Sector, Laurentide Ice Sheet, Arthur S. Dyke, 1974.

